

Ocean-atmosphere Fresh Water Flux in Global Hydrological Balance

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- **Two Way to Estimate**
- **Validation with Rawinsonde**
- **Atmospheric Water Balance**
- **Ocean Mass Balance**
- **Continental Mass Balance**
- **Sea Level Balance**
- **Salinity Balance**
- **Meridional Transport**

Importance of Water Cycle

- **Critical to existence of human life**
- **Essential to weather and climate**
- **Tightly coupled with energy cycle**
- **Heat storage with high heat capacity and latent heat form clouds and affect radiative balance**

HYDROLOGIC BALANCE

$$\frac{\partial W}{\partial t} + \nabla \cdot \Theta = E - P$$

$$\Theta = \frac{1}{g} \int_0^{p_0} q U dp$$

$$W = \frac{1}{g} \int_0^{p_0} q dp$$

$$\Theta = U_e W$$

Θ is equivalent to column water vapor W advected by U_e .

U_e is the depth-averaged wind weighted by humidity

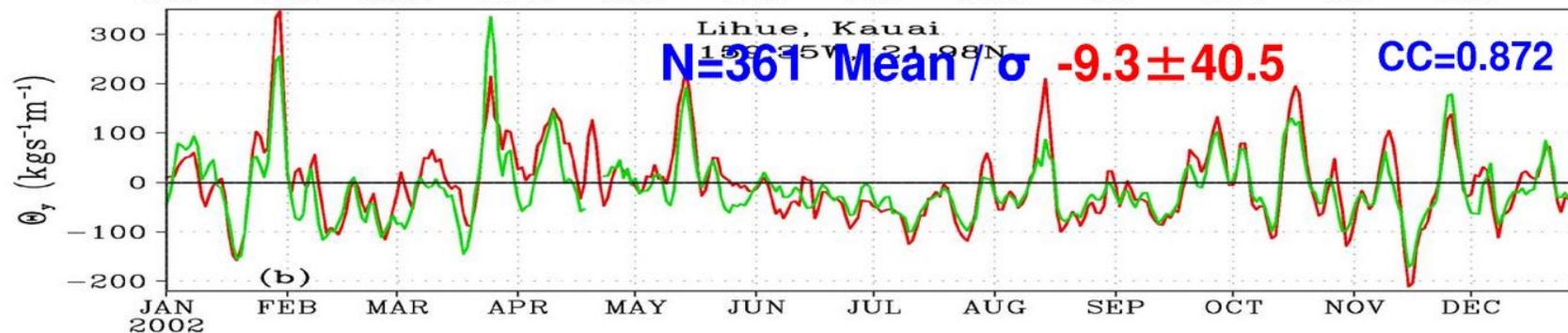
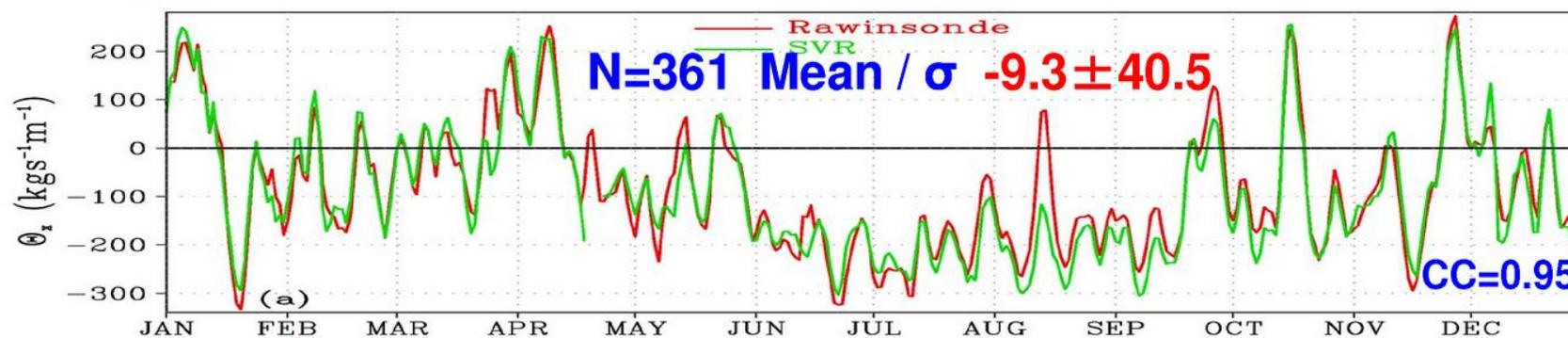
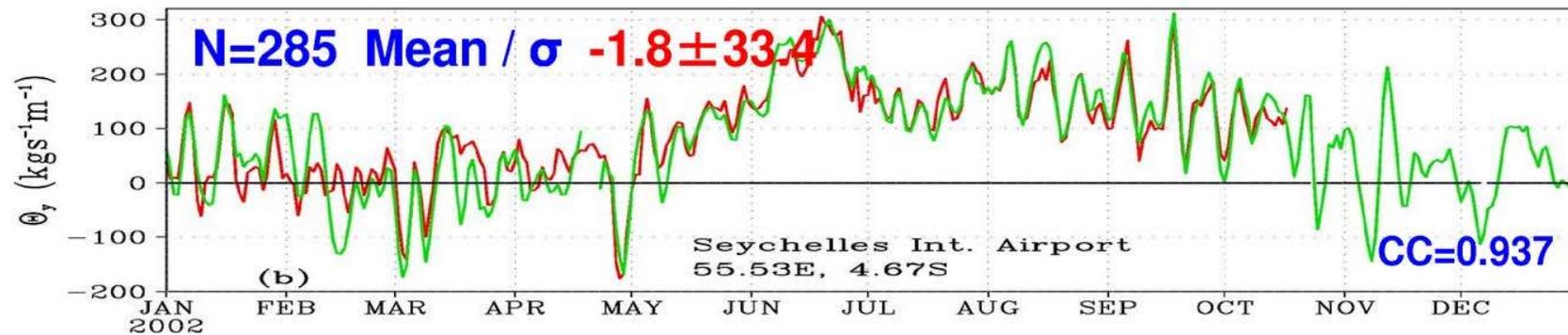
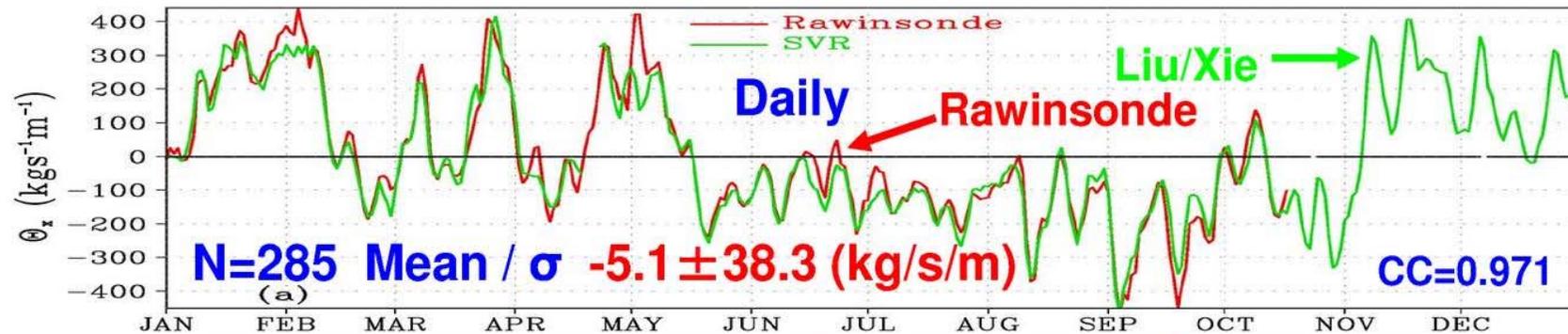
We use SVR to relate U_e to wind at two levels:

