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# Symmetric instability in cross equatorial western boundary currents

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# The global conveyor

Interactions between the sub-mesoscale and global scale circulations

- Northward flowing surface current
- Southward flowing deep western boundary current
- Global scale current, drives
  cross-equatorial flow
- Can sub-mesoscale instability at the equator lead to a bottleneck?





Image credit: Greg Holloway, Dan Wright

# **Tropical circulations**

- Northward flowing surface currents
  - North Brazil Current & rings
  - Equatorial Counter Current
- Southward flowing deep western boundary current





Dengler et al. (2004)

 $) \bigcirc \bigcirc$ 

# **Potential vorticity** And a necessary condition for symmetric instability

• PV is materially conserved:



- SI if fQ < 0
  - If initially stable in SH, not stable in NH
- Growth rate  $\sigma^2 \approx -f(f + dV/dx)$ 
  - Need large horizontal shear







## What does symmetric instability look like?

- Predictions from a linear stability analysis
- Stacked overturning cells
- Localised in regions of negative PV
- Could be an important mixing mechanism





# An idealised model

- Simplifications made:
  - Brazil is a straight line
  - No topography
  - Open boundaries (sponged)
- Unstable regions have negative PV in the northern hemisphere



# An idealised model **Two types of instability**

- PV at 50 m
- What's going on?
  - 1. Eddy field develops as fluid crosses the equator — e.g. Edwards & Pedlosky, 1998; Goes et al. 2009.





# An idealised model **Two types of instability**

- PV at 50 m
- What's going on?
  - 1. Eddy field develops as fluid crosses the equator — e.g. Edwards & Pedlosky, 1998; Goes et al. 2009.
  - 2. SI is excited from 300 km North of the equator.





# PV and relative vorticity



## Sign of RV and PV uncorrelated $\implies$ planetary vorticity dominates



# Conclusions

- flows.
- currents.
- We can see the effects of the instability on:
  - potential vorticity of eddy cores
  - correlations between potential and relative vorticity
- Next steps:
  - Deep western boundary currents
  - LLC4320 model
  - Existing glider datasets?



## • From theoretical considerations we might expect to observe symmetric instability in cross-equatorial

• Symmetric instability has been observed in an idealised model of cross-equatorial western boundary

• For more details see Goldsworth et al. (2021), *Journal of Physical Oceanography* (early online release)

# Supplementary slides

# Instability & power spectra

- Spatially Fourier transform  $w^2/2$  at 50 m
- Plot at week long intervals
- SI causes flattening and lifting of the spectra
- See also Yankovsky & Legg (2019)











# **Evidential summary**

- 1. Regions of negative PV are unstable
- 2. Vertical scale
- 3. Horizontal scale
- 4. Time scale
- 5. Viscosity dependence
- 6. Structure of overturning
- 1. 3 dimensional models
- 8. Power spectra of the vertical KE
- 9. Correlations between PV and relative vorticity

General agreement between linear stability analysis, simplified two dimensional models and

## Why symmetric instability?

## **Predictions from a hierarchy of models**





## 2D numeric overturning streamfunction



![](_page_13_Figure_7.jpeg)

![](_page_13_Figure_9.jpeg)

Longitude (km)

Longitude (km)

![](_page_13_Picture_12.jpeg)

## Water mass formation The Wallin framework

![](_page_14_Picture_1.jpeg)

![](_page_14_Figure_2.jpeg)

 $\Delta \psi = \psi_N - \psi_S$ 

![](_page_14_Figure_4.jpeg)

 $y_N$ 

![](_page_15_Figure_0.jpeg)

# Water mass formation The Wallin framework

![](_page_15_Figure_2.jpeg)

 $(\times 10^{12})$ formation lass water m

![](_page_15_Figure_4.jpeg)

![](_page_15_Figure_5.jpeg)

![](_page_15_Figure_6.jpeg)

![](_page_15_Figure_7.jpeg)

![](_page_15_Figure_8.jpeg)

![](_page_15_Figure_10.jpeg)

![](_page_15_Figure_11.jpeg)

![](_page_15_Figure_12.jpeg)

![](_page_15_Figure_13.jpeg)

250

-2 -4 02

80

![](_page_15_Figure_14.jpeg)

![](_page_15_Figure_15.jpeg)

40

16

![](_page_15_Figure_18.jpeg)