

Mixing in the Equatorial Cold Tongues

A TAO / PIRATA Comparison of χ pod Adventures

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D. Cherian⁵, W.D. Smyth¹ and E.L. Shroyer¹

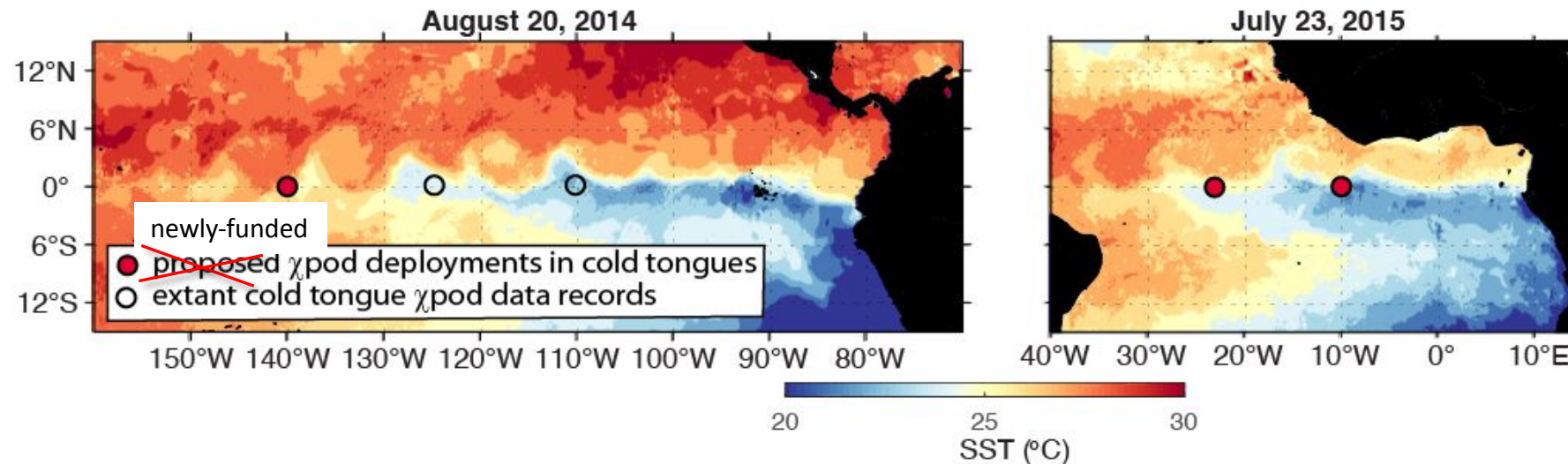
¹OSU ²Brandeis University ³IRD ⁴GEOMAR/Kiel ⁵NCAR



National Oceanic and Atmospheric Administration's
National Data Buoy Center
Center of Excellence in Marine Technology



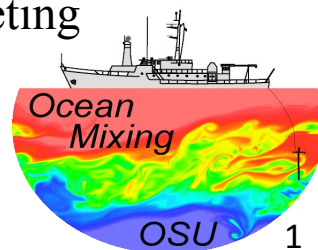
IRD
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PIRATA-24/TAV

24th Prediction and Research Moored Array in the Tropical Atlantic /Tropical Atlantic Variability Meeting

<https://cpaess.ucar.edu/meetings/2021/pirata-24tav-abstract-information>



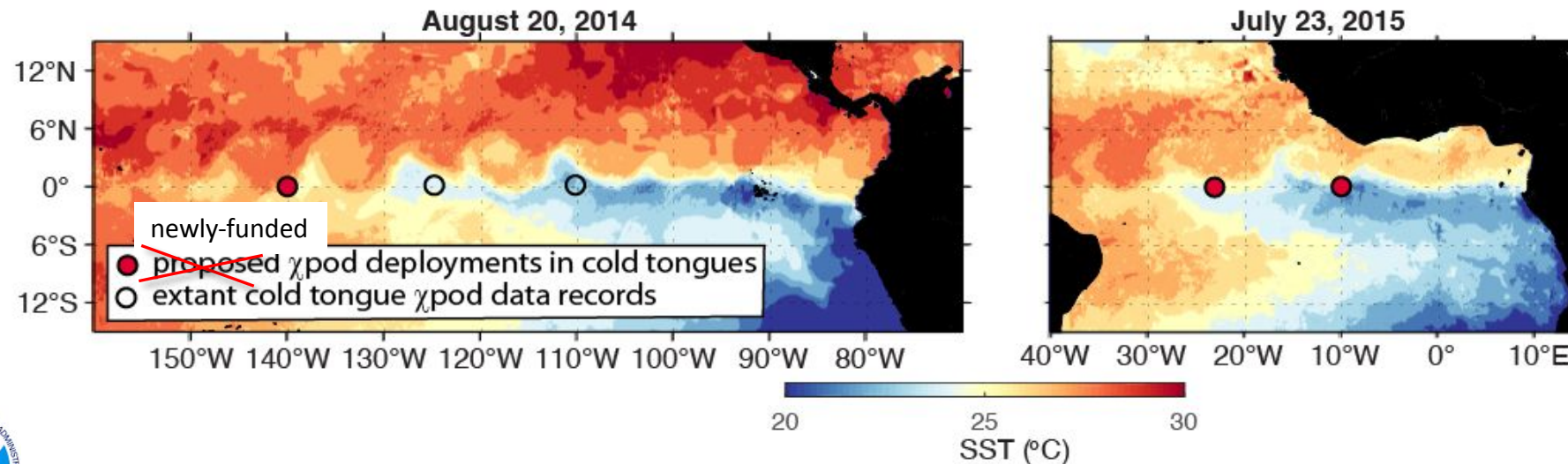
Cold Tongue Mixing / NSF

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³Department of Marine and Environmental Sciences, Hampton University



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TAO



PIRATA
RAMA

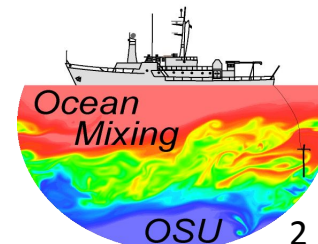


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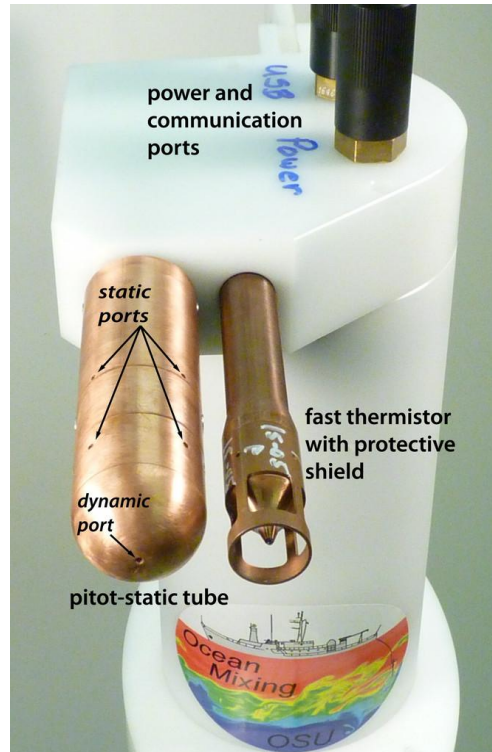
New 5y NSF grant to deploy xpods at

- 0°, 10°W
- 0°, 23°W
- 0°, 140°W

and to push on the analysis

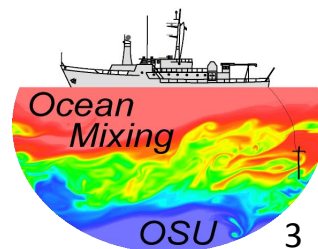


χ pod



Technical publications define methods/processing:

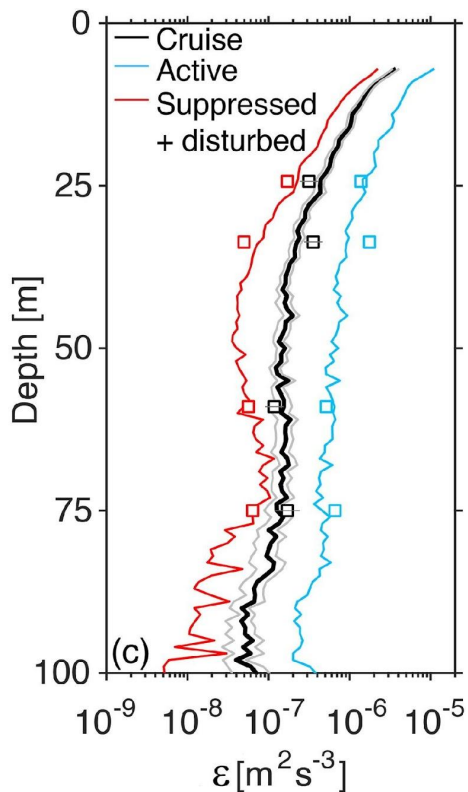
Moum & Nash, 2009 (*J. Atmos. Ocean Technol.*)
 Zhang & Moum, 2010 (*J. Atmos. Ocean Technol.*)
 Perlin & Moum, 2012 (*J. Atmos. Ocean Technol.*)
 Moum, 2015 (*J. Atmos. Ocean Technol.*)
 Becherer & Moum, 2017 (*J. Atmos. Ocean Technol.*)



How we know χ pods work

0, 80 E / 2011 DYNAMO

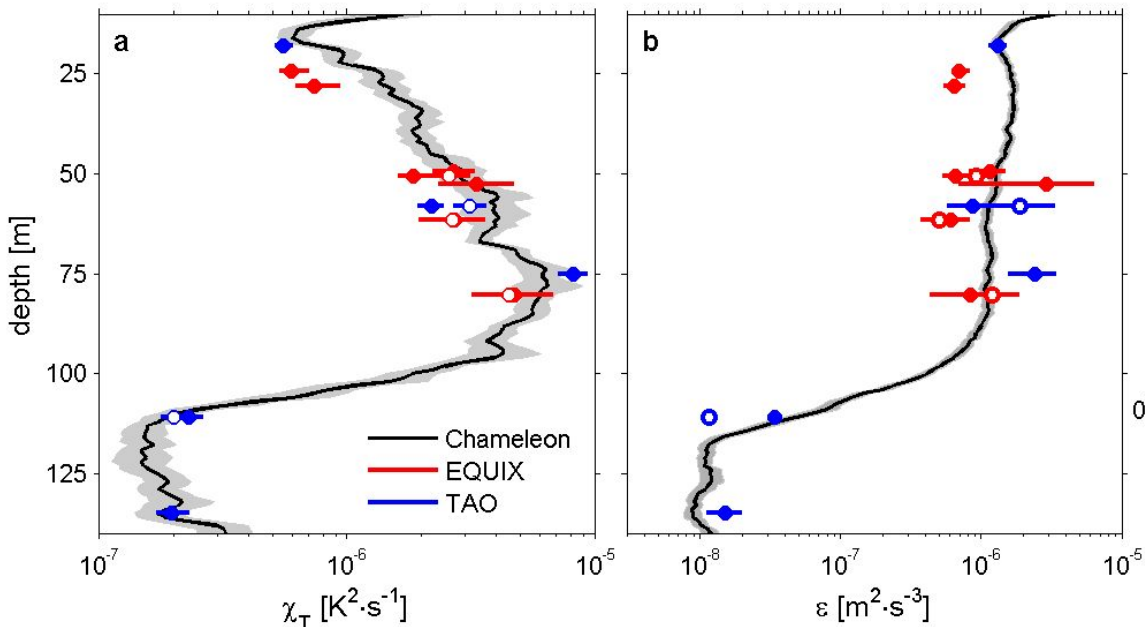
variation of ϵ with MJO phase



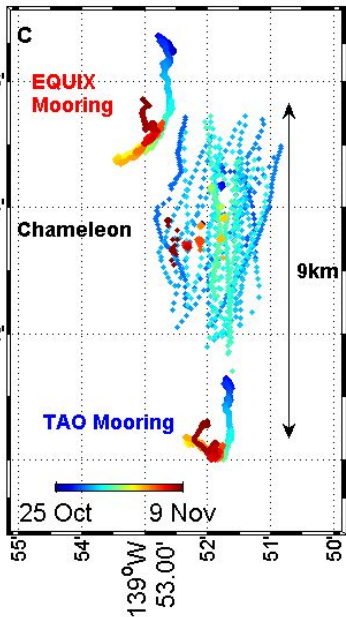
Pujiana et al 2018 (JPO)

0, 140W / boreal autumn 2008

variation of ϵ with depth



Perlin & Moum, 2012 (J. Atmos. Ocean Technol.)