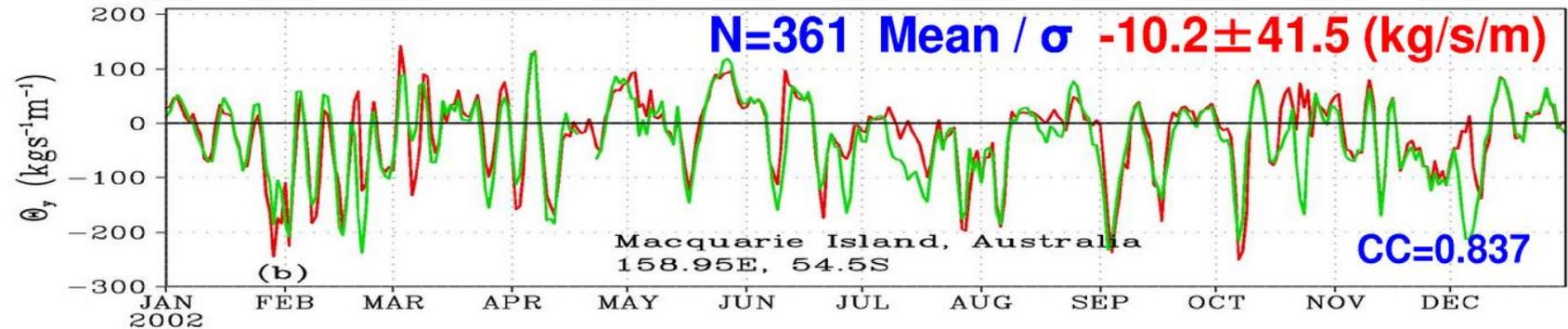
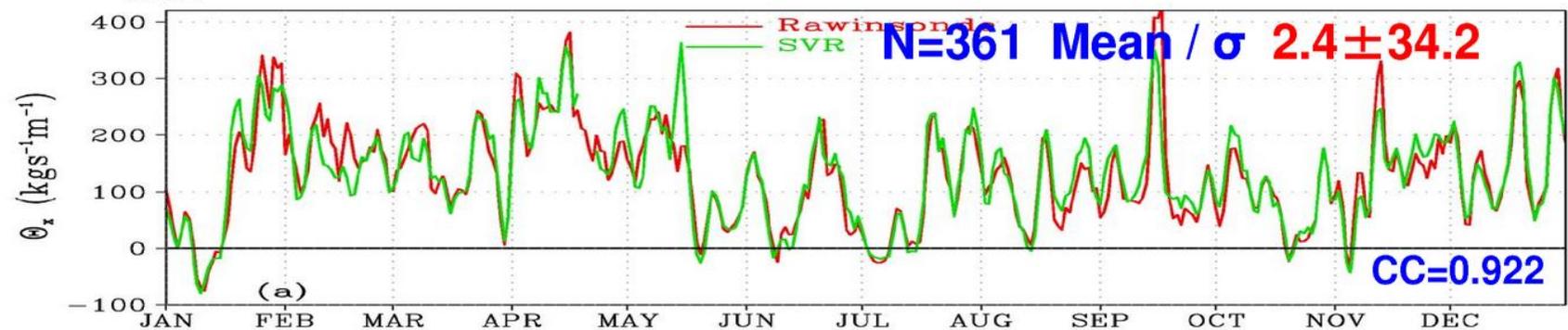
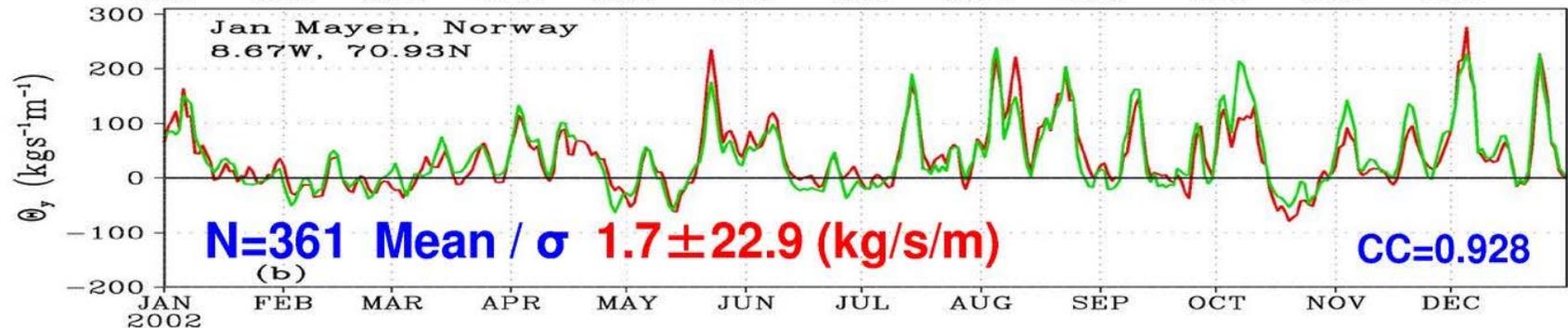
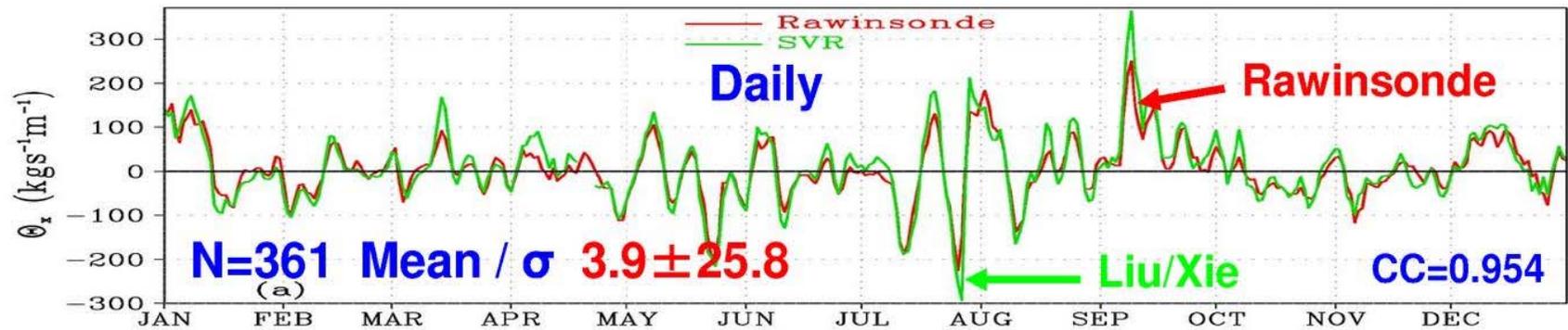
1. U_N : scatterometer surface wind stress

