

# Purely Data-Driven Sub-Weekly Mass Change Models from GRACE Residuals

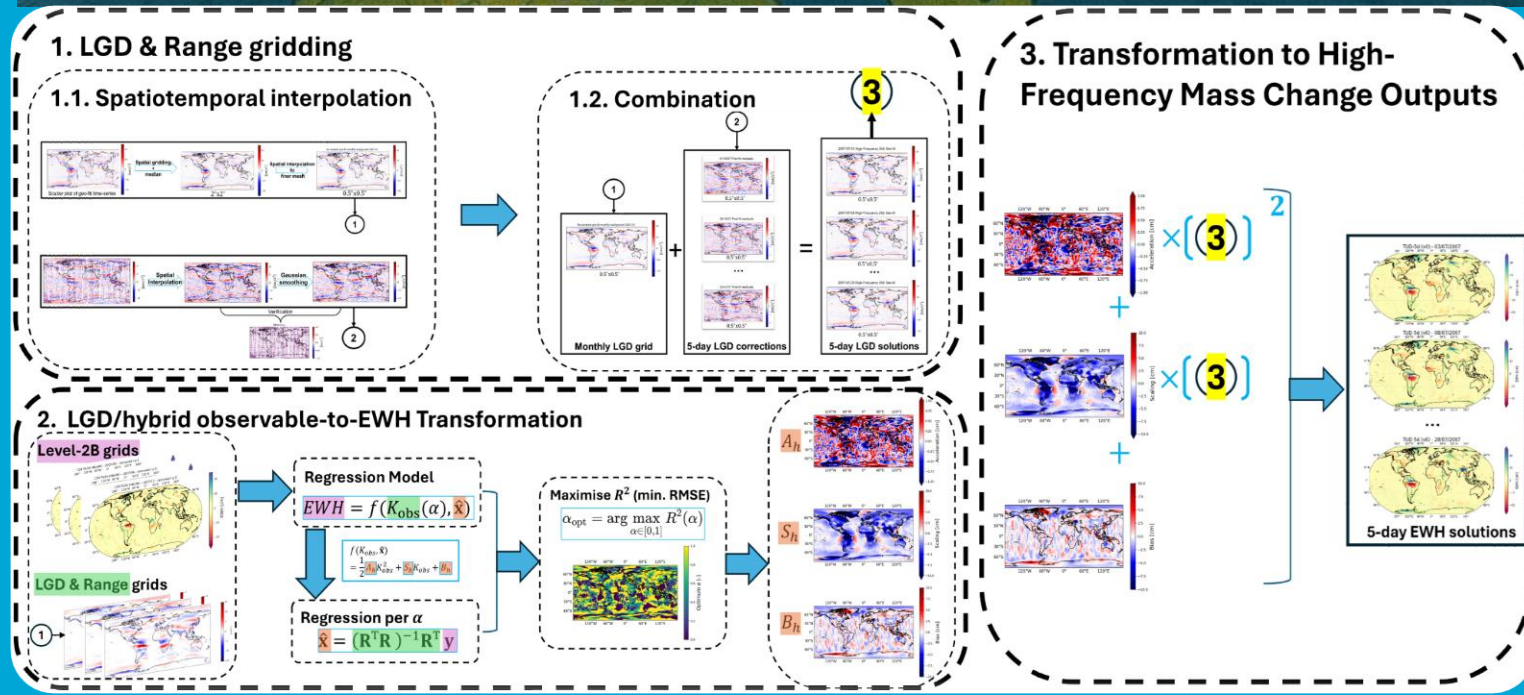
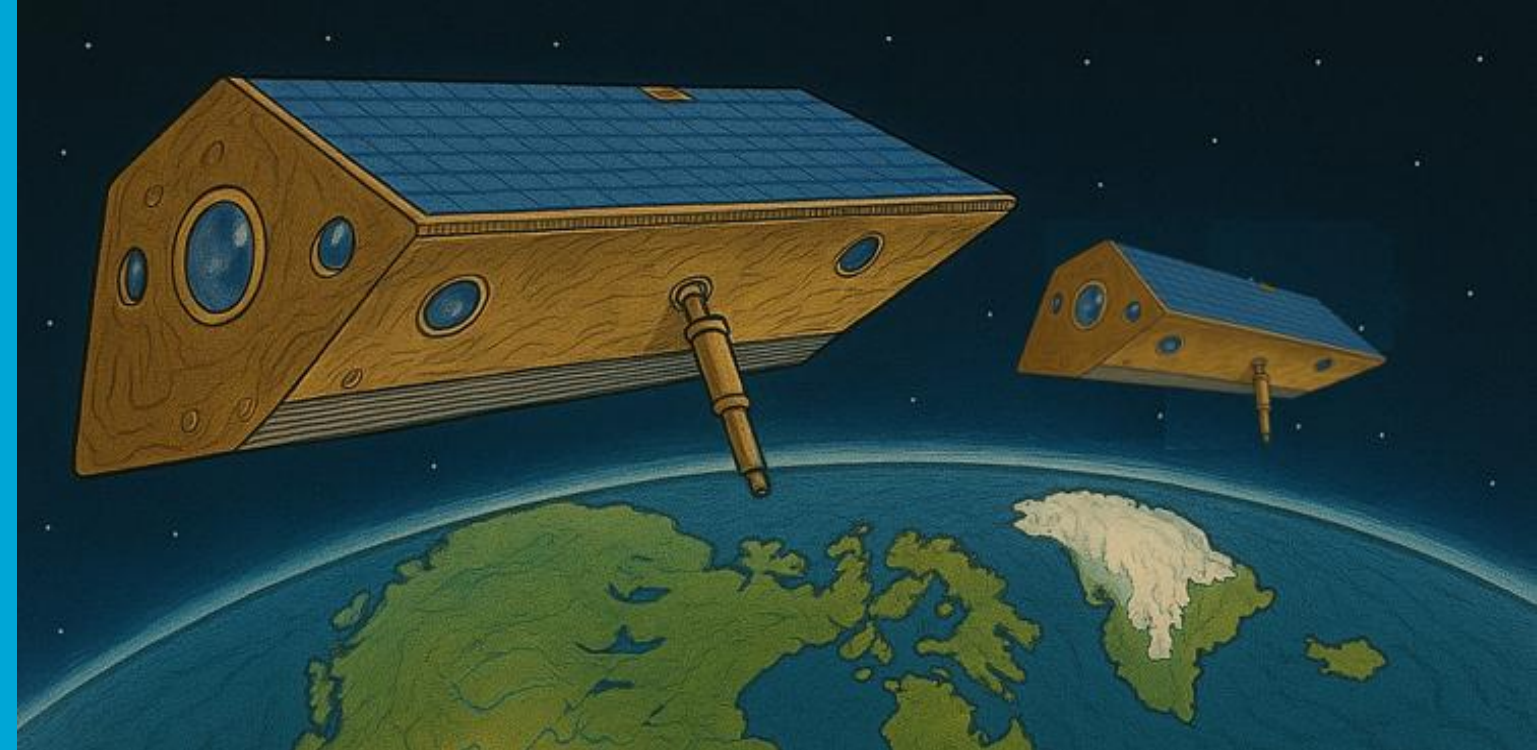
M. Cuadrat-Grzybowski

Joao Teixeira da Encarnacao

GRACE-FO 2025 Science Team Meeting



JWG C10. Tailored  
Parameterization  
Strategies for  
Climate Applications of  
Satellite Gravimetry



# Introduction

## Overall research goal

Design high-frequency mass change models and constrain them with GRACE residual Level-1B data.