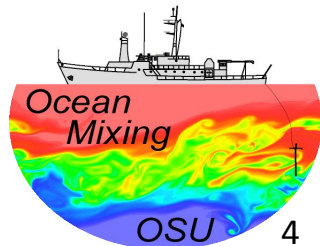


comparison of χ , ϵ from χ pods on moorings with turbulence profiling on shipboard field experiments

turbulence profiler *Chameleon* ϵ from shear probes

moored χ pods

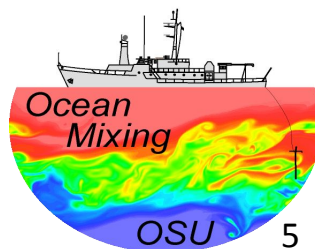
ϵ_χ inferred from $N^2 \chi / (2 \Gamma T_z^2)$

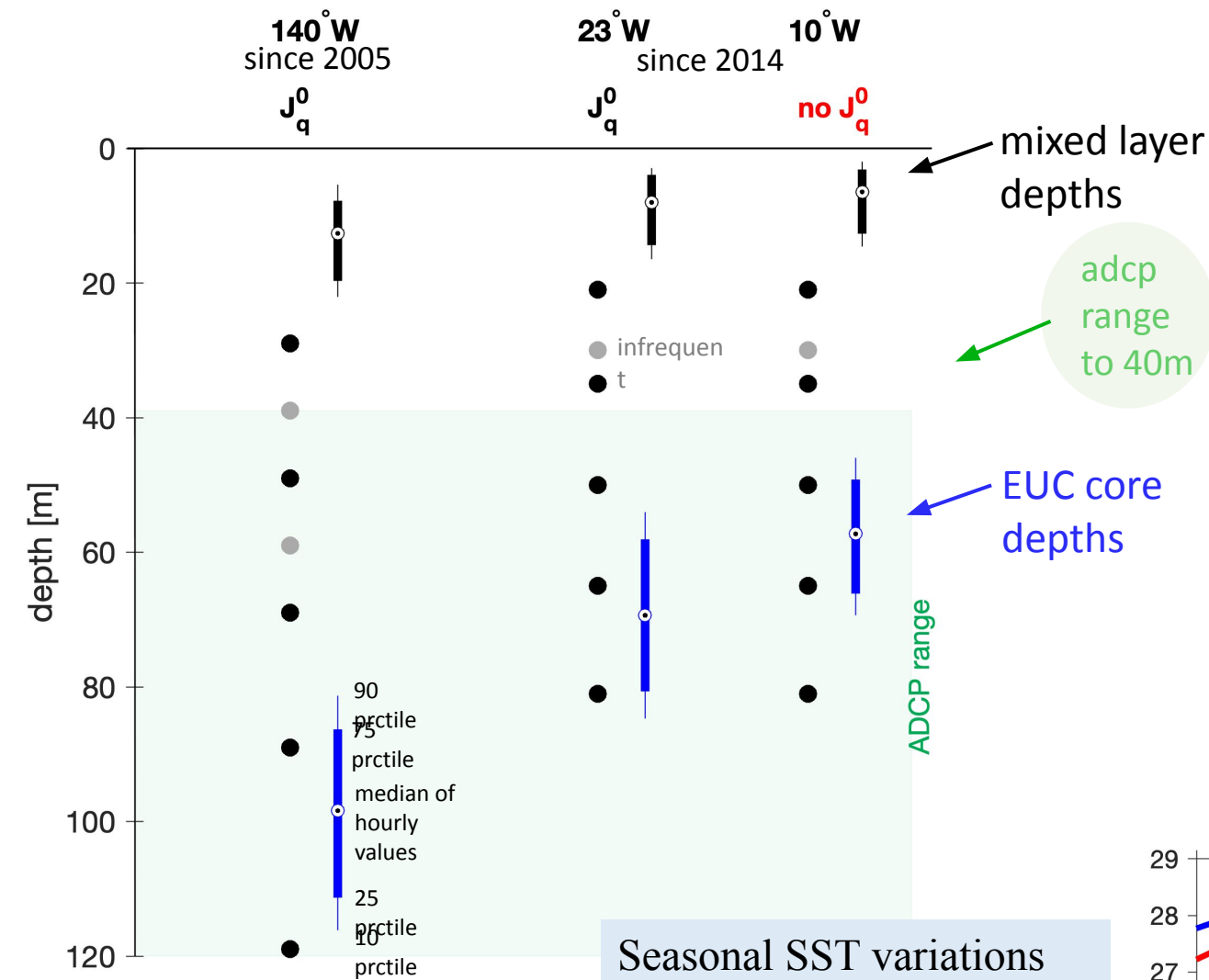


Objective: quantify the role of mixing of energy, heat and momentum on Cold Tongue Dynamics

To this end we derive from these measurements

1. turbulence diffusivities for heat and momentum, K_t and K_m ← these are sort of what modelers think they need
2. turbulence dissipation rates, ε ← these are rate variables that appear in equations governing rates of change of energy, heat and momentum
3. turbulence heat flux, J_q^t ← these tell us what the state variables **will** be
4. turbulence momentum flux, J_m^t ←





Seasonal SST variations

0, 140W 2°C
0, 23W >3°C
0, 10W >5°C

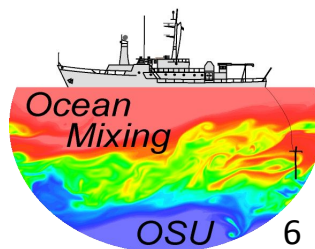
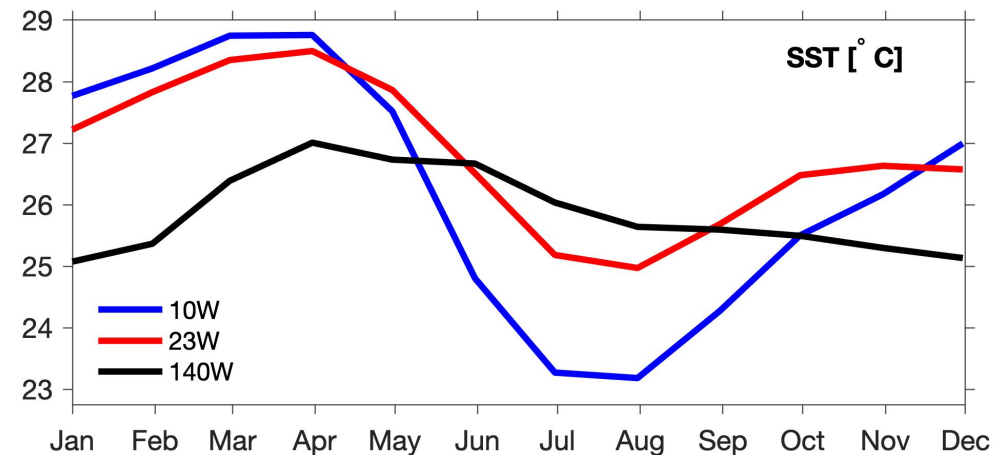
important *commonalities* of the 3 sites

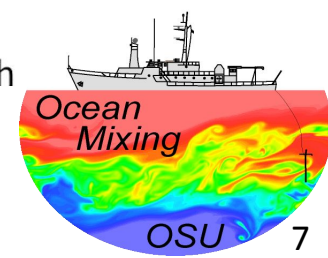
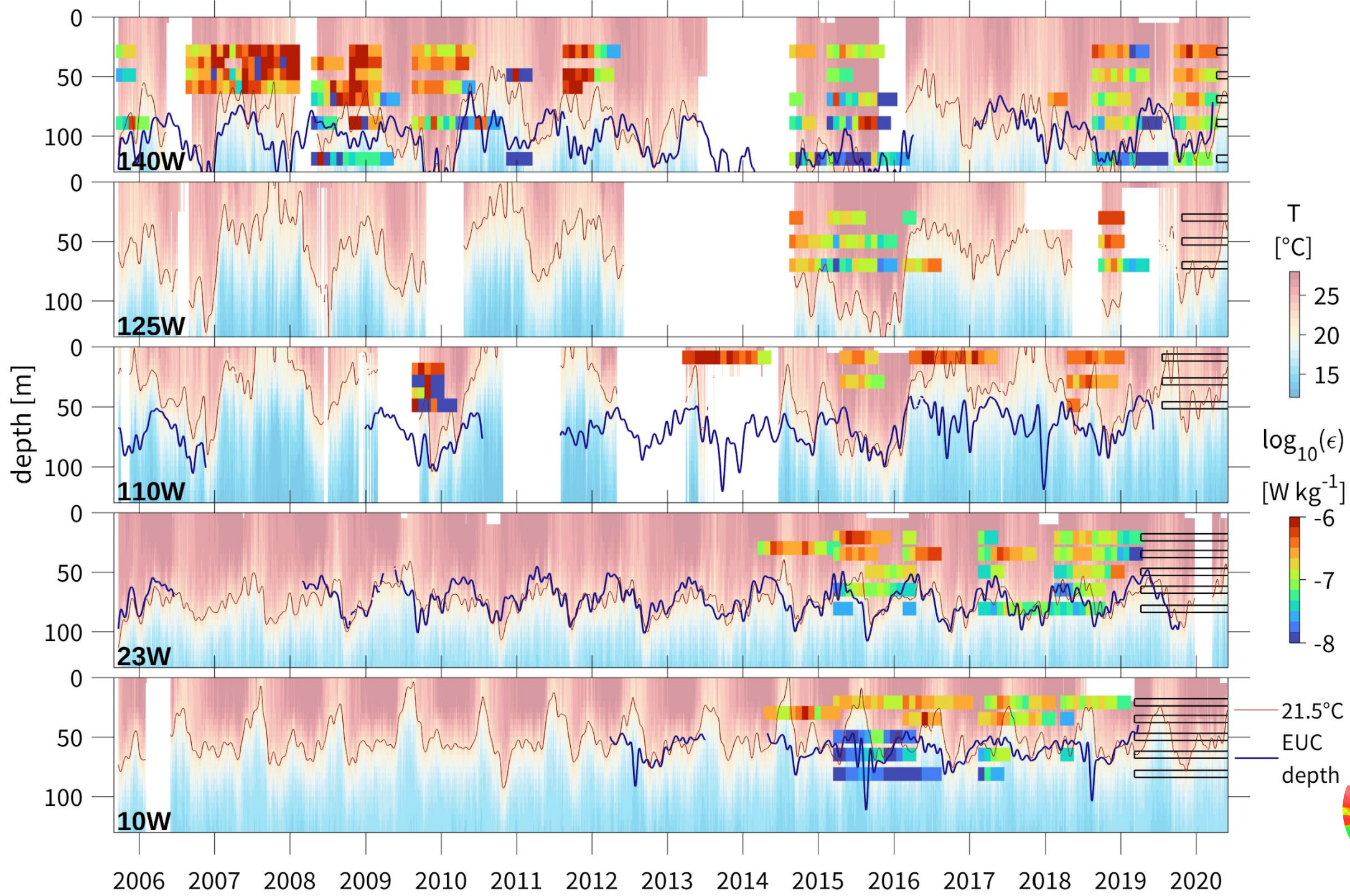
- EUC
- highly-sheared near surface flows
- $Ri \approx 1/4$ over MLD to EUC / marginal instability (MI)

important *distinctions* between the 3 sites

at 0, 140W

- windier
- greater TIW KE
- deeper ML, EUC
- thicker MI layer





Combined shipboard and moored χ pod measurements at 0, 140W have shown:

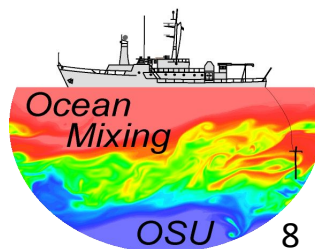
- diurnal deep cycle (below the mixed layer) of mixing / Moum et al 1989 (*JGR*); Lien et al, 1995 (*JGR*)
- marginal instability ($Ri \cong 1/4$) between MLD and EUC core / Smyth & Moum 2013 (*GRL*)
- descent rate of deep cycle (in S^2 , N^2 , ε) / Smyth et al, 2013 (*JPO*)
- seasonal variability of turbulence and heating that accounts for seasonal SST changes / Moum et al 2013 (*Nature*)
- feedbacks to ENSO phase change / Warner & Moum, 2019 (*GRL*)

today, we compare χ pod measurements

PCT (0, 140W)

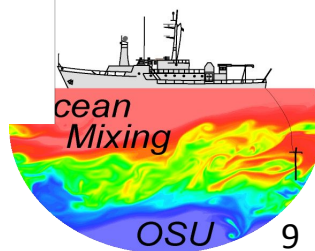
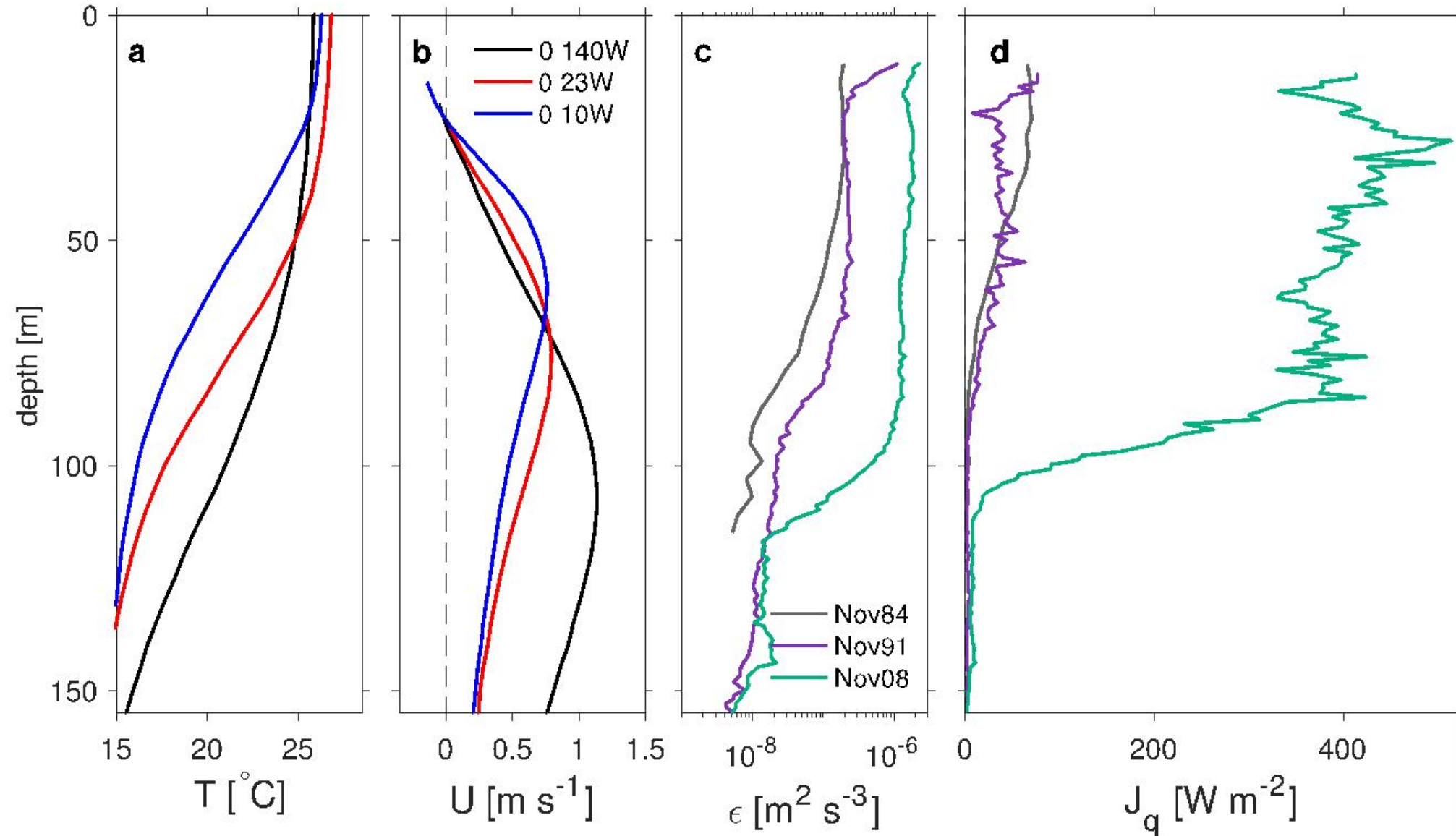
ACT (0, 10W; 0, 23W)