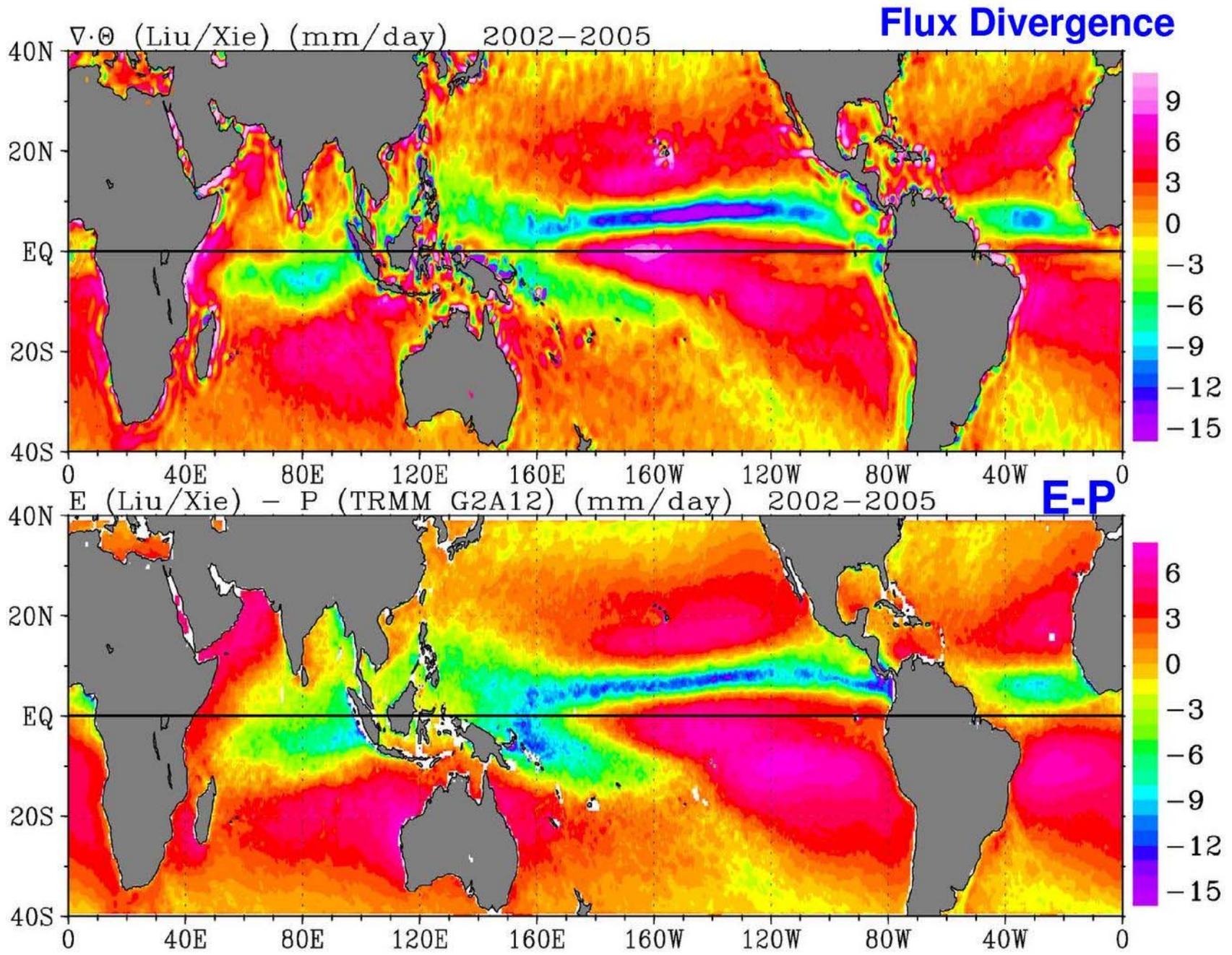
2. U_{850mb} : cloud drift wind (free-stream wind)

Two ways of estimating air-sea water flux- can they be reconciled?