## Current goals

1. Construct *GRACE-only* sub-weekly mass change models using *purely* post-fit range-rate data.
2. Estimate associated uncertainties.

## Presentation outlook

- Along-orbit residual Line-of-sight Gravity Difference
- Novel High-Frequency Methodology
- 5-day solutions results
- Preliminary uncertainty estimations

$$\hat{l} = A\hat{x}$$

$\hat{l}$ : post-fit residual range rate,  
 $A$ : design matrix,  
 $\hat{x}$ : estimated parameter vector.

# LGDs from L1B post-fit data

## Residual Level-1B (post-fit) data (RL06 CSR\*):

- Geo-fit range-rate:  
→ part of the post-fit range rate **only** containing the monthly gravity signal.
- Post-fit residuals:  $\delta l = l - A\hat{x}$ .  
→ Contain sub-monthly gravity signals (w.r.t monthly mean) and background model errors.

## Obtaining residual Line-of-sight-Gravity Difference (LGD):

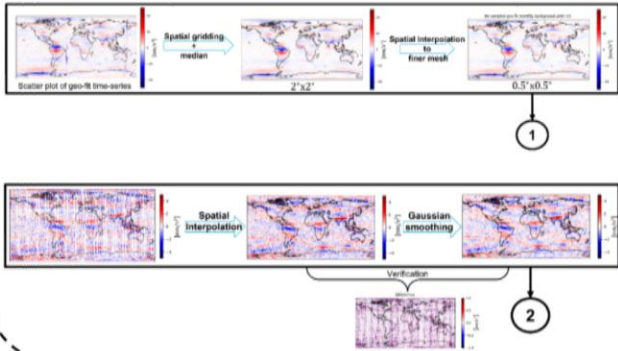
- Step 1: apply numerical differentiation → range-accelerations  $\delta\ddot{\rho}$ ,
- Step 2: apply band-pass filter to remove KBR noise and long wavelengths [0.9, 11] mHz,
- Step 3: apply frequency-based Transfer Function by [Ghobadi-Far, K. et al 2018](#), denoted as  $Z(f)$ , to remove residual centrifugal acceleration.

$$\delta g_{AB}^{\text{LoS}}(t) \approx F^{-1}\{Z(f)F\{\delta\ddot{\rho}(t)\}\}$$

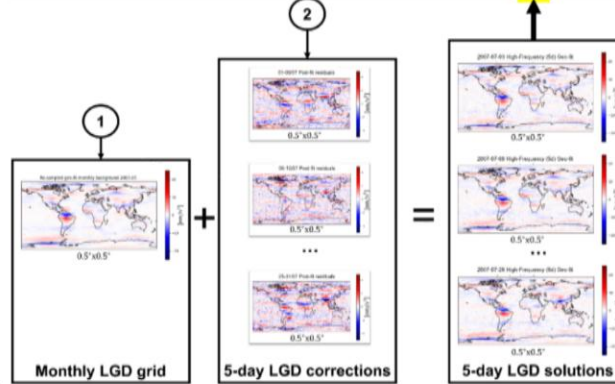


## 1. LGD & Range gridding

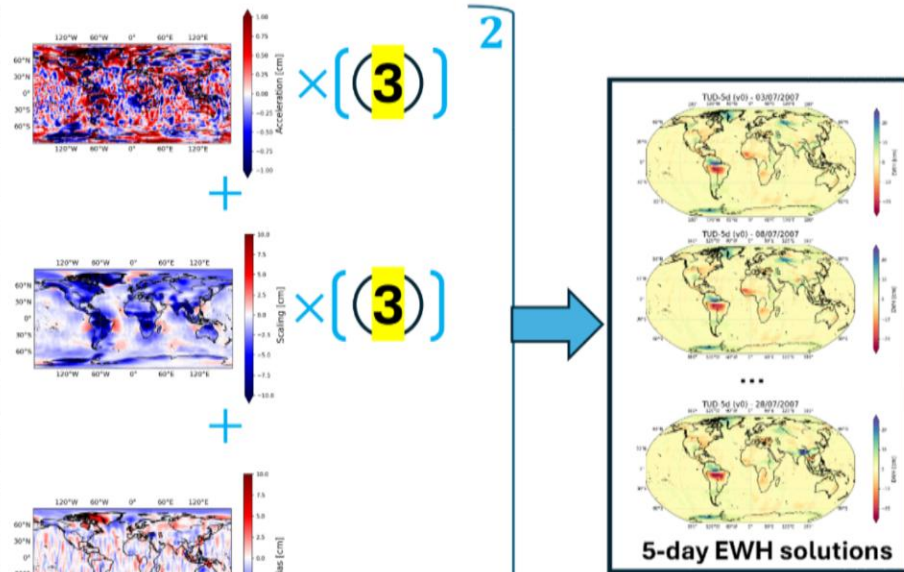
### 1.1. Spatiotemporal interpolation



### 1.2. Combination

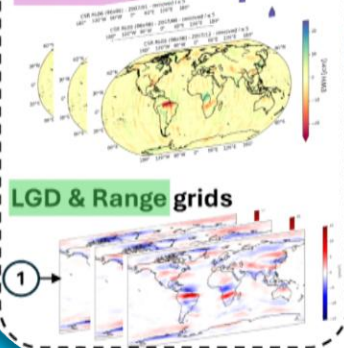


## 3. Transformation to High-Frequency Mass Change Outputs



## 2. LGD/hybrid observable-to-EWH Transformation

### Level-2B grids



### Regression Model

$$EWH = f(K_{obs}(\alpha), \hat{x})$$

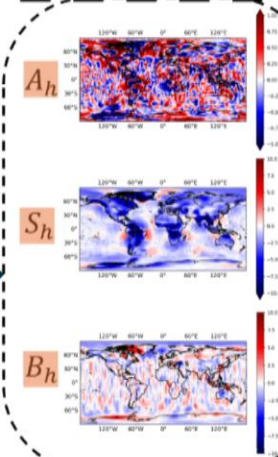
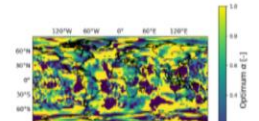
$$f(K_{obs}; \hat{x}) = \frac{1}{2} \hat{x}^T K_{obs}^2 \hat{x} + S_{\hat{x}}^T K_{obs} \hat{x} + B_{\hat{x}}$$