- *Averaged vertical profiles*
- *Diurnal variability of mixing*
- *Seasonal variability of turbulence and heating*

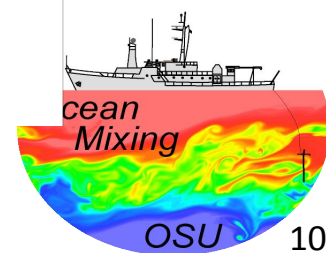
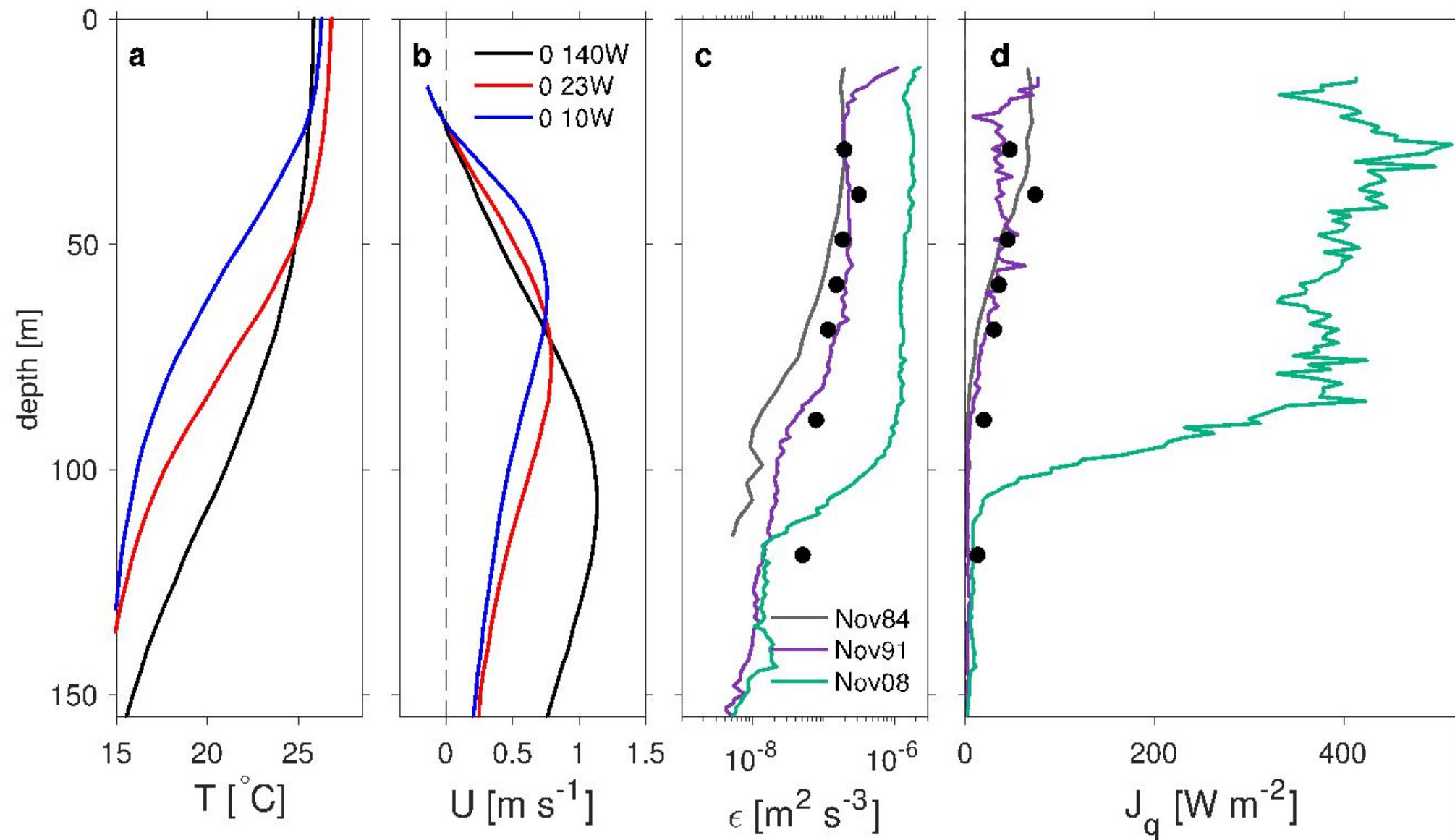


Averaged vertical profiles (2014-2019)

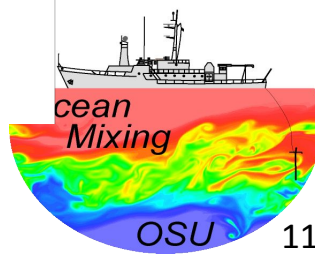
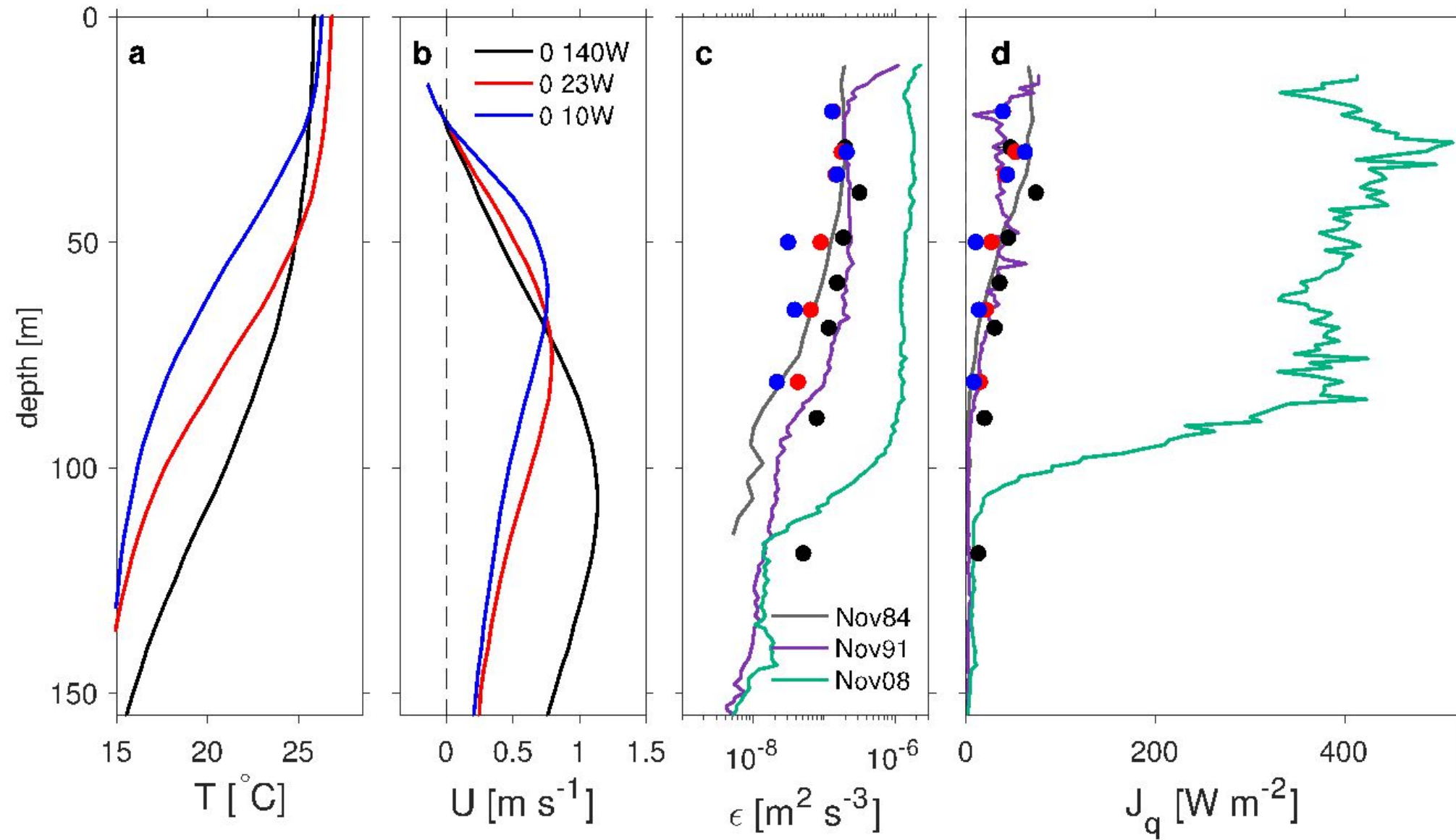
results from turbulence profiling experiments at 0 140W



add full 0,140W χ pod record (2005-2019)



add 0,10W and 0,23W χ pod records

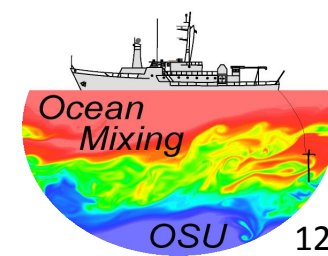


b

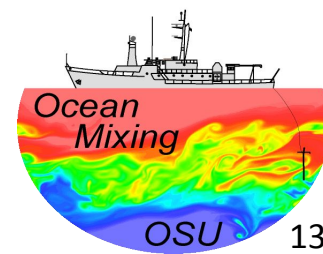
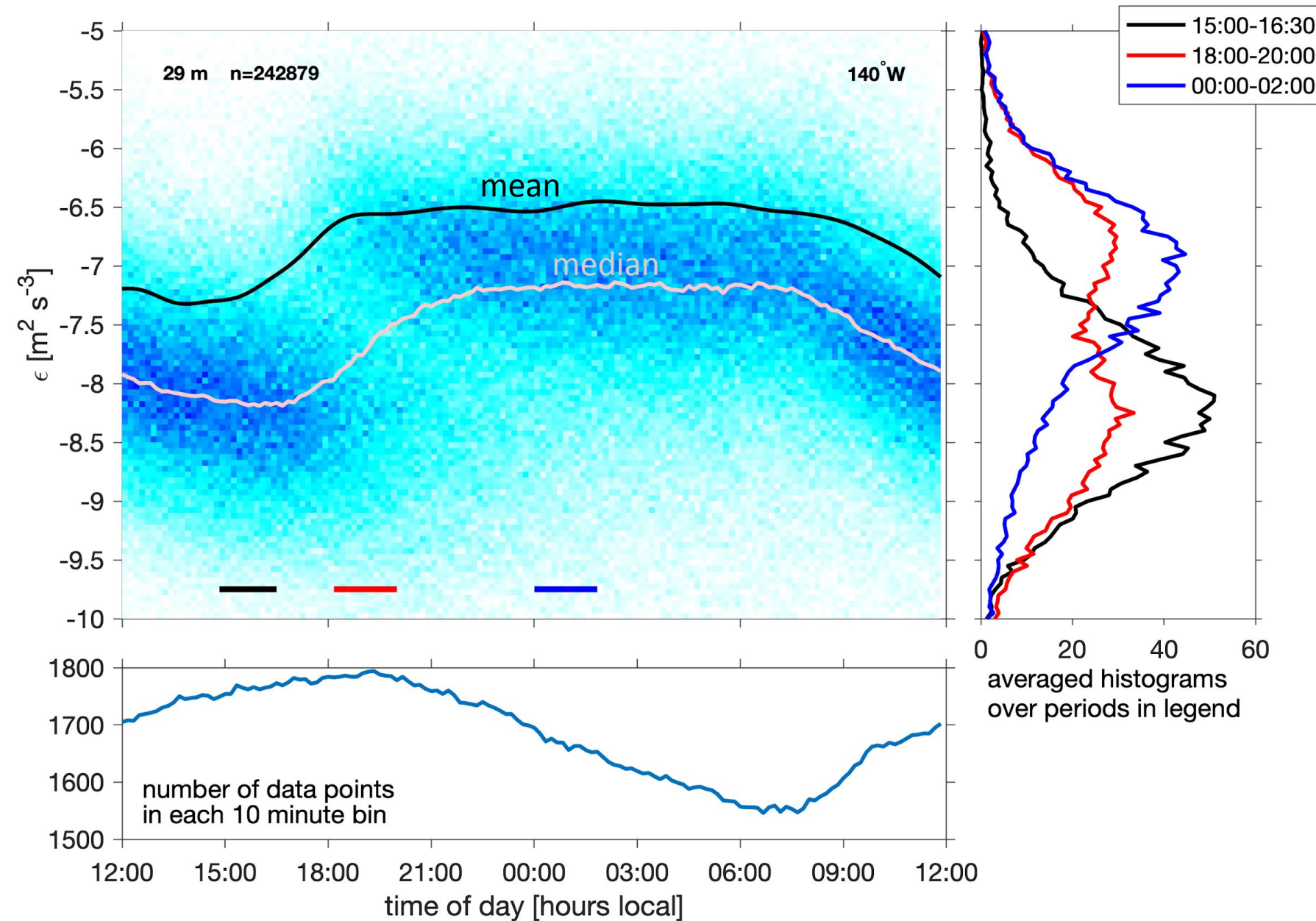
— 0 140W
— 0 23W
— 0 10W