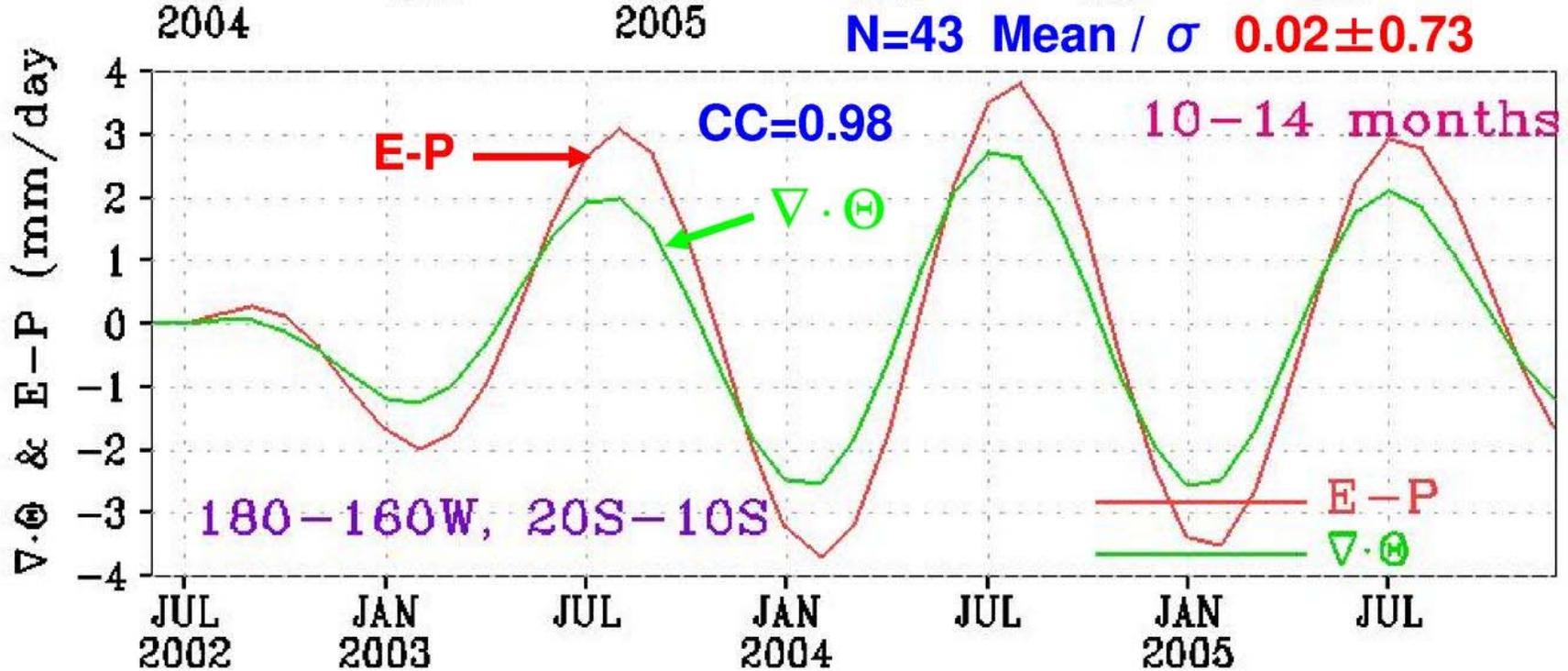
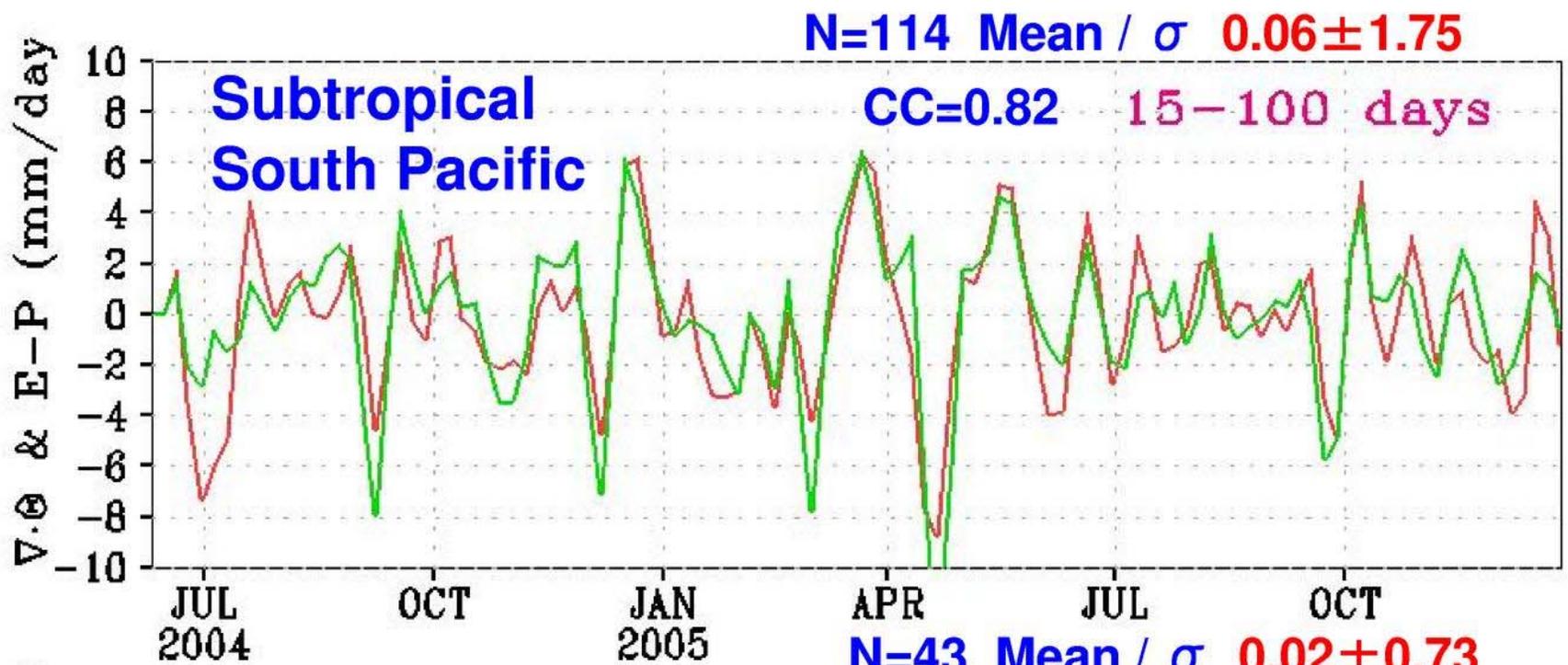
- § Evaporation is transported by turbulence. At small scale of turbulence, factors of atmospheric circulations, such as, Coriolis force, pressure gradient force, baroclinicity, cloud entrainment, are not important
- § All these factors are important in the divergence of moisture transport integrated through atmos. depth
- § Bulk parameterization was first used as a zero order approximation of what we wanted from what we had, bulk measurements. We hid our ignorance or incapability in the coefficient and we need to understand its limit