### Regression per $\alpha$

$$\hat{x} = (R^T R)^{-1} R^T y$$

### Maximise $R^2$ (min. RMSE)

$$\alpha_{opt} = \arg \max_{\alpha \in [0,1]} R^2(\alpha)$$



Novel Framework (core method)

# Uncertainty estimation

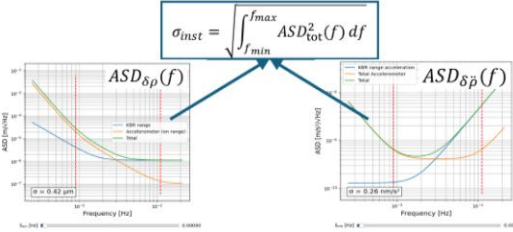
$$\sigma_{EWH}^2 = \left( \frac{\partial EWH}{\partial K_{obs}} \right)^2 \sigma_{K_{obs}}^2 + p^T \Sigma_{\hat{x}} p + \sigma_{EWH,L2}^2$$

Empirical model

$$EWH = \frac{1}{2} A K_{obs}^2 + S K_{obs} + B$$

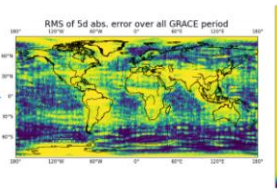
## 1. LGD/hybrid observable uncertainty

Instrument error (analytical spectra)



Interpolation error (5-day tracks)

Simulate 5-day observables



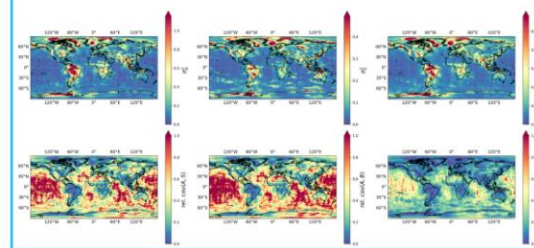
## 2. Regression uncertainty

Estimates

$$\hat{x} = \begin{bmatrix} \hat{A} \\ \hat{S} \\ \hat{B} \end{bmatrix} = (R^T R)^{-1} R y$$

Estimate Covariance matrix

$$\Sigma_{\hat{x}} \approx \hat{\sigma}^2 (R^T R)^{-1}$$

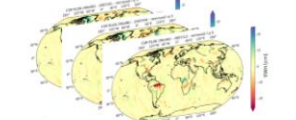


Compute Model Sensitivities

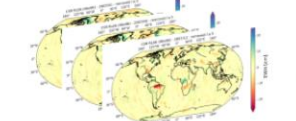
$$p = \begin{bmatrix} \partial EWH / \partial A \\ \partial EWH / \partial S \\ \partial EWH / \partial B \end{bmatrix}$$

## 3. Level-2 EWH uncertainty

CSR



GFZ



...

ITSG

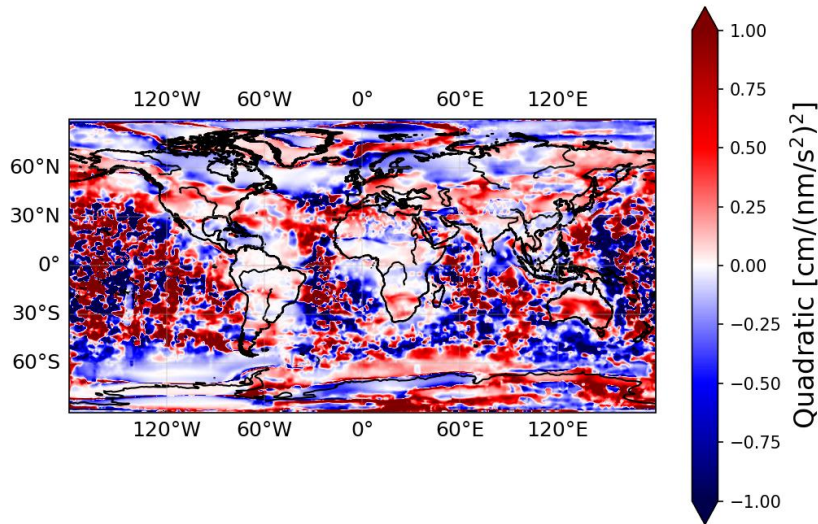


Spread between different analysis centers/solutions

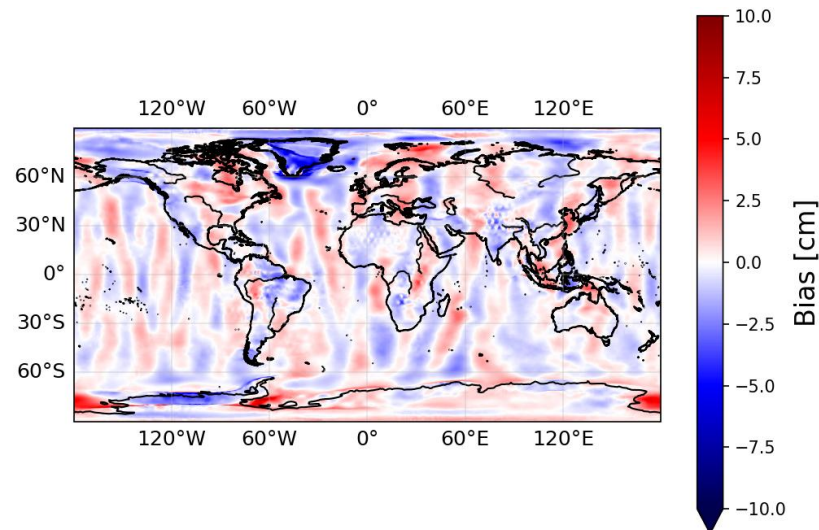
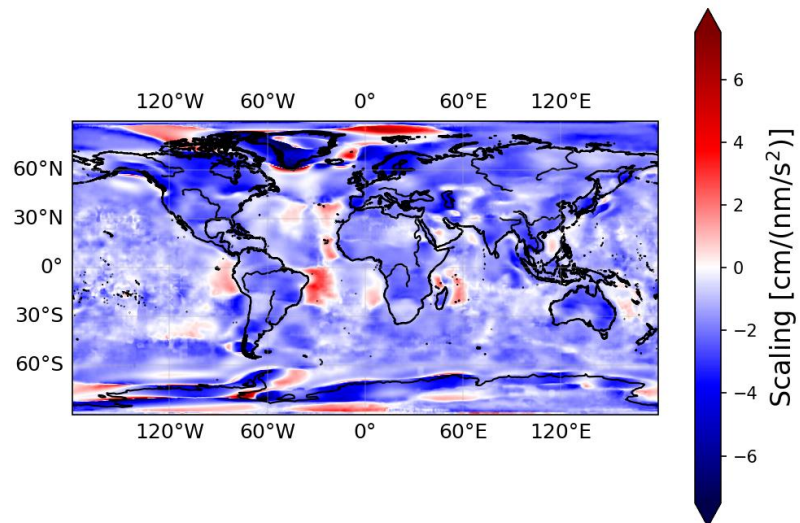
# Novel Framework (uncertainties)