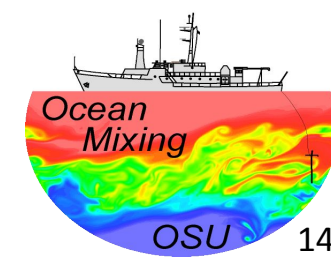
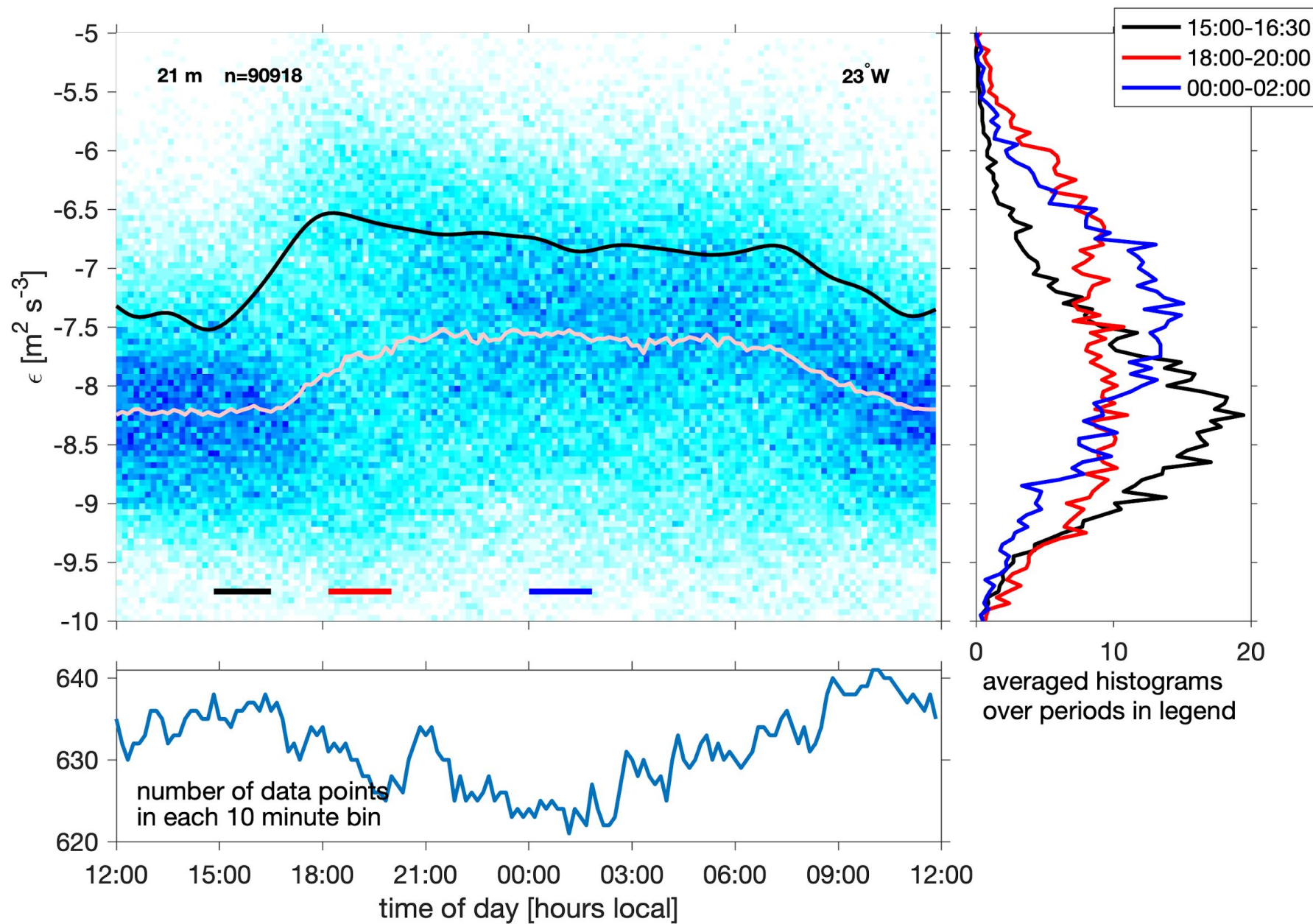
50 m

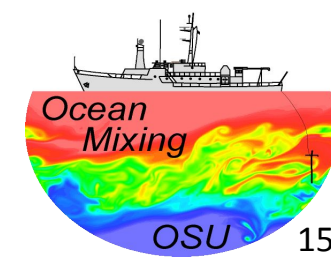
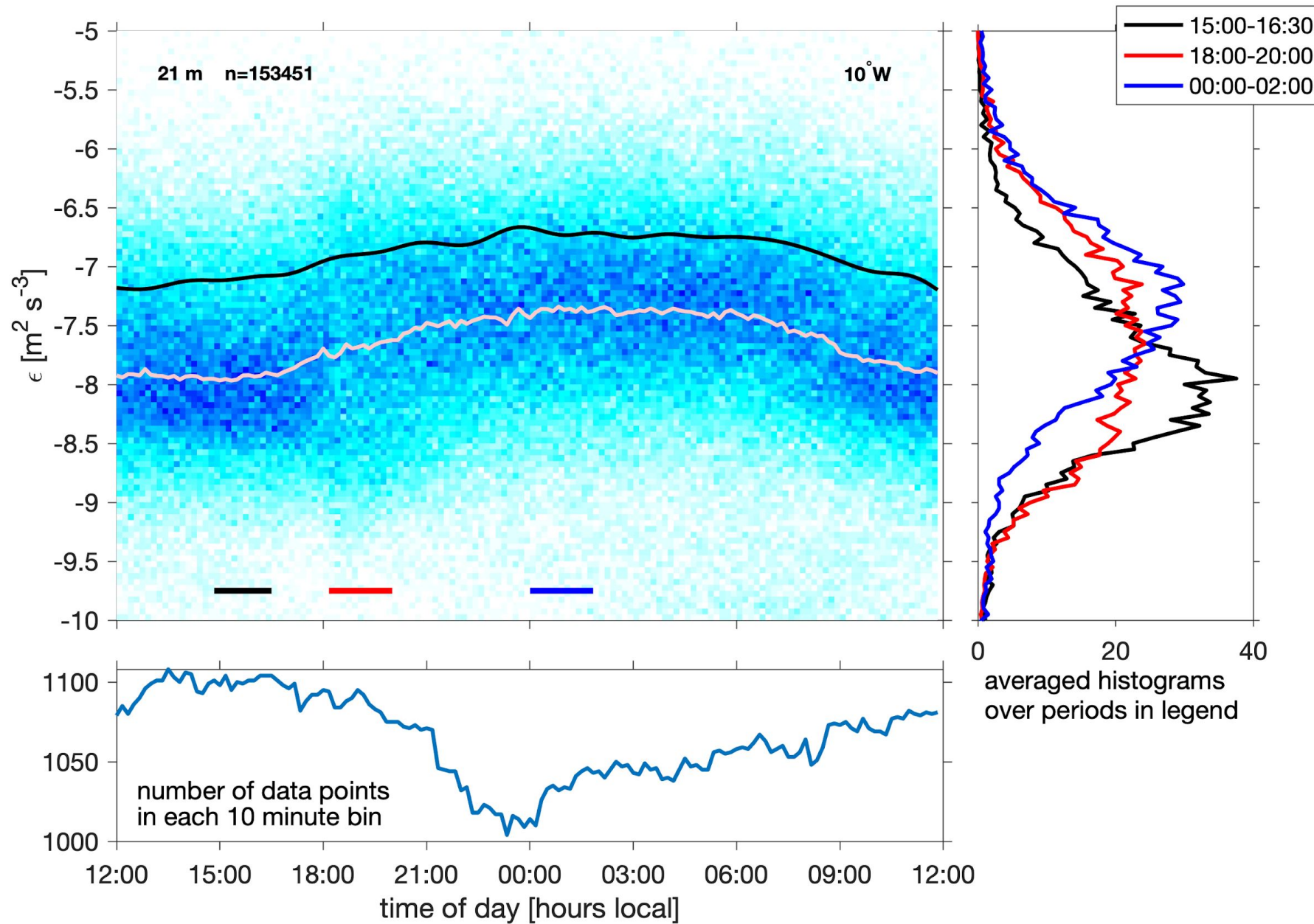
100 m

c $10^{-8} \text{ m}^2 \text{ s}^{-3}$ 10^{-6} ϵ **d**200 W m^{-2} 400 J_q^t 

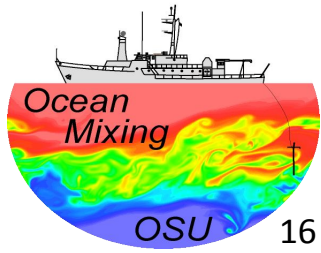
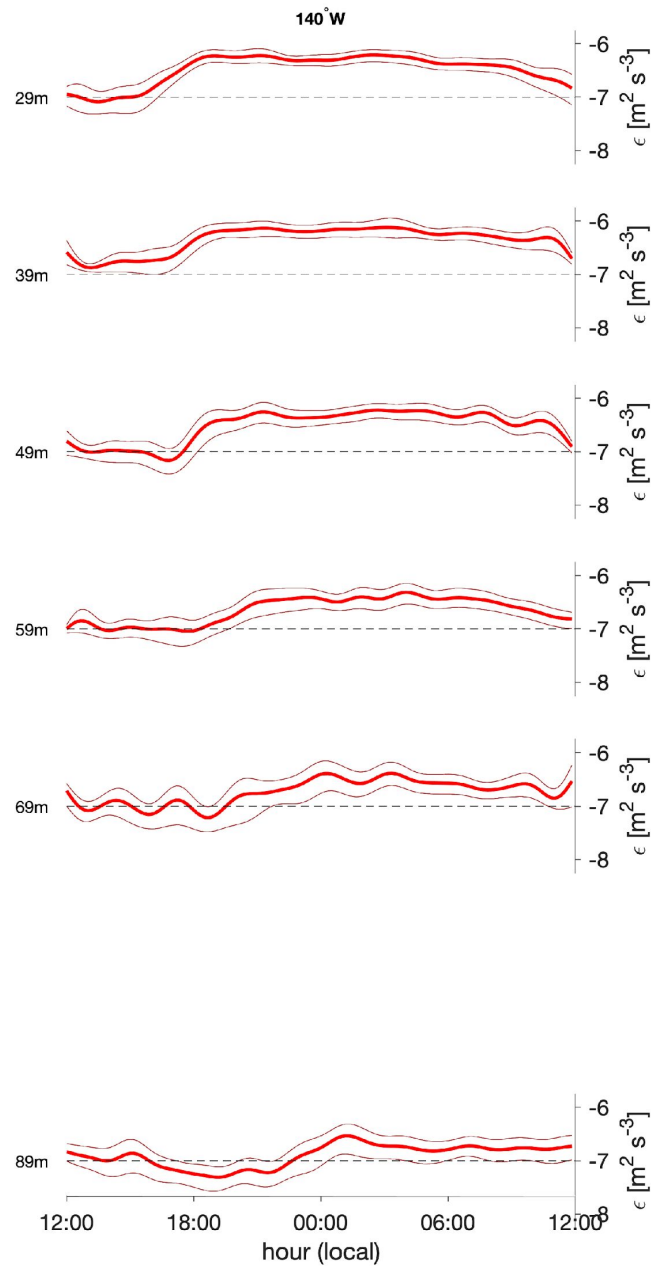
Diurnal variability of mixing





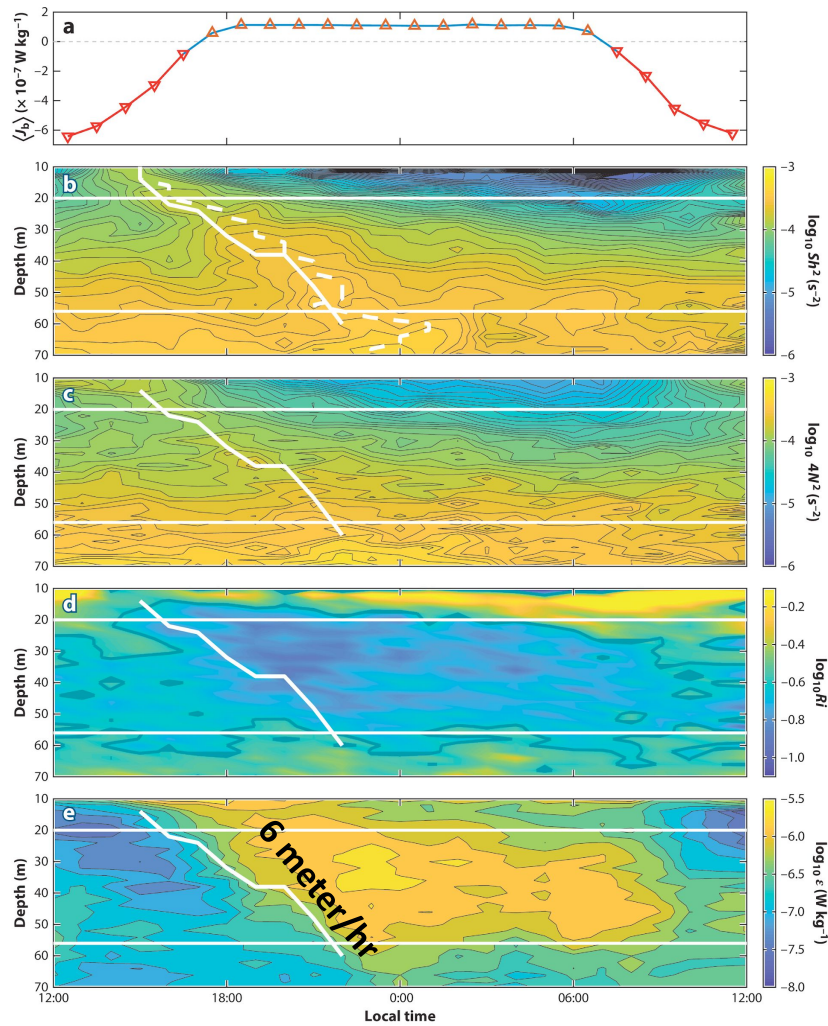


$\tau > 0.075 \text{ Pa}$



$\tau > 0.075 \text{ Pa}$

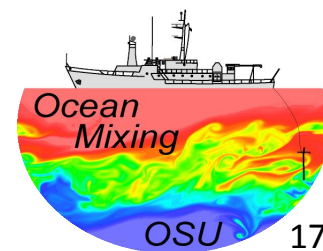
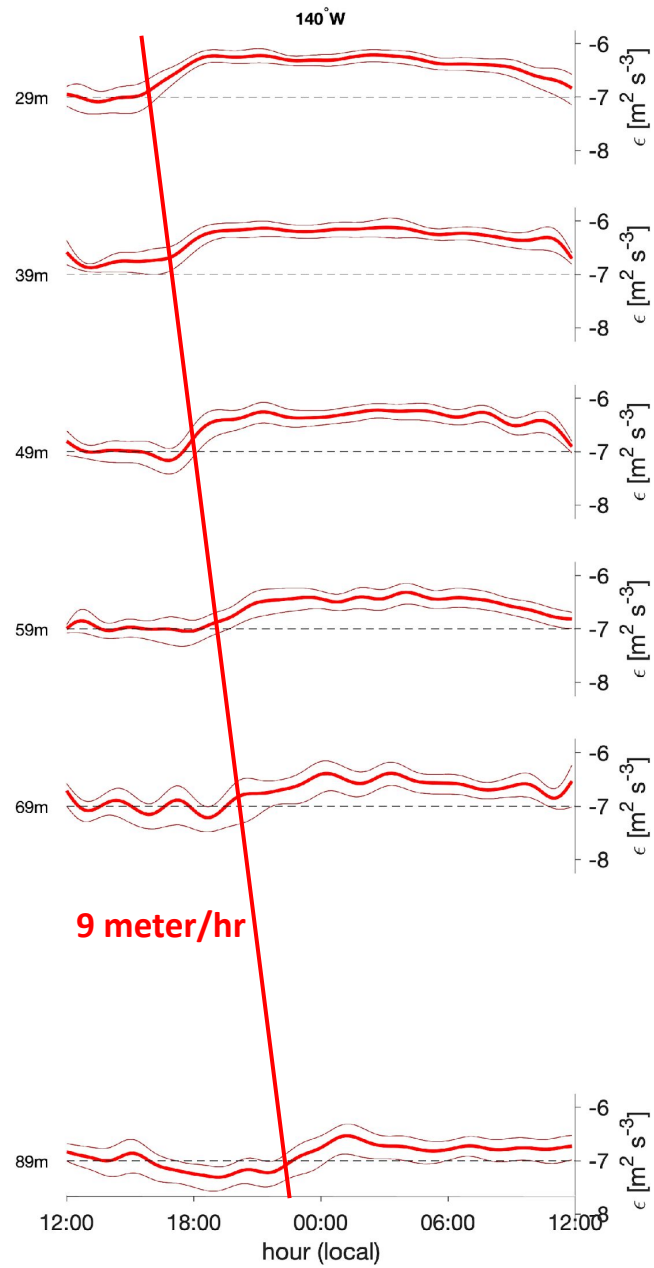
diurnal descent of S^2 , N^2 , ϵ layers