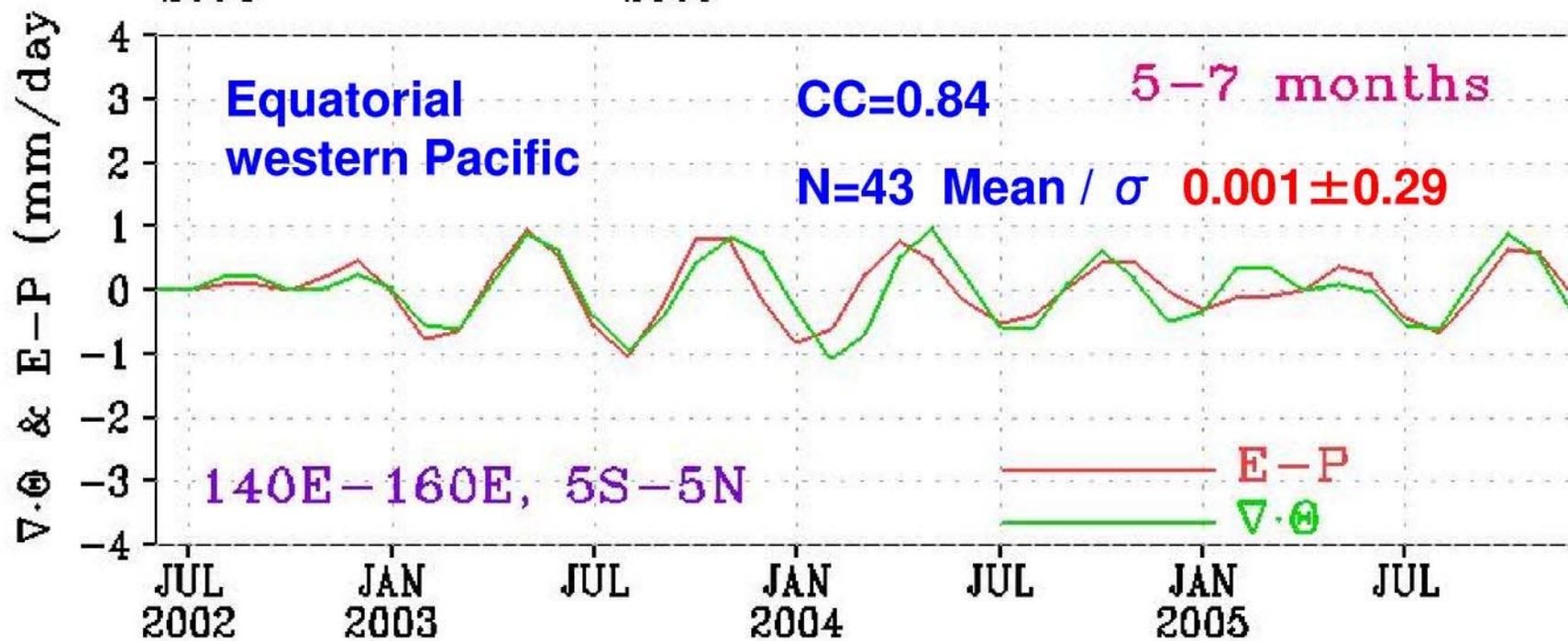
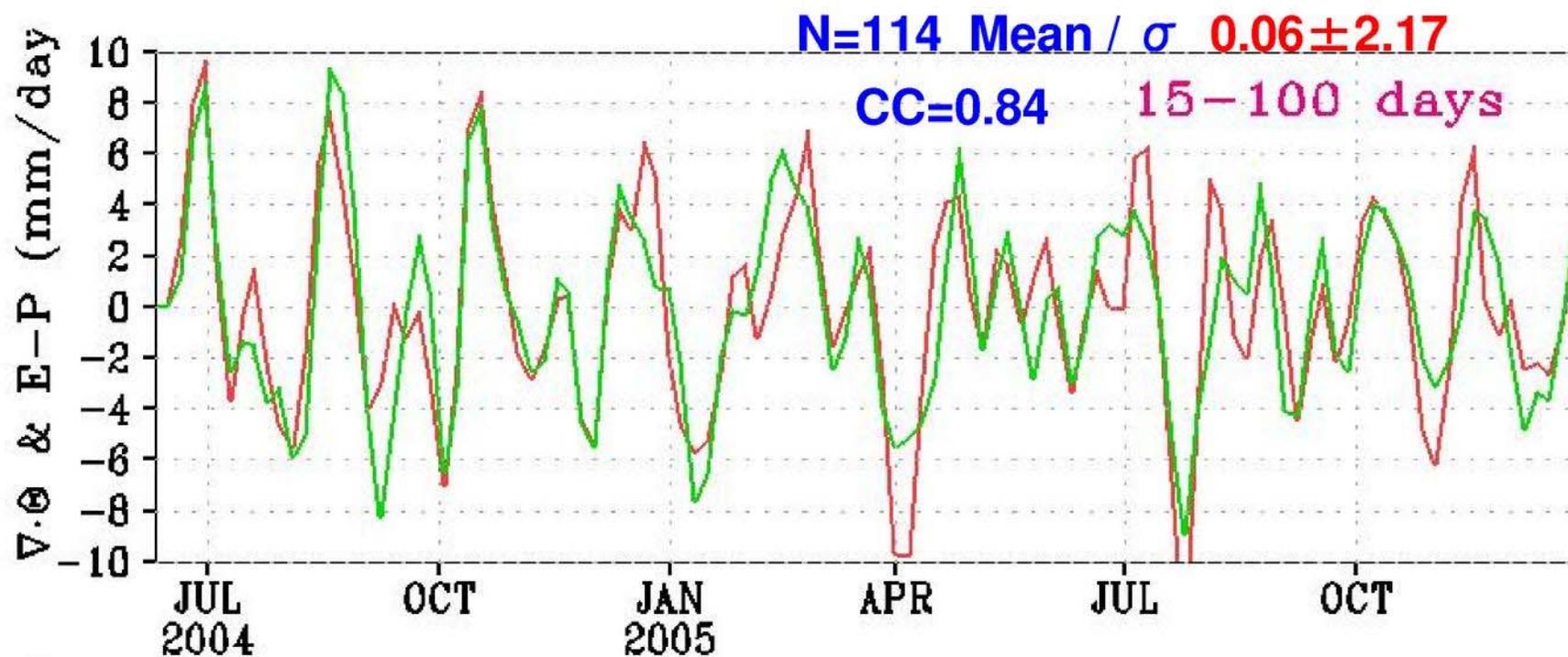




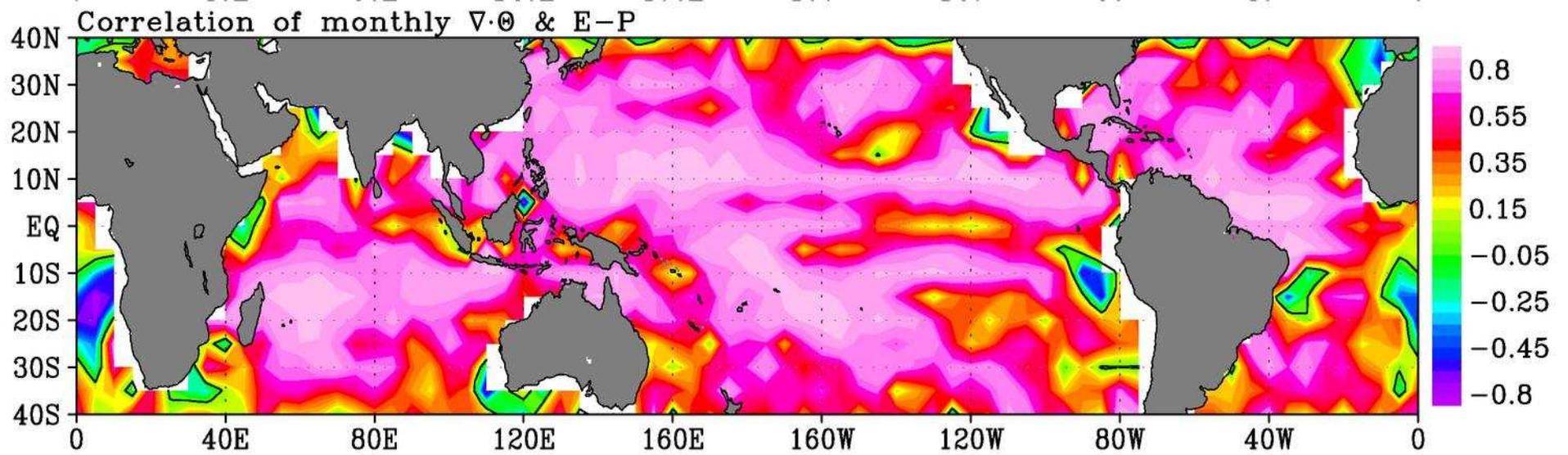


N=82589 Mean / standard deviation -0.06 ± 0.81 (mm/day)