# Estimated parameters



$$EWH \approx \frac{1}{2} \mathbf{A} LGD^2 + \mathbf{S} LGD + \mathbf{B}$$

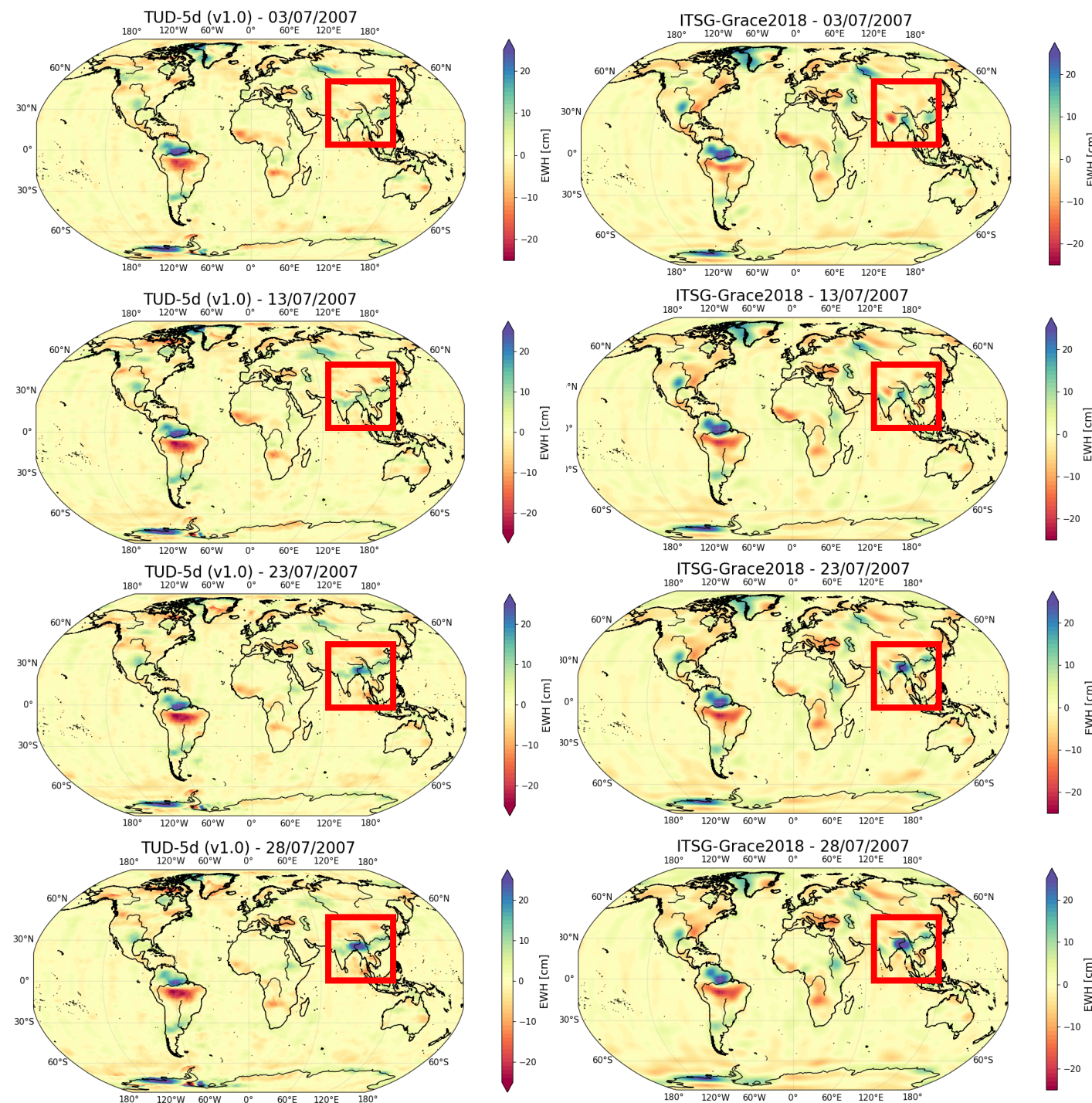


Period: 2003-2016.