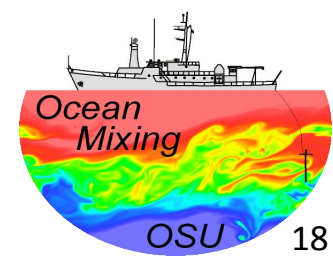
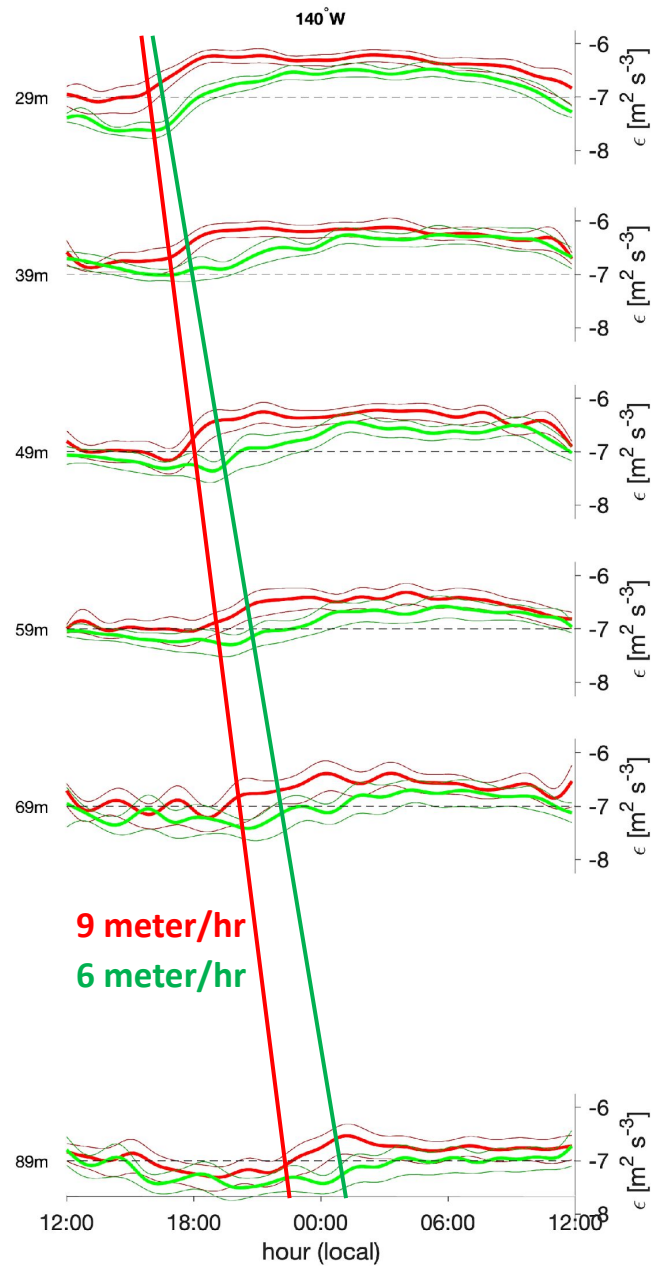
Smyth et al, 2013 (*JPO*)

computed from 8 day record

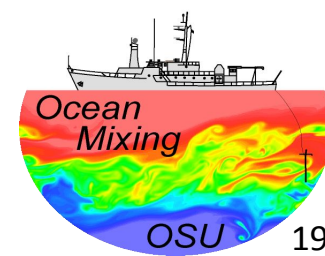
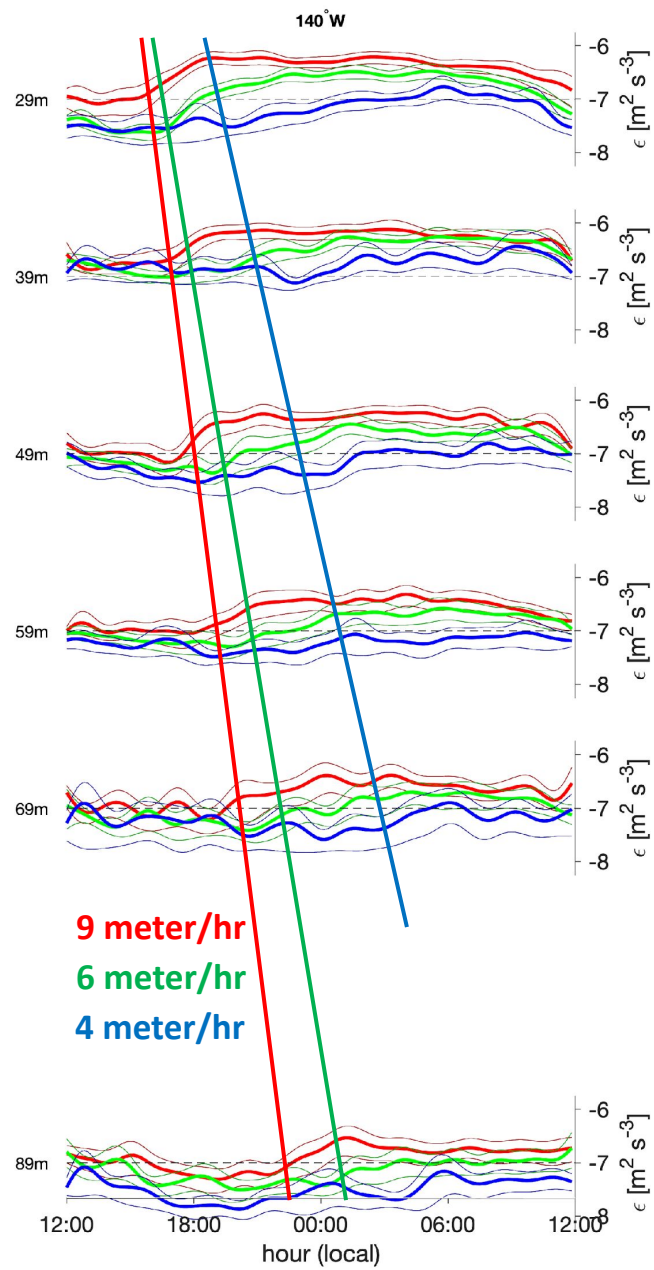
computed from 15 year record



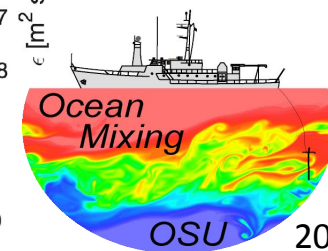
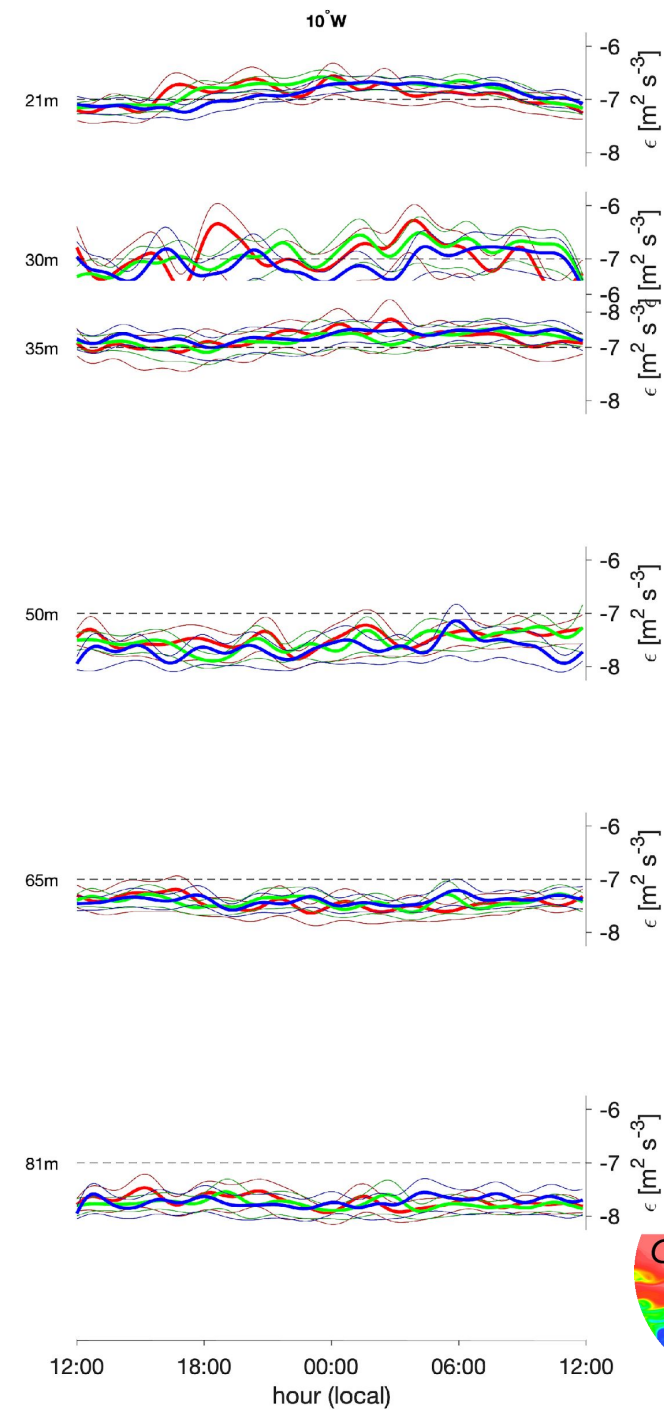
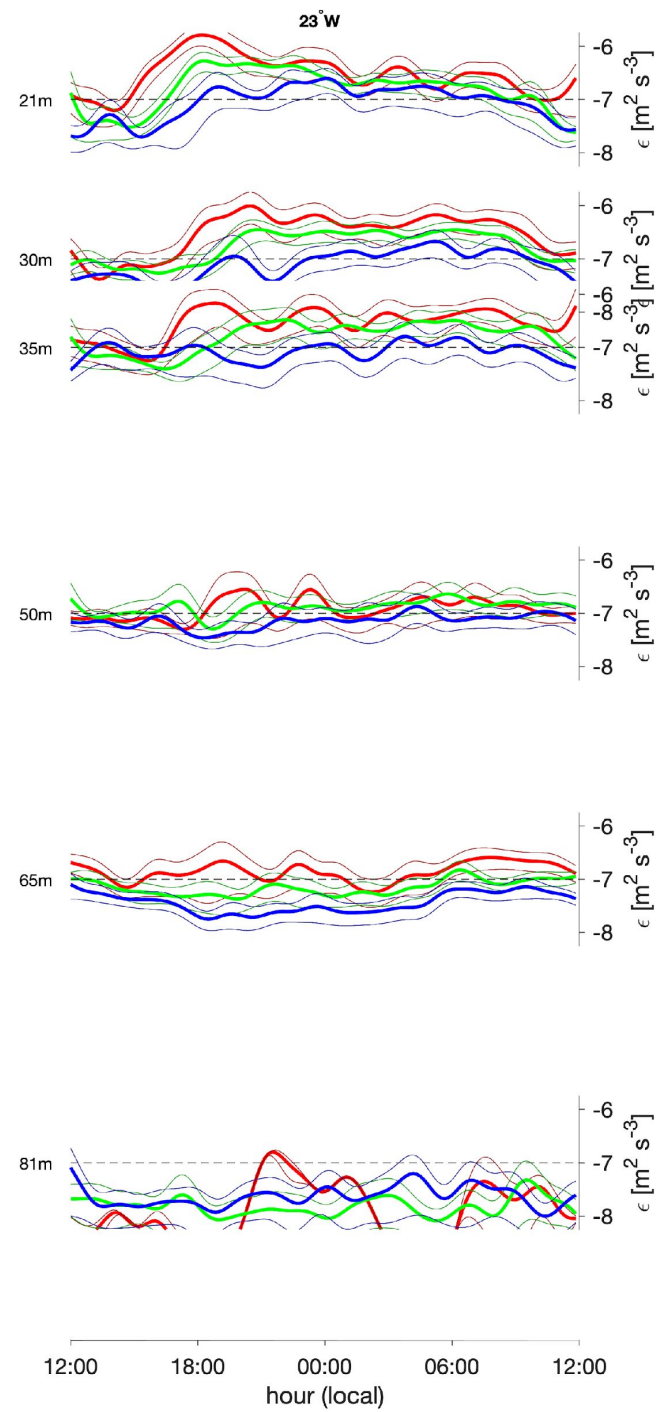
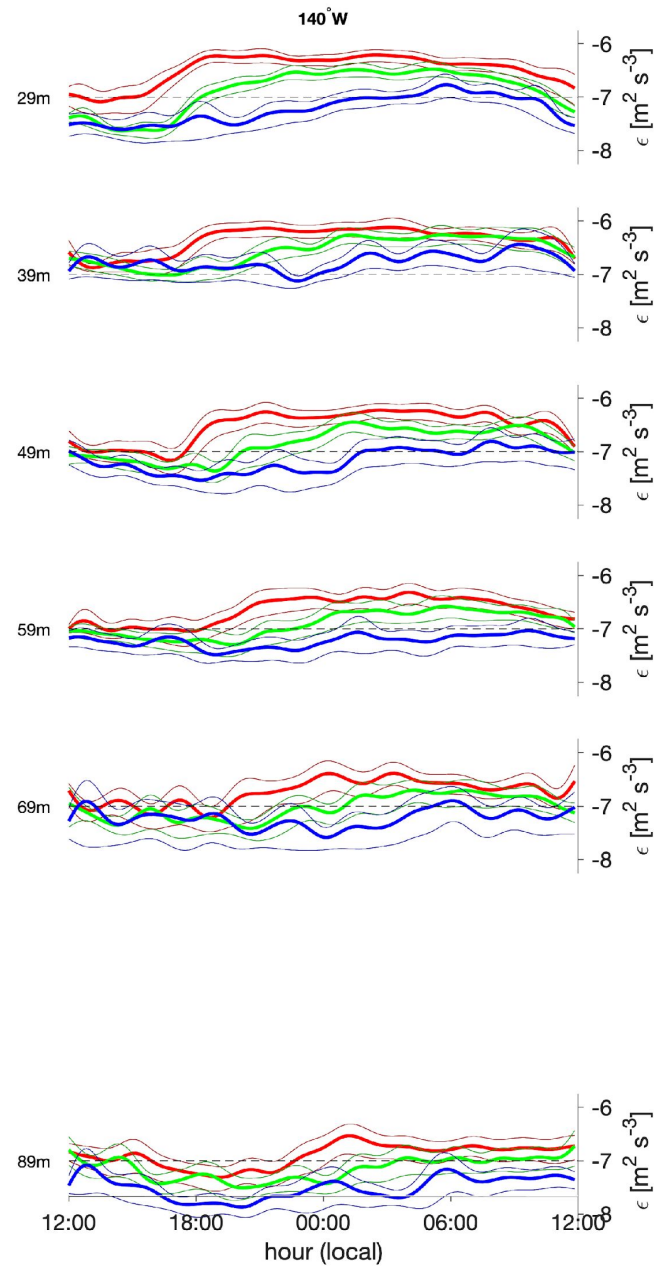
$\tau > 0.075 \text{ Pa}$
 $0.04 \text{ Pa} < \tau < 0.075 \text{ Pa}$