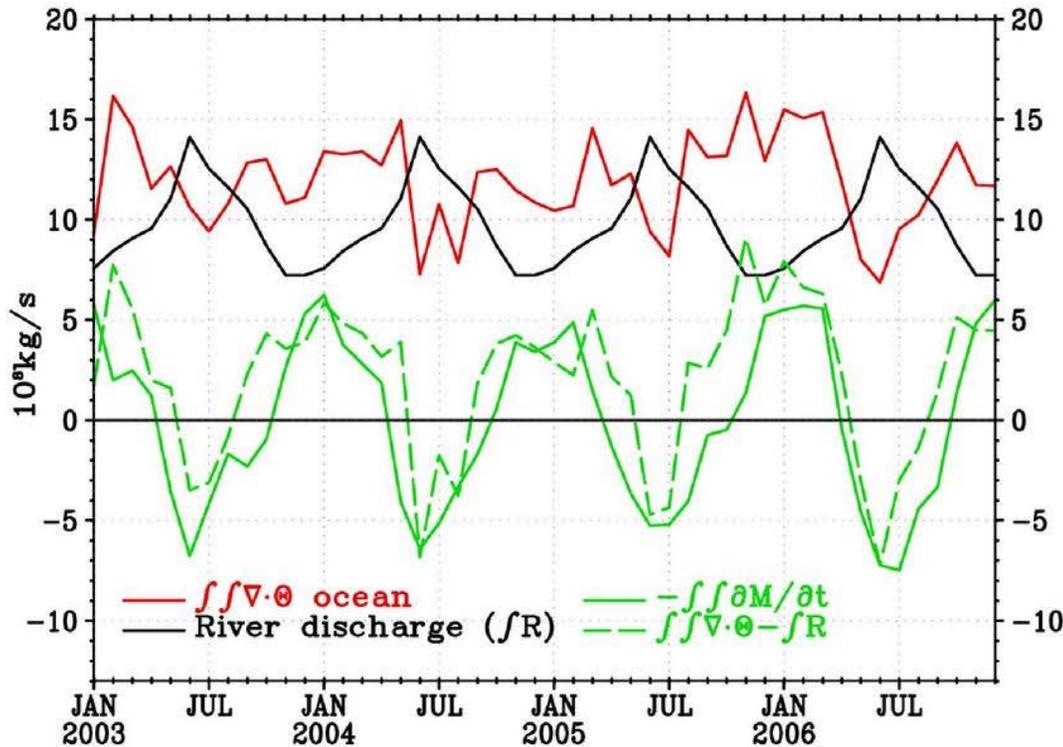




Correlation of $\nabla \cdot \Theta$ and E-P



Global Ocean Water Balance



Mean / standard deviation
2.14 ± 2.62 (10⁸kg/s) 20%

$$\iint \frac{\partial M}{\partial t} = \int R - \iint \nabla \cdot \Theta$$

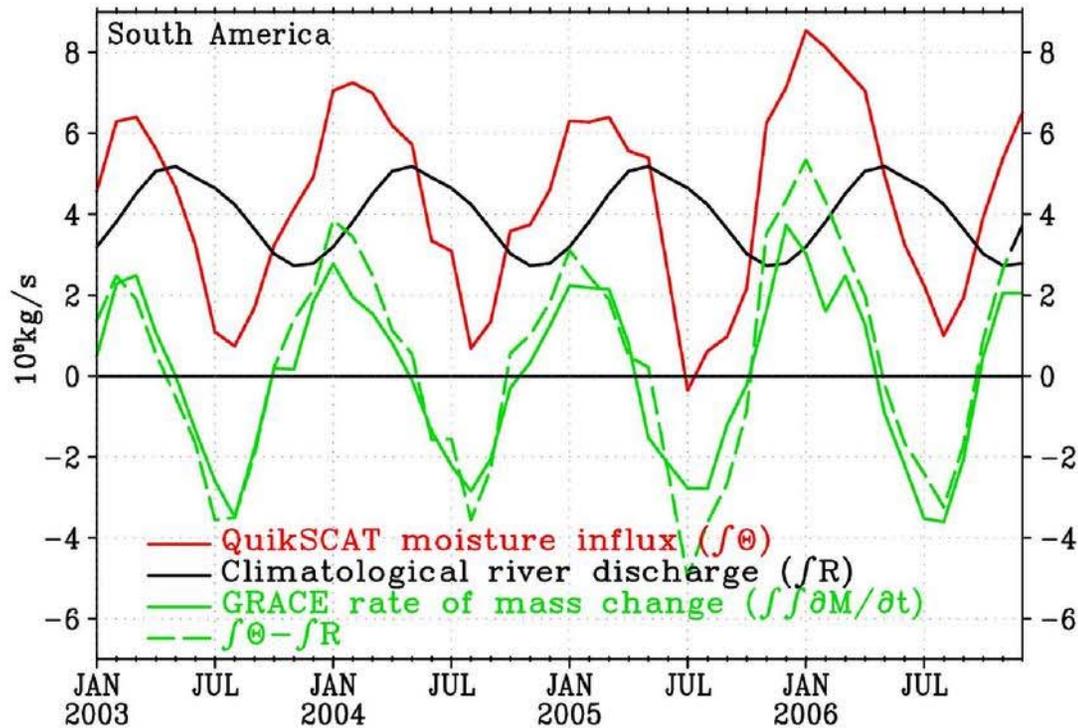
GRACE Dai/Trenberth Liu/Xie

$$\iint (E-P) = \int \Theta = \iint \nabla \cdot \Theta$$

Four-year means in cm/yr
div of water transport

Liu (2010)	10.6
Hilburn	21.9
E-P	
Merra (2008)	10.6
NEWS	10
Budyko (1974)	12
River discharge	
Dai (2002)	8.6

Water Balance over South America



Mean / standard deviation
 0.36 ± 0.93 (10^8 kg/s) 13%

$$\iint \frac{\partial M}{\partial t} = \int \Theta - \int R$$

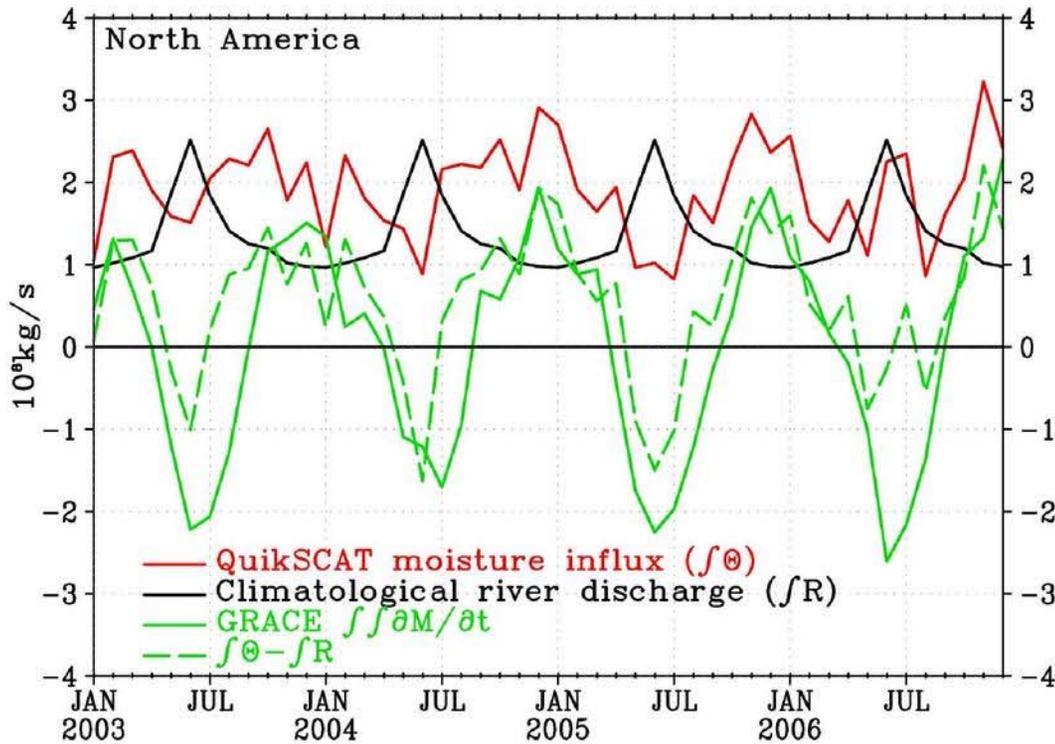
GRACE Liu/Xie Dai/Trenberth

$$\iint (P - E) = \int \Theta$$

Four-year means in cm/yr
Moisture into continent

Liu	76.1
Hilburn	154.5
P-E	
NEWS	61.3
Budyko (1974)	73
River discharge	
Dai (2002)	69.2