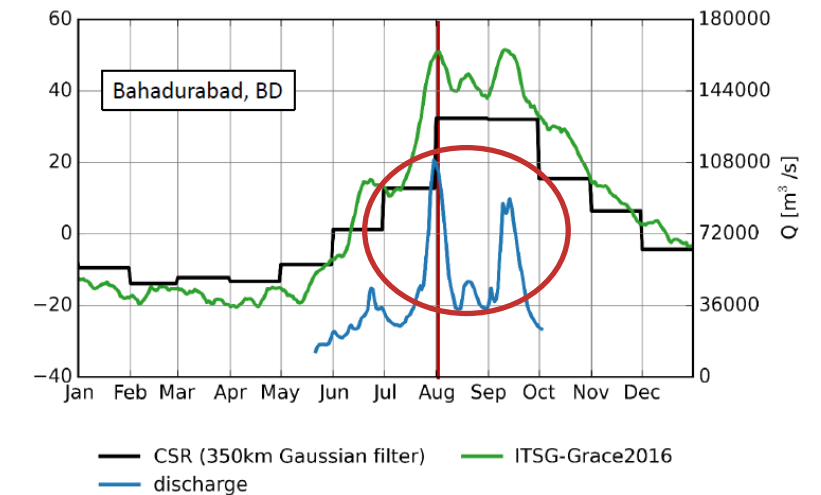
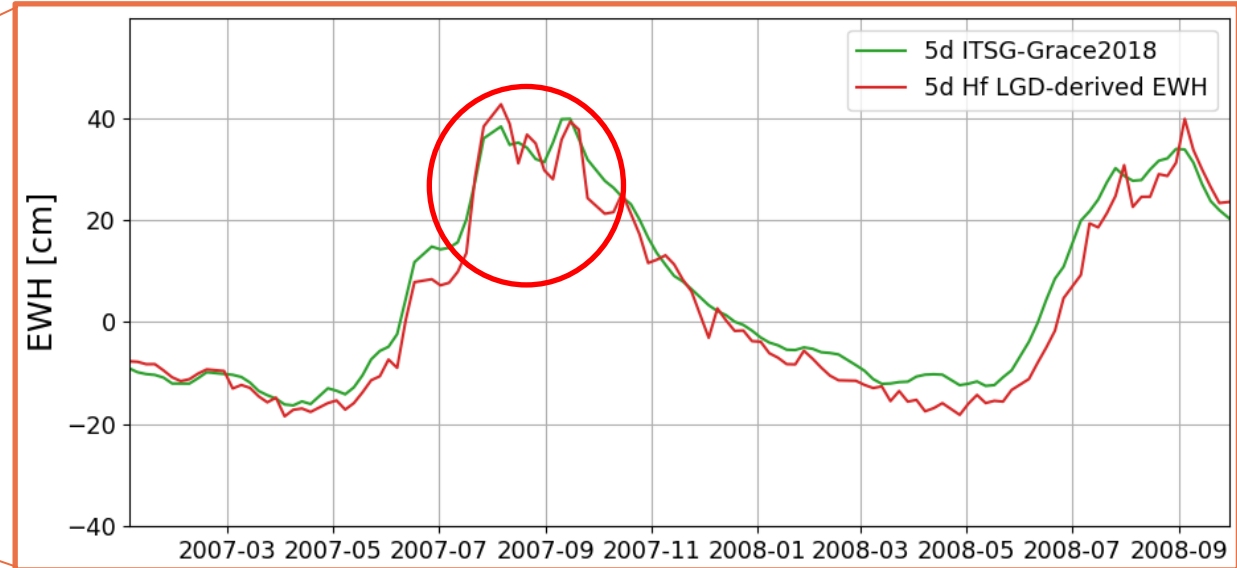
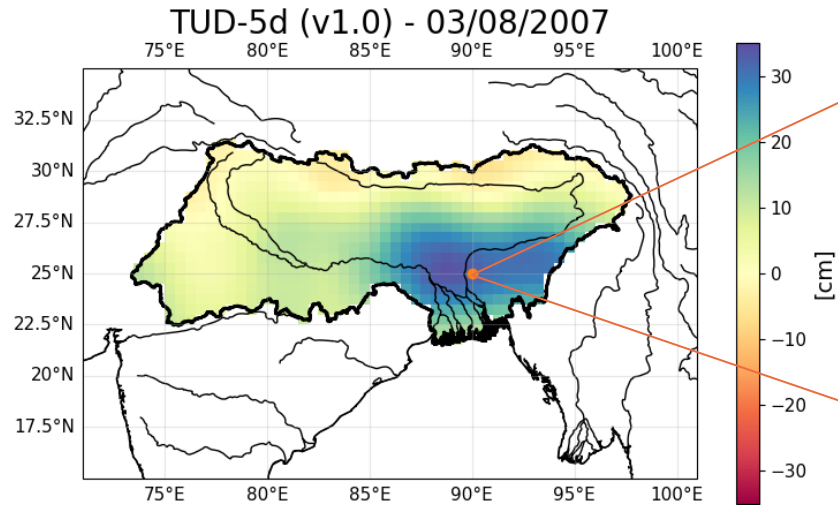
- Quadratic factor shows high variability in oceans, and very low (near zero values) in major river basins.
- Scaling factor shows overall negative values (expected negative LGD peak due to positive mass anomaly).
- Bias shows the mean difference between LGDs and EWH:
  - i. mean N-S stripe noise,
  - ii. Larger difference in polar regions.

# ITSG vs. TUD v1.0

- Evident hydrological sub-weekly signals observed in highly hydrologically active river basins.
- Differences:
  - Higher frequency content up to degree 57.
  - monthly model set as background,
  - background model errors in polar regions.



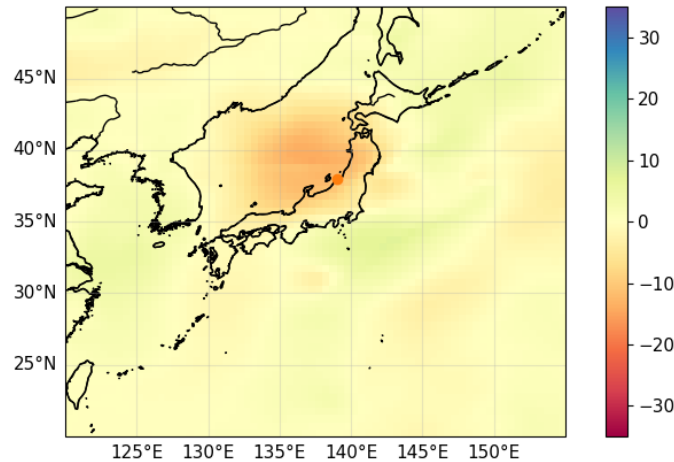
# Bangladesh 2007 floods



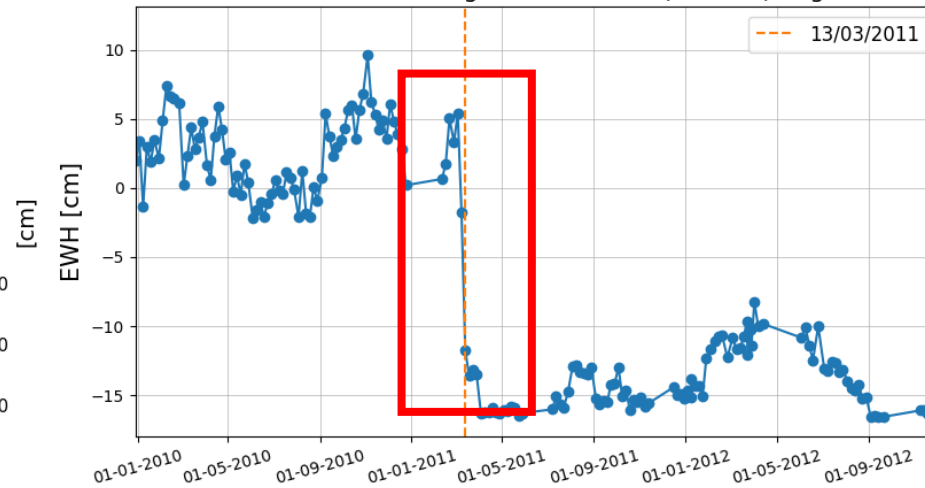


# Co-seismic sub-monthly signals

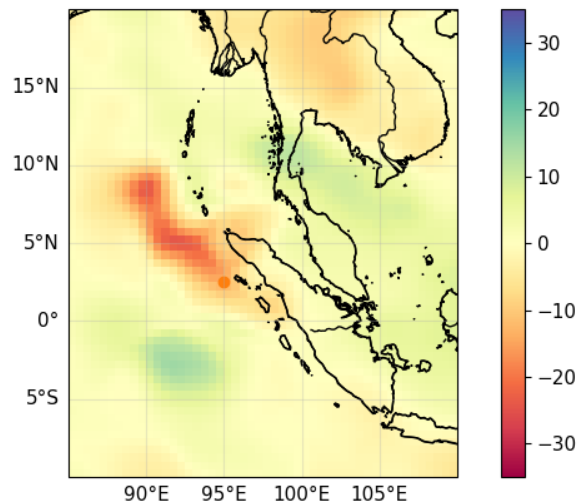
TUD-5d (v1.0) - 13/03/2011  
125°E 130°E 135°E 140°E 145°E 150°E