$\tau > 0.075 \text{ Pa}$
 $0.04 \text{ Pa} < \tau < 0.075 \text{ Pa}$
 $\tau < 0.04 \text{ Pa}$



$\tau > 0.075$ Pa
 $0.04 \text{ Pa} < \tau < 0.075$ Pa
 $\tau < 0.04$ Pa



Seasonal variability of turbulence and heating

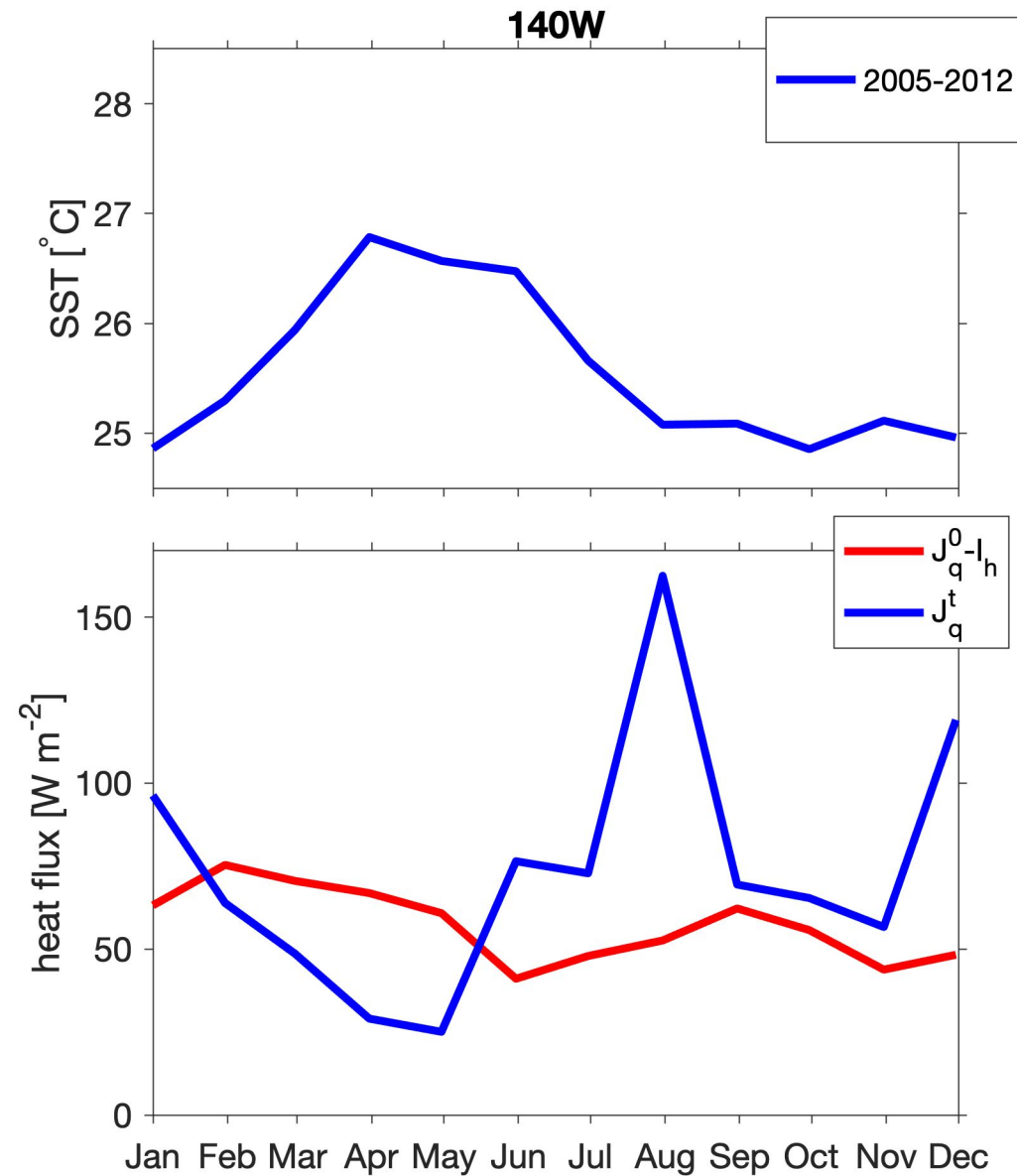
$$\frac{\partial T_{mld}}{\partial t} = \frac{1}{\rho C_p} \frac{\partial J_q}{\partial z}$$

$$\approx \frac{1}{\rho C_p} (J_q^0 - I_h - J_q^t)/h$$

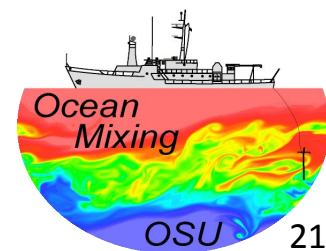
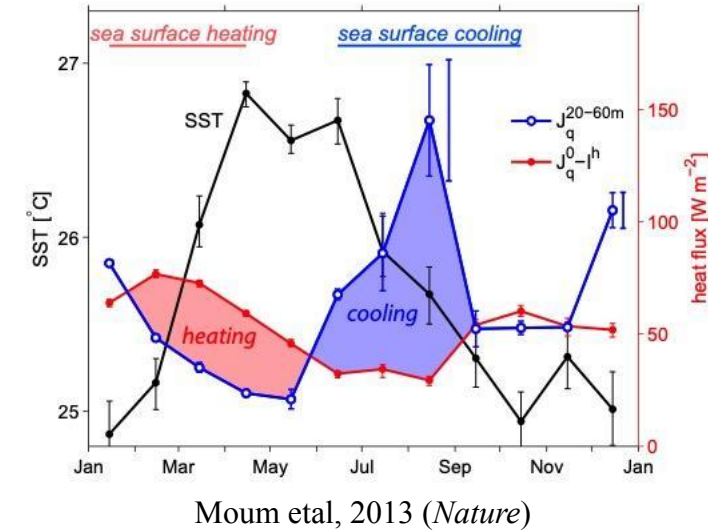
J_q^0 - net surface heat flux

I_h - short wave penetration through mixed layer base h

J_q^t - turbulent heat flux at h



Seasonal sea surface cooling in the equatorial Pacific cold tongue controlled by ocean mixing



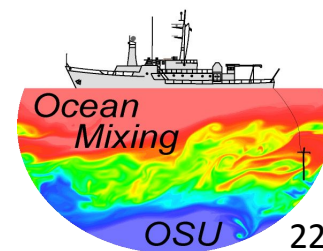
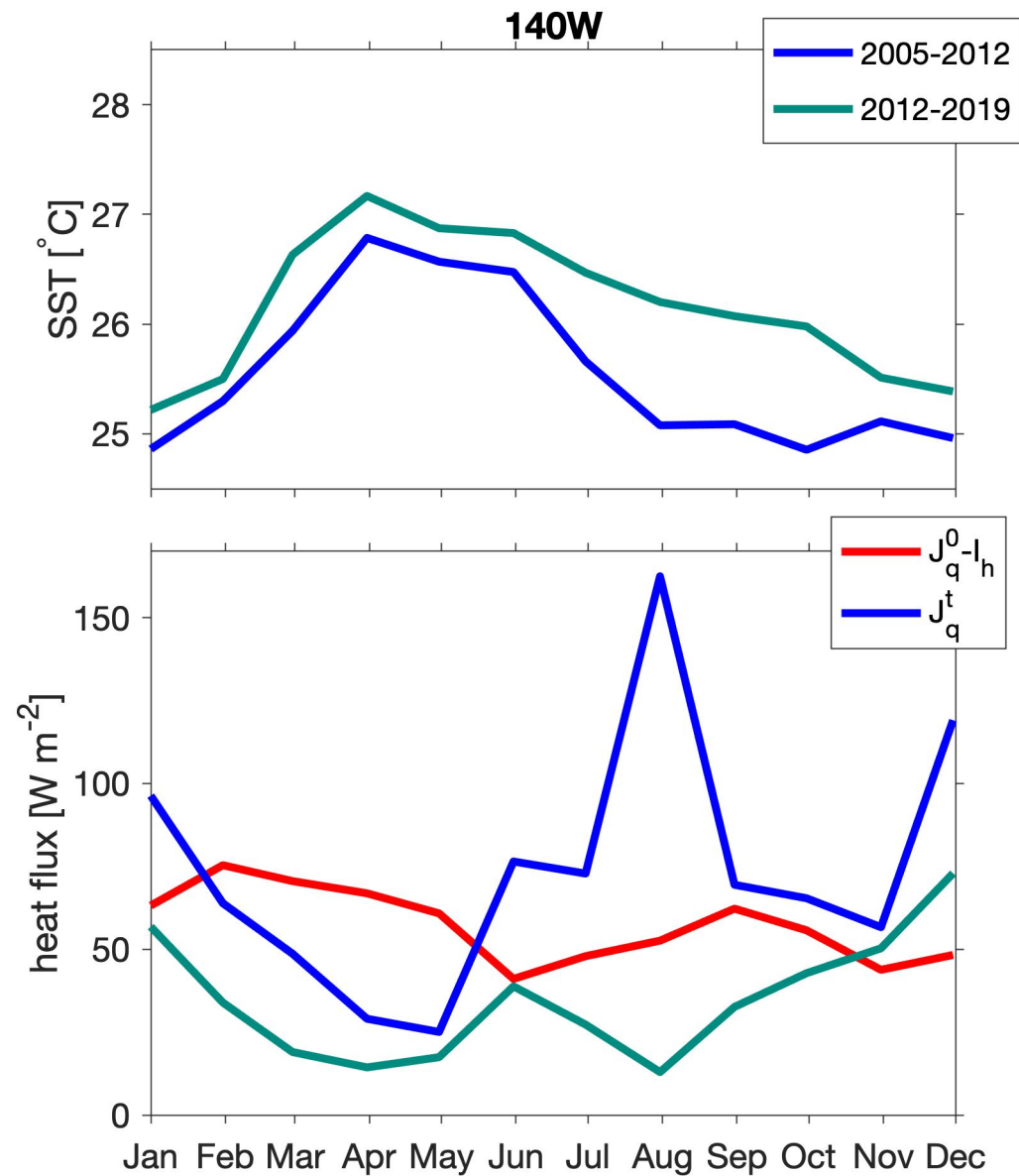
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$$\approx \frac{1}{\rho C_p} (J_q^0 - I_h - J_q^t) / h$$