Water Balance over North America



Mean / standard deviation

1.92 ± 1.31 (10^8 kg/s) 17%

$$\iint \frac{\partial M}{\partial t} = \int \Theta - \int R$$

GRACE

Liu/Xie Dai/Trenberth

$$\iint (P - E) = \int \Theta$$

Four-year Means in cm/yr

Moisture into continent

Liu 28.9

Hilburn -5.9

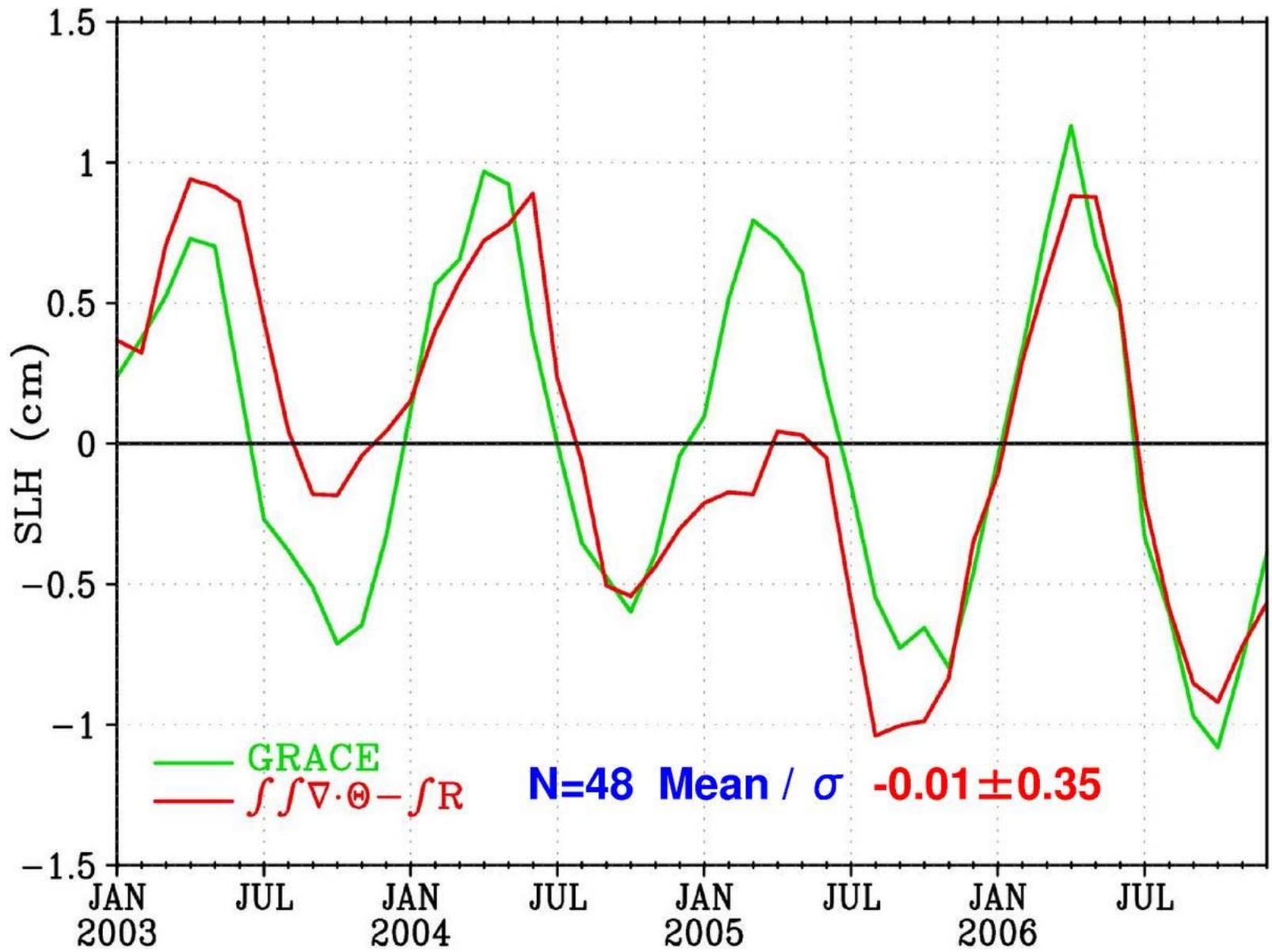
P-E

NEWS 20.9

Budyko (1974) 34

River Discharge

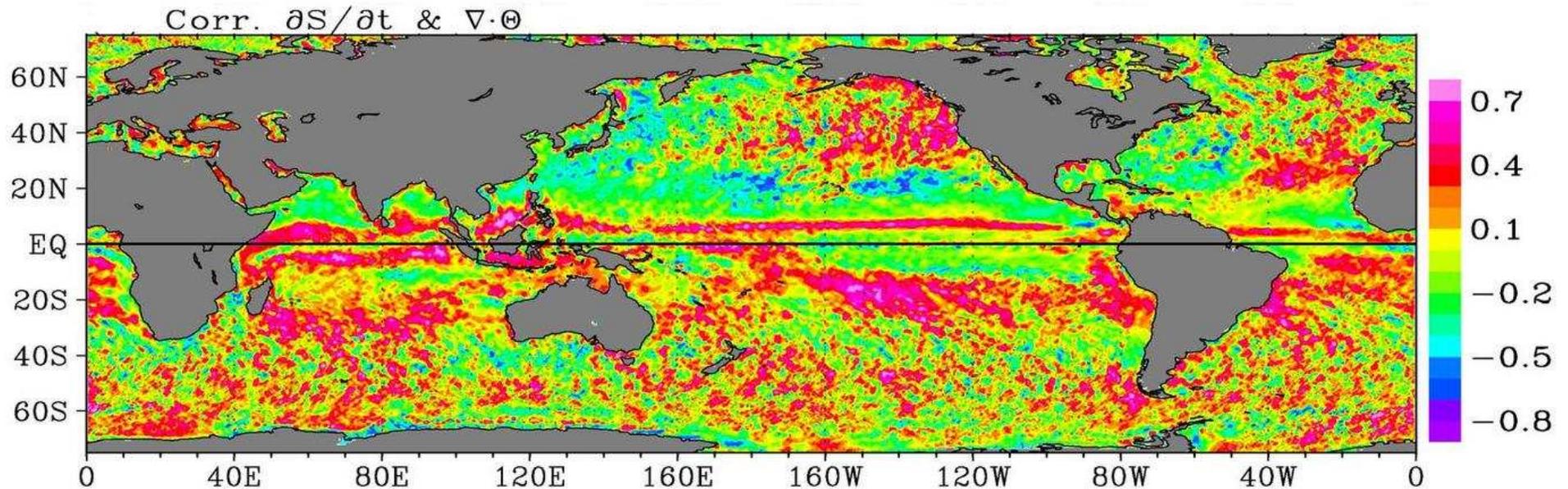
Dai (2002) 20.5

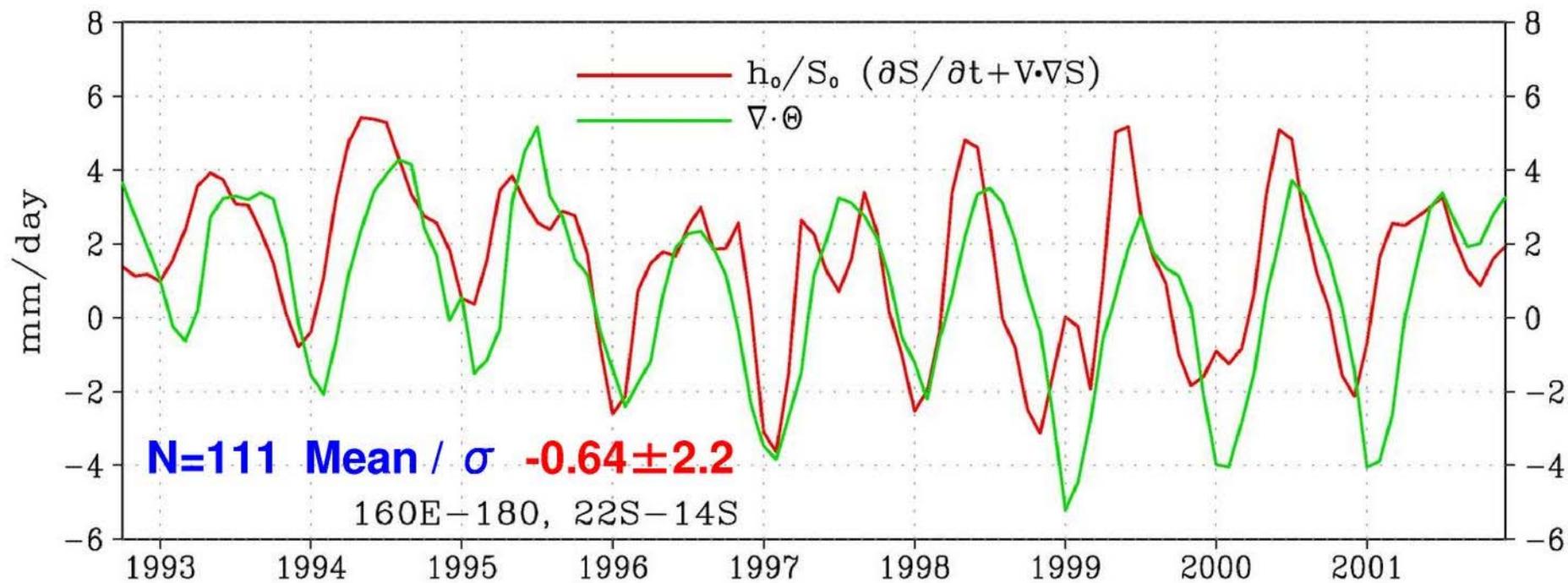
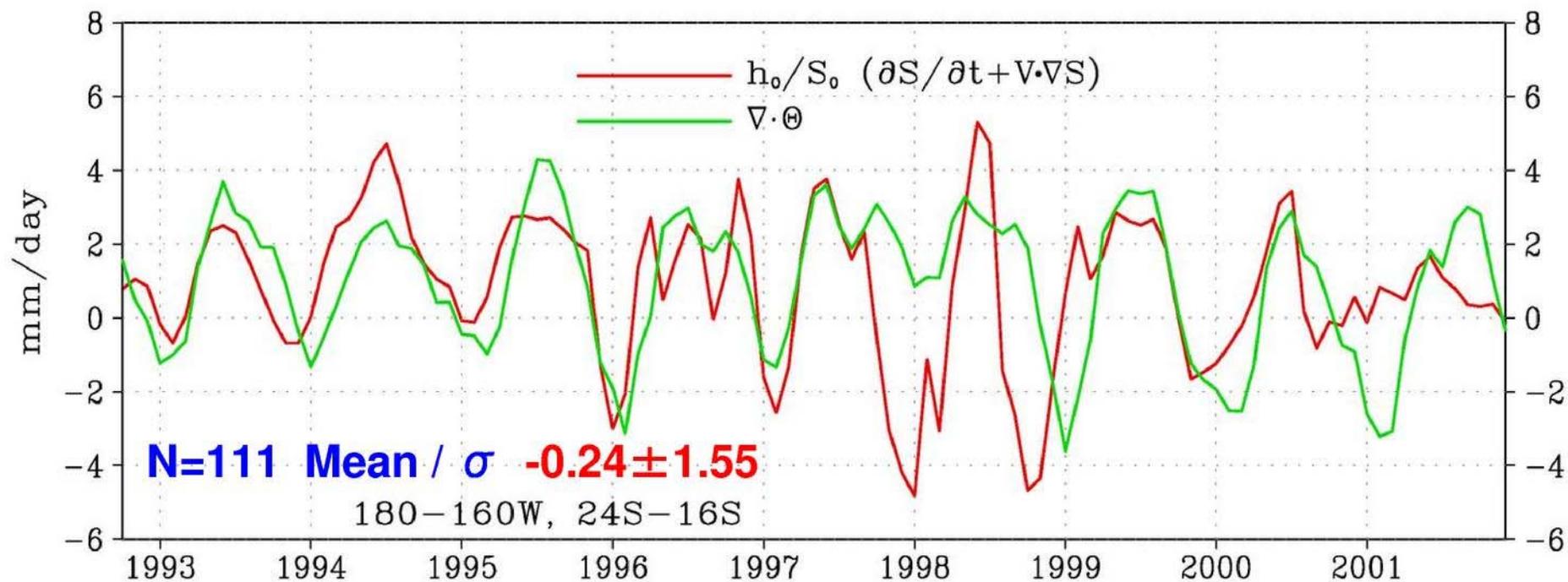


Sea level variation calculated from GRACE & moisture transport

Water and Salinity Balance

$$E - P = \frac{h_0}{S_0} \left(\frac{\partial S}{\partial t} + \mathbf{V} \cdot \nabla S \right)$$





Meridional Water Transport (MWT)

Conservation of water mass

$$\frac{\partial M}{\partial t} + \nabla \cdot \psi = P - E$$

By Green's theorem

$$\text{MWT}(\theta) = \int_{\theta}^{\theta_0} \int_{x_1}^{x_2} \left(\frac{\partial M}{\partial t} + E - P - R \right) dx dy$$

Ekman water transport

$$\text{EWT}(\theta) = \int_{x_1}^{x_2} -\frac{\tau_x}{\rho f} dx$$

P: Precipitation

E: Evaporation

ψ : Horizontal mass flux

R: River discharge

τ_x : Zonal stress