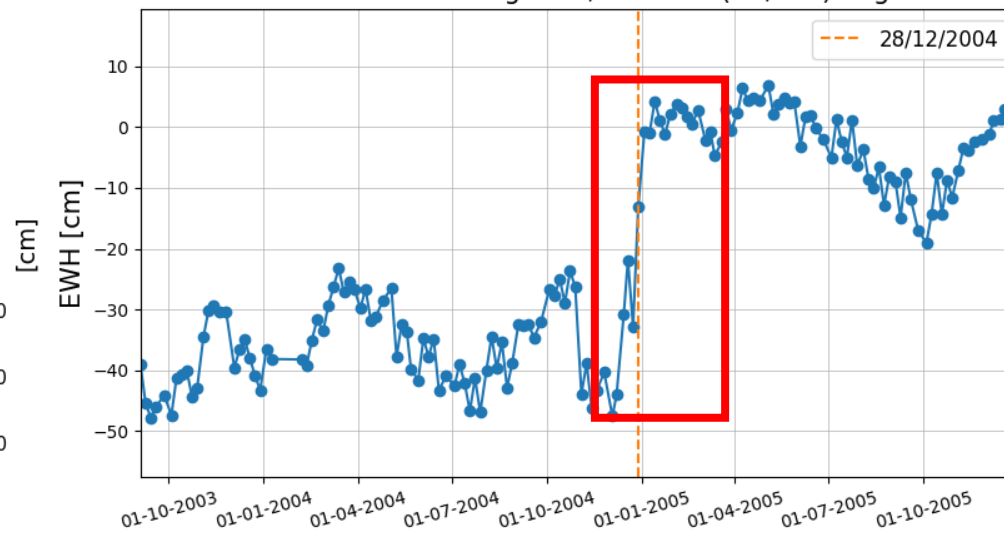
Time-series for longitude, latitude: (139, 38) deg.



TUD-5d (v1.0) - 28/12/2004  
90°E 95°E 100°E 105°E



Time-series for longitude, latitude: (95, 2.5) deg.



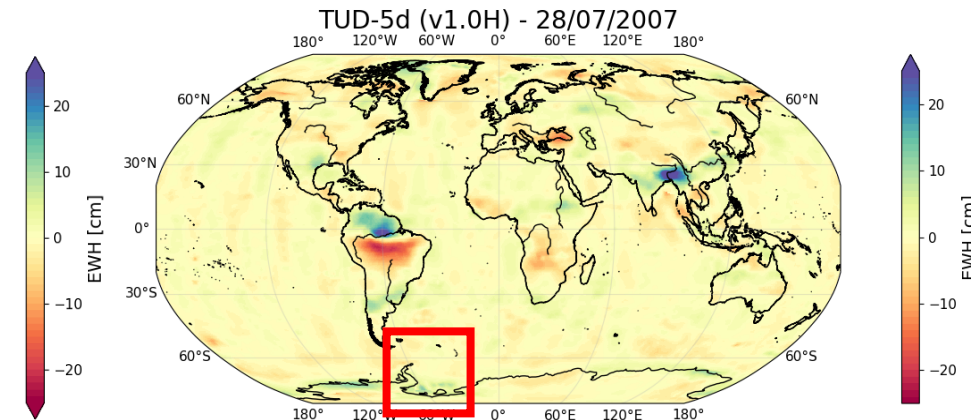
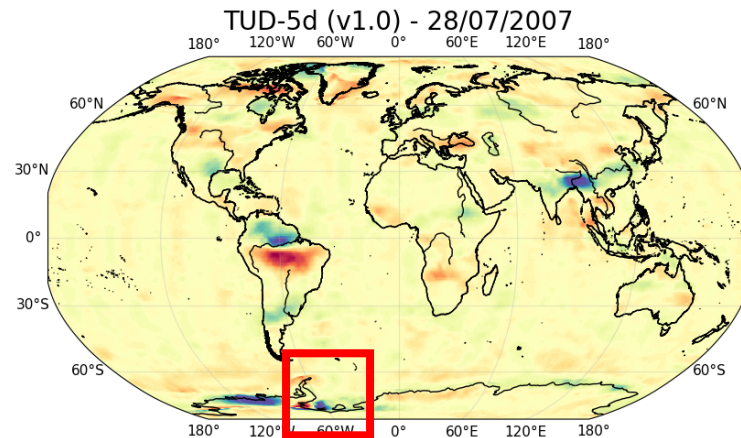
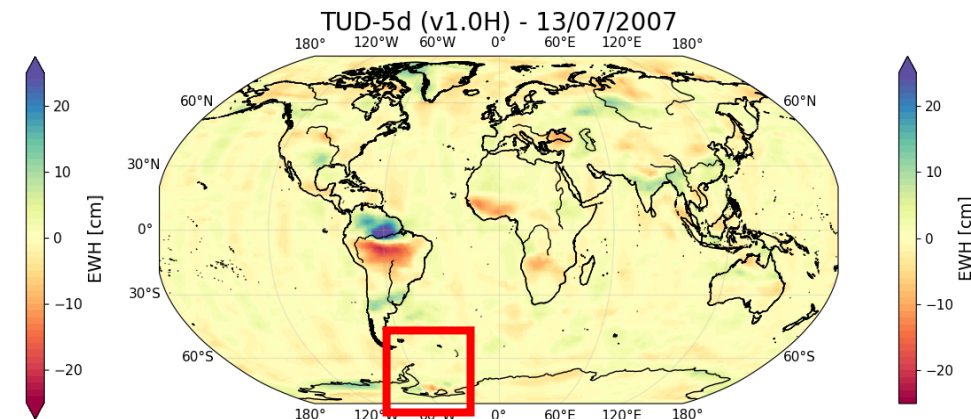
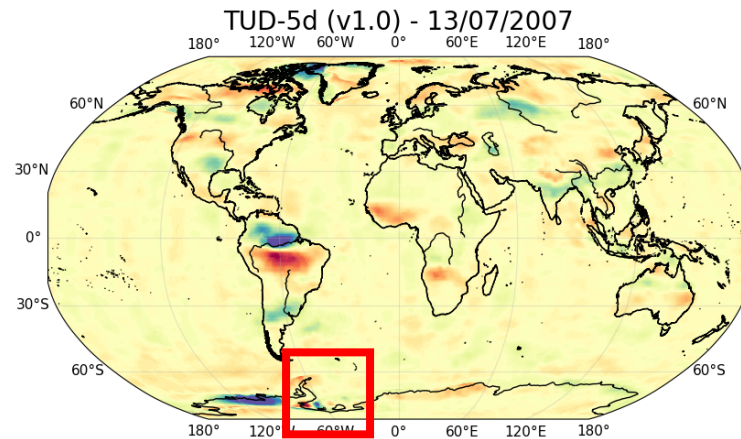
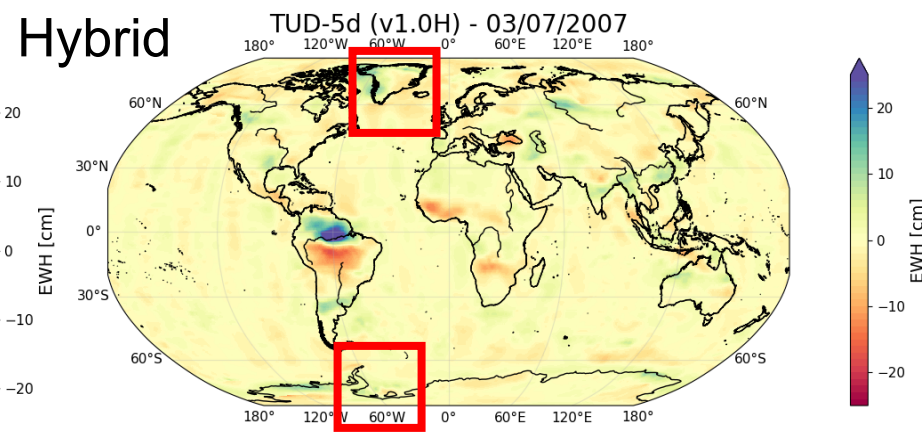
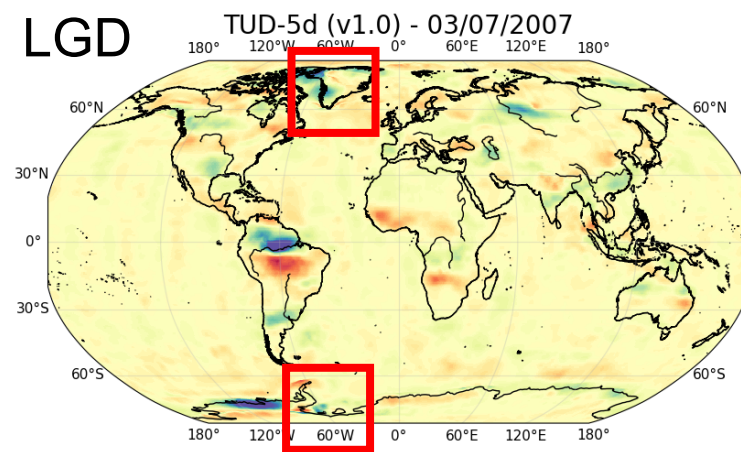
- Earthquake events:
  - i. Sumatra 2004,
  - ii. Tohoku 2011.
- Tohoku 2011  
Clear co-seismic sub-monthly signal in high-frequency models.
- Sumatra 2004  
Less clear as the event happened near the beginning of January 2005.