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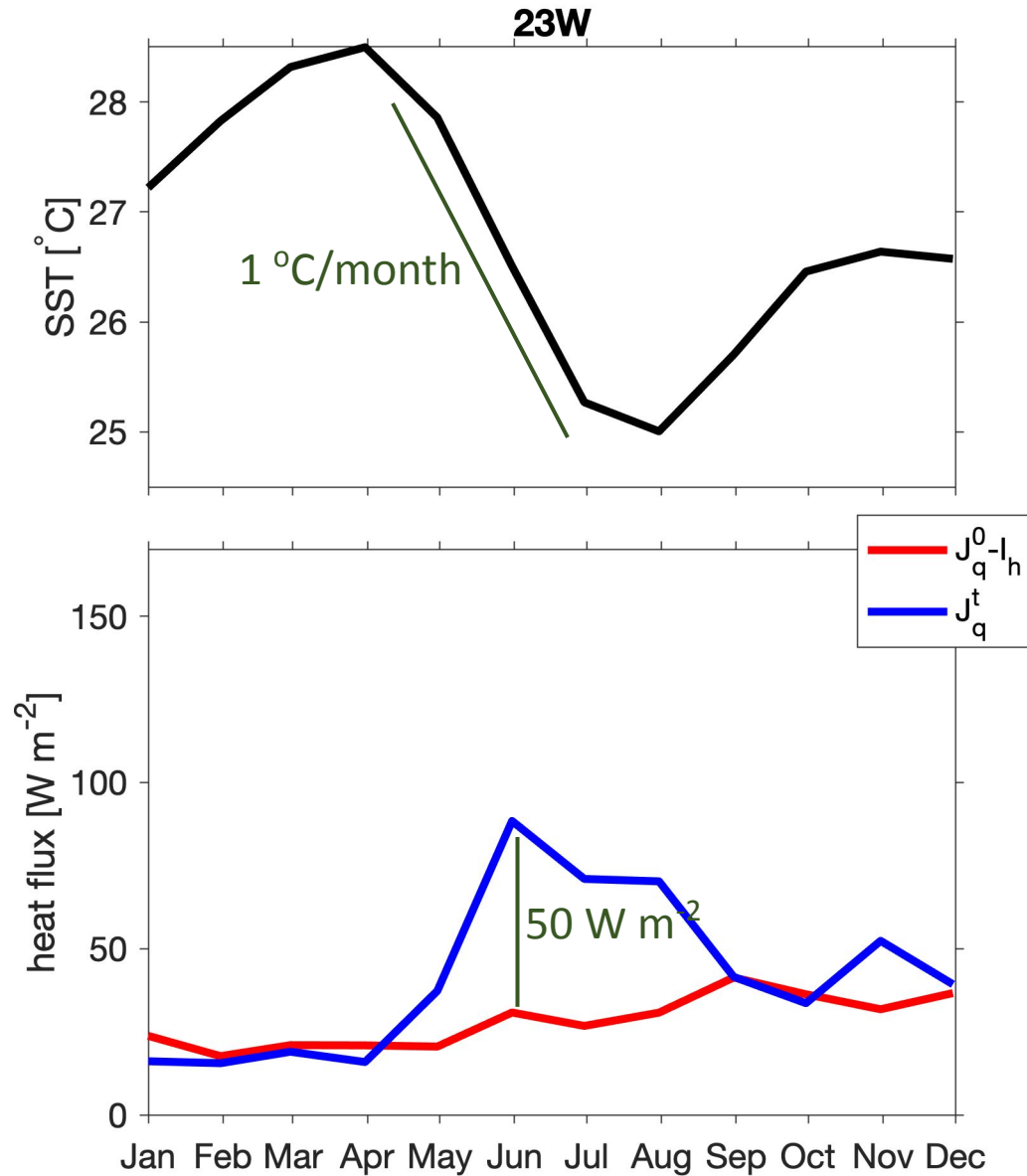
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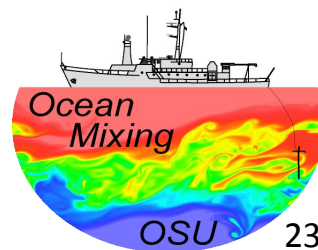


$$\frac{\partial T}{\partial t} = \frac{1}{\rho C_p} \frac{\partial J_q^t}{\partial z}$$

$$\approx \frac{1}{4 \times 10^6} \frac{50}{21}$$

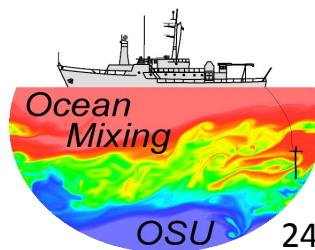
→ 1.5 °C/month

0, 110W
no hourly flux values



Summary

- deep diurnal mixing cycle present at 10W, 23W – similar to deep cycle at 140W
- wind-dependent phasing with depth
- near-surface magnitudes are similar at all 3 sites
- depth dependencies unclear in ACT because mixing is confined to thinner, shallow layers
- turbulent heat transport through ML at 23W > required for seasonal SST changes
- ADCP marginally measures MI layer at 10W, 23W



Proposed New Deployments

shallower sensors
reduced vertical spacing
(nominal 10 m separations)

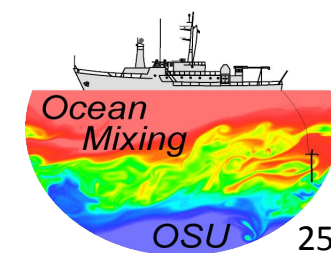
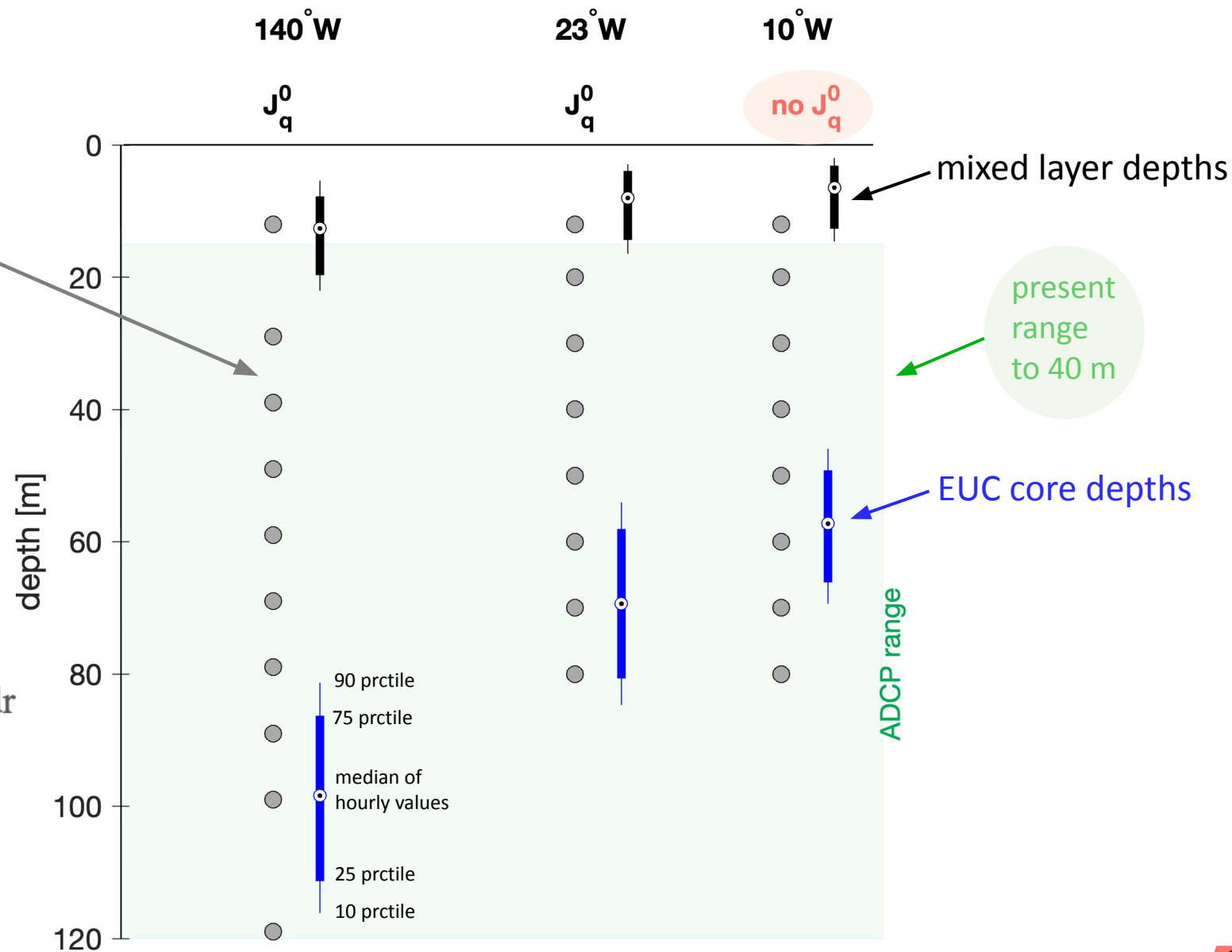
- more shallow sensors
- better resolution of $\frac{\partial J_q^t}{\partial z}$

Also, some issues to consider ?

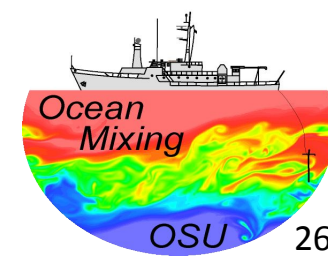
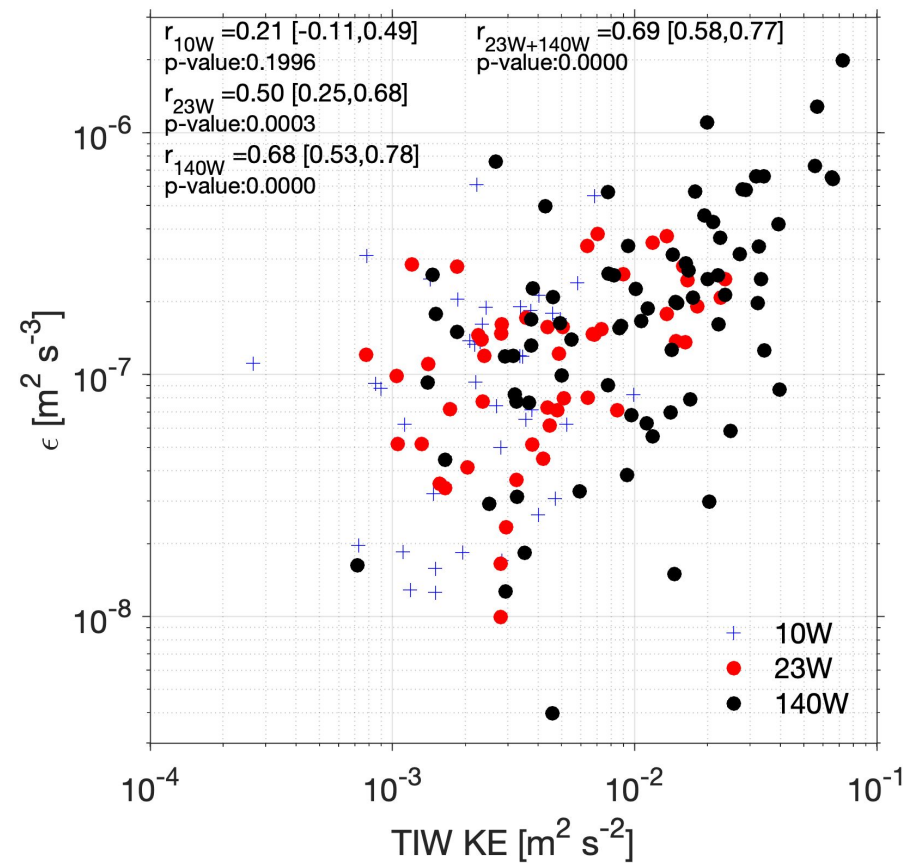
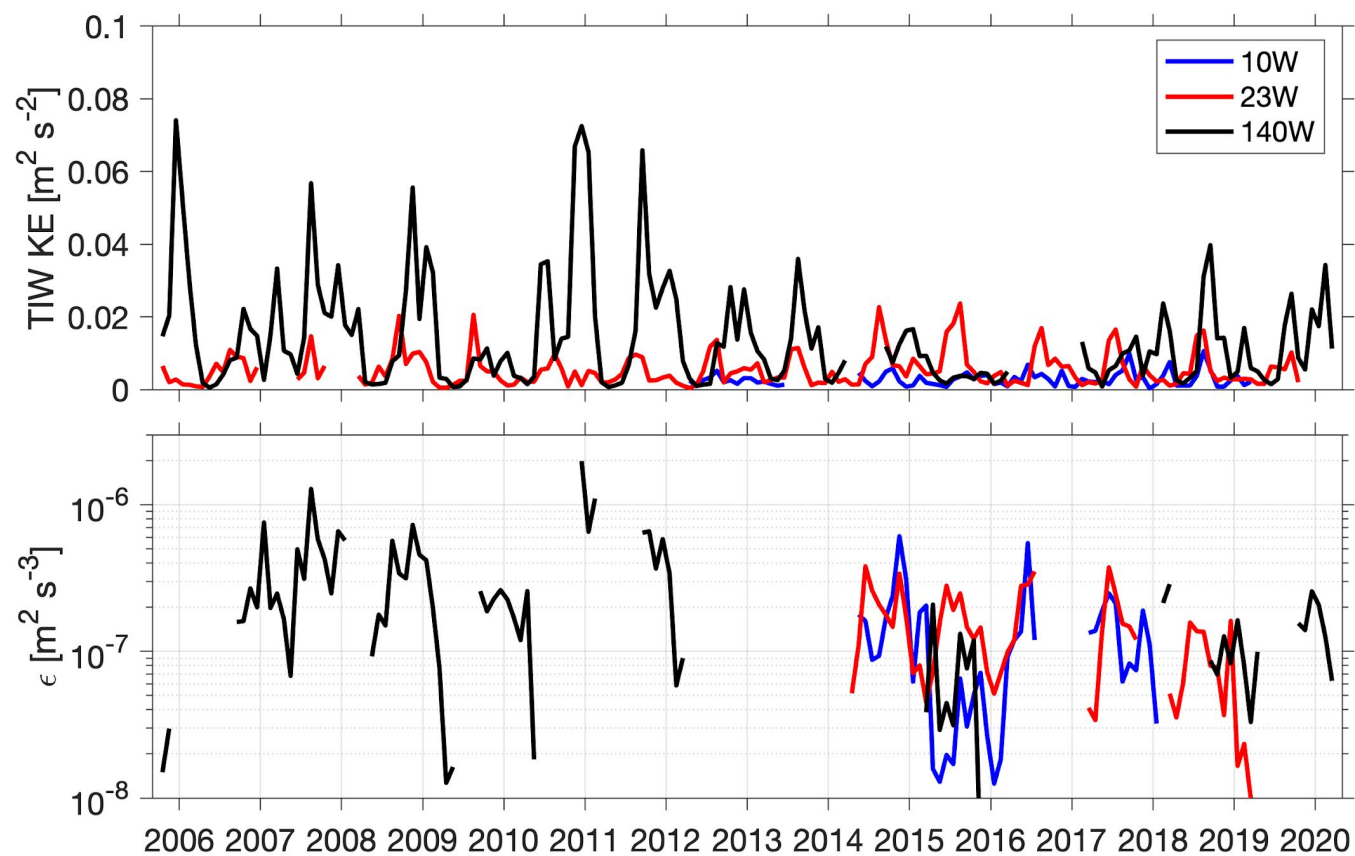
are there ways to get better near-surface velocities?