# Comparison of LGD-only vs Hybrid

## Observations:

1. Spatial patterns similar to  $EWH^{L2}$ ,
2. Dampening observed:  
 $|\widehat{EWH}^{hyb}| \leq |\widehat{EWH}^{LGD}|$

→ Polar region and other lobe artifacts are dampened!



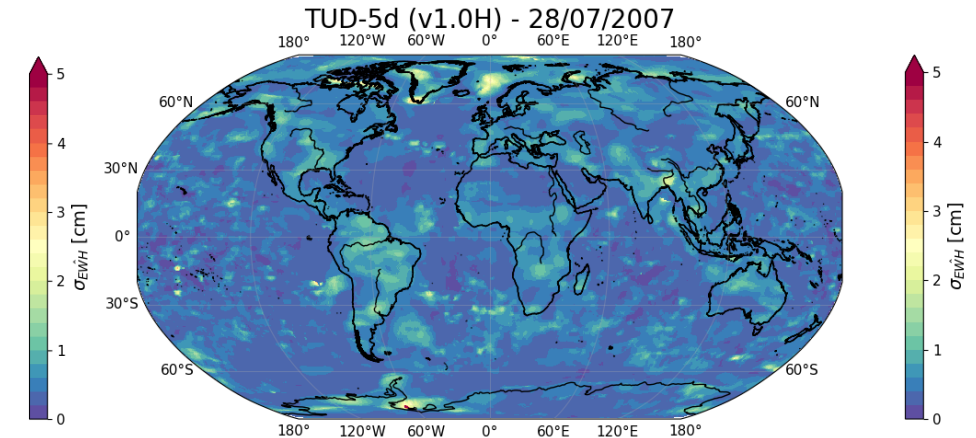
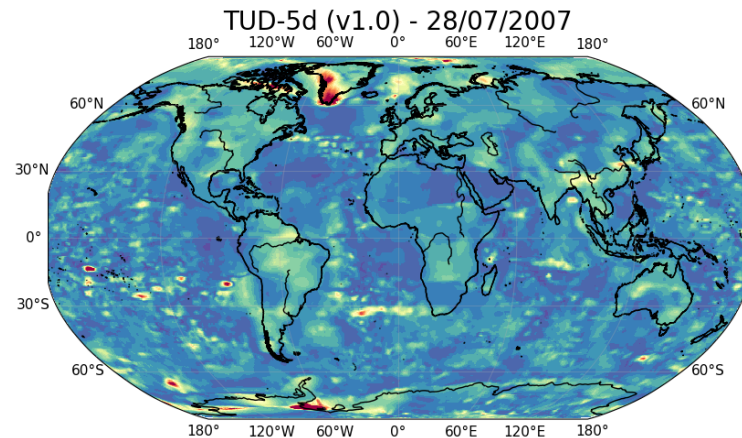
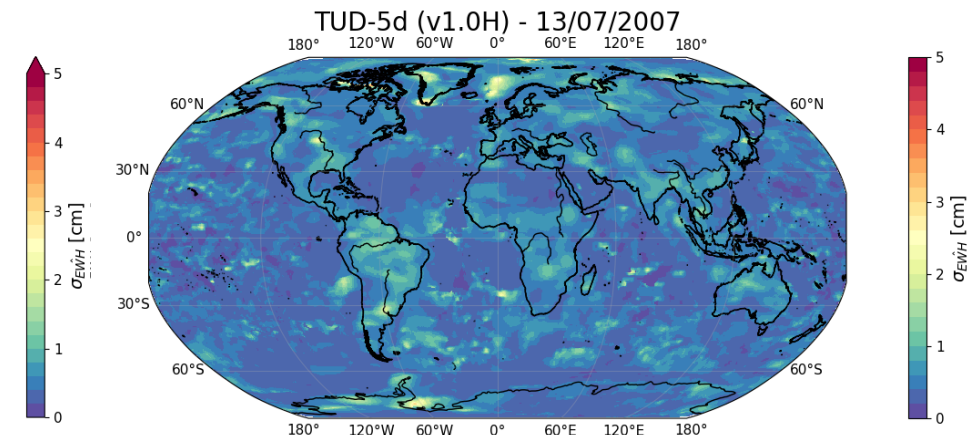
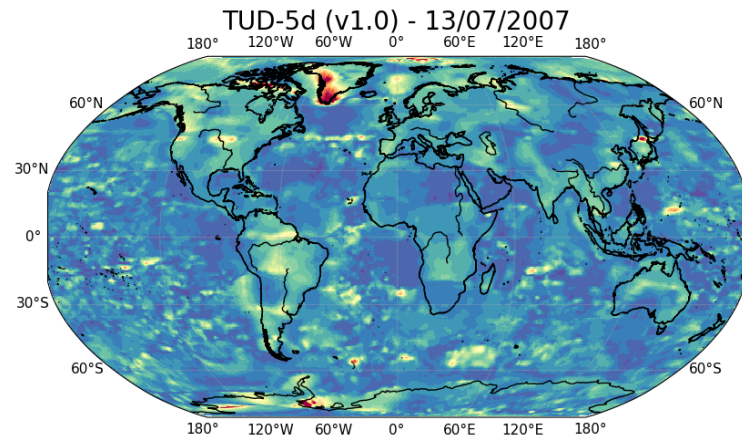
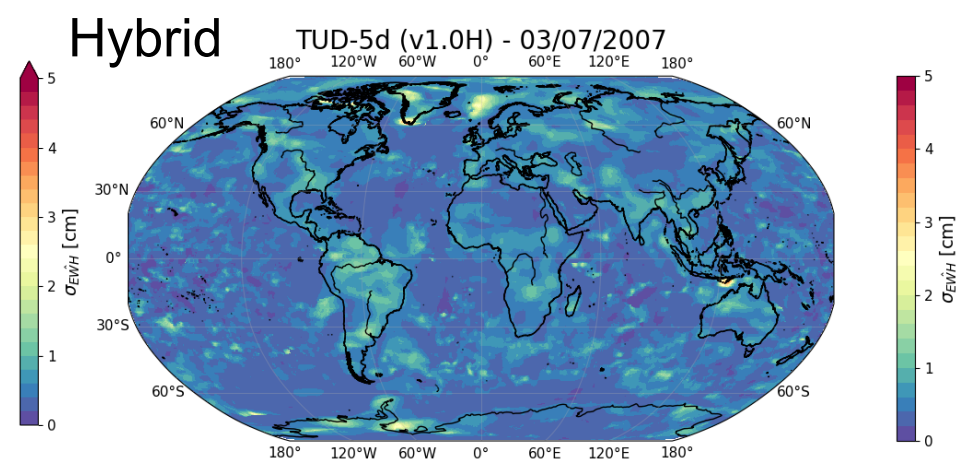
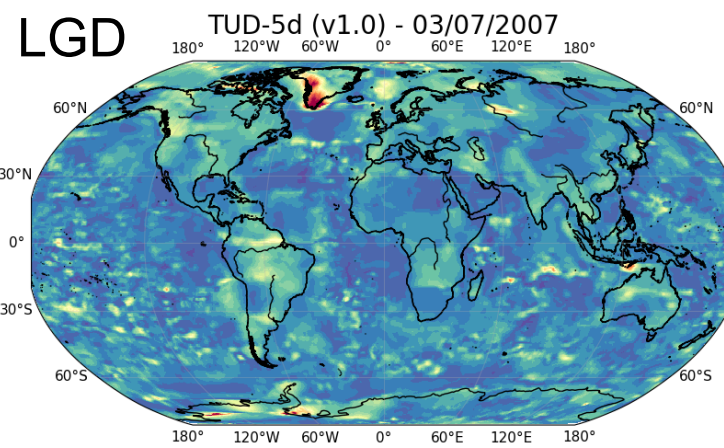


# Uncertainty estimates (LGD vs Hybrid)

## Observations:

1. Hybrid formulation unc. is approx. half of the LGD-only.
2. Hybrid unc. < 2.5 cm.
3. LGD-formulation → overestimates.

\* Currently only regression uncertainties, transfer function and instrument errors are included in the model uncertainties.





# Summary and looking forward

## Goal

Construct *GRACE-only* sub-weekly mass change models using purely post-fit range-rate data.

## Novel Methodology (purely data-driven)

- Usage of along-track res. KBR LGDs and range.
- Derivation of an ***empirical and spatially distributed model*** between LGDs into EWH.
- Two different model formulations (LGD-only vs. KBR hybrid form).
- Error estimation based on analytical noise models and regression errors.

## Comparison with ITSG-Grace2018 daily solutions

## Further work: towards a complete version.

Finalising error estimation:

- spatial interpolation error,
- L2 product errors.

# References

Mayer-Gürr, T., Behzadpour, S., Kvas, A., Ellmer, M., Klinger, B., Strasser, S., & Zehentner, N. (2018). ITSG-Grace2018: Monthly, Daily and Static Gravity Field Solutions from GRACE. Data set/Database. <https://doi.org/10.5880/ICGEM.2018.003>

Ghobadi-Far, K., Han, S.-C., Weller, S., Loomis, B. D., Luthcke, S. B., Mayer-Gürr, T., & Behzadpour, S. (2018). A transfer function between line-of-sight gravity difference and GRACE intersatellite ranging data and an application to hydrological surface mass variation. *Journal of Geophysical Research: Solid Earth*, 123, 9186–9201. <https://doi.org/10.1029/2018JB016088>

Kling, H., Fuchs, M., and Paulin, M.: Runoff conditions in the upper Danube basin under an ensemble of climate change scenarios, *J. Hydrol.*, 424–425, 264–277, <https://doi.org/10.1016/j.jhydrol.2012.01.011>, 2012.

Cuadrat-Grzybowski, M., Encarnacao, J.G.T.D., Visser, P.N.A.M. (2025). Potential of Ka-band Range Rate Post-fit Residuals for High-frequency Mass Change Applications. arXiv preprint arXiv:2503.04227. (**already submitted** to *JGR: Solid Earth* for review).