- side lobe suppression / bigger xdr
- reduce beam angle (20°)
- move ADCPs shallower in east

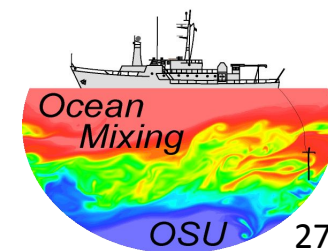
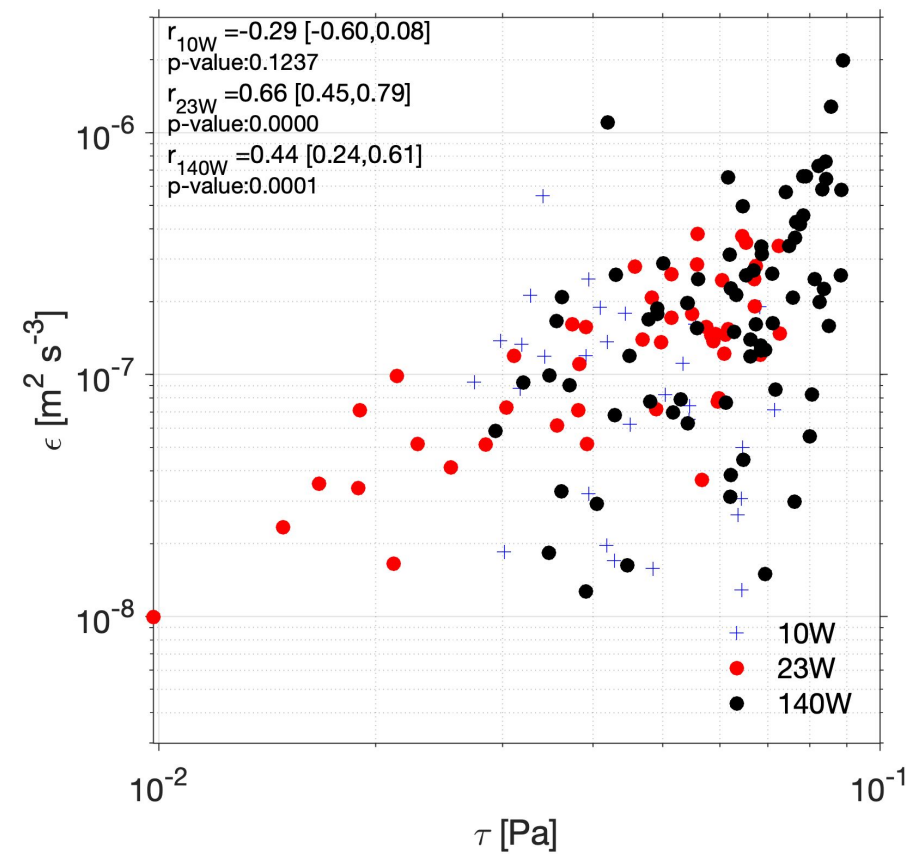
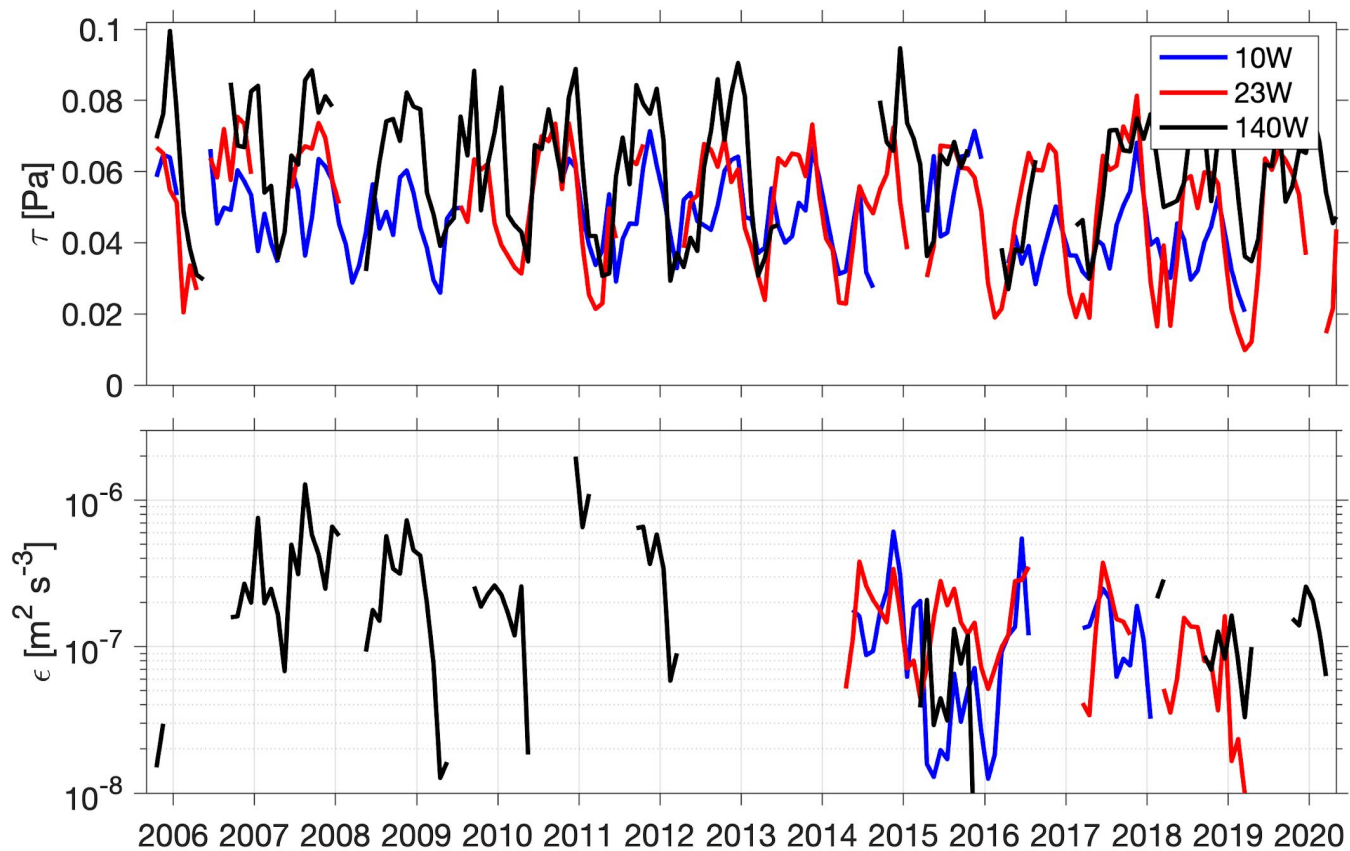
can 0, 10W be made a flux site?



TIW KE and Mixing (ϵ)



Wind stress (τ) and Mixing (ϵ)



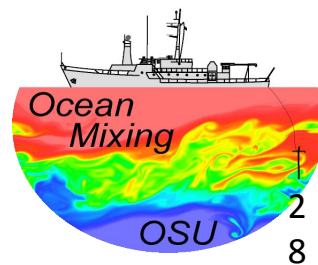
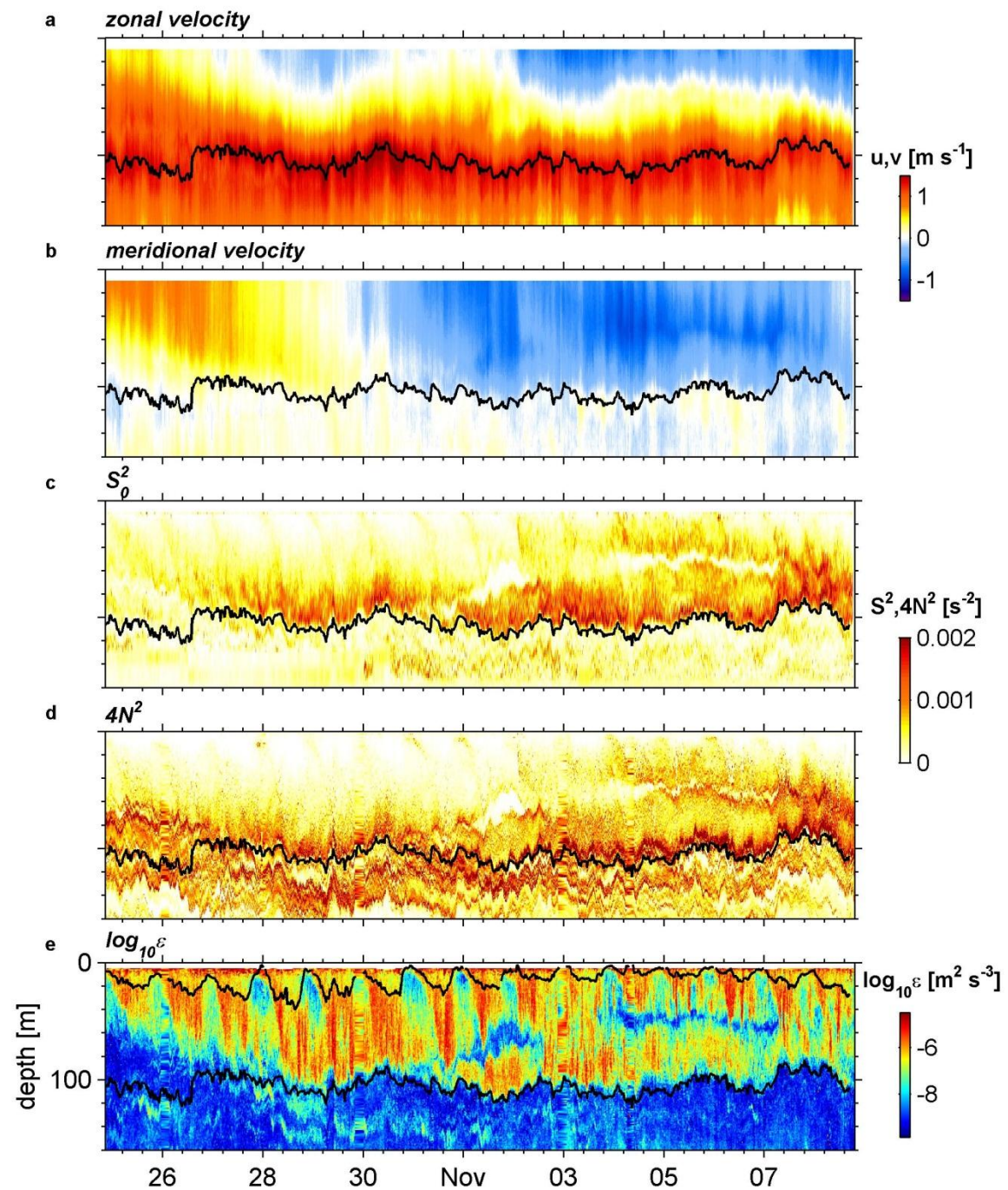
Patterns of Mixing 2008

0, 140W

an example of profiling time series from ship – 16 days

mixing during passage of a tropical instability wave

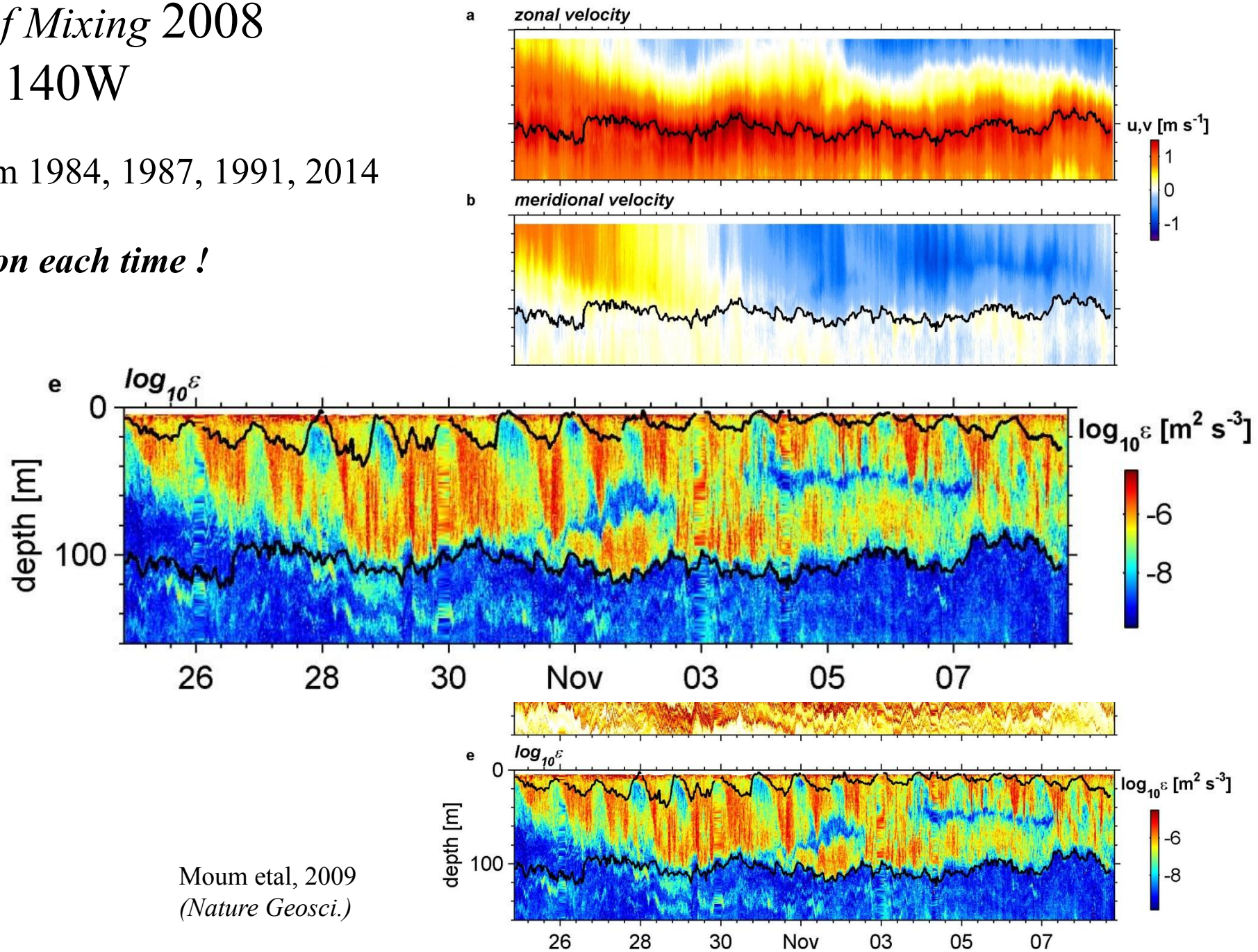
Moum et al, 2009
(*Nature Geosci.*)



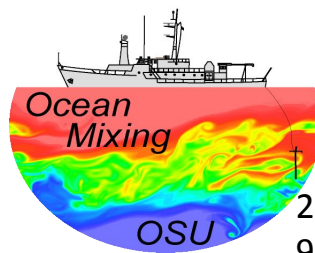
Patterns of Mixing 2008 0, 140W

unique from 1984, 1987, 1991, 2014

a new lesson each time !



Moum et al, 2009
(*Nature Geosci.*)



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$$\approx \frac{1}{\rho C_p} (J_q^0 - I_h - J_q^t) / h$$

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