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1. Overview

Accurate accounting of surface CO₂ flux is crucial for policymaking toward mitigation of greenhouse gases emissions for limiting global warming below 1.5 or 2.0 °C by 2100. This ambition requires accurate and spatiotemporally dense coverage of satellite CO₂ monitoring and application of top-down inversion system. Our study reports (probable) systematic biases in OCO-2 satellite retrievals (version 10r) through comparison with highly accurate aircraft CO₂ measurement in an inversion system.

2. Data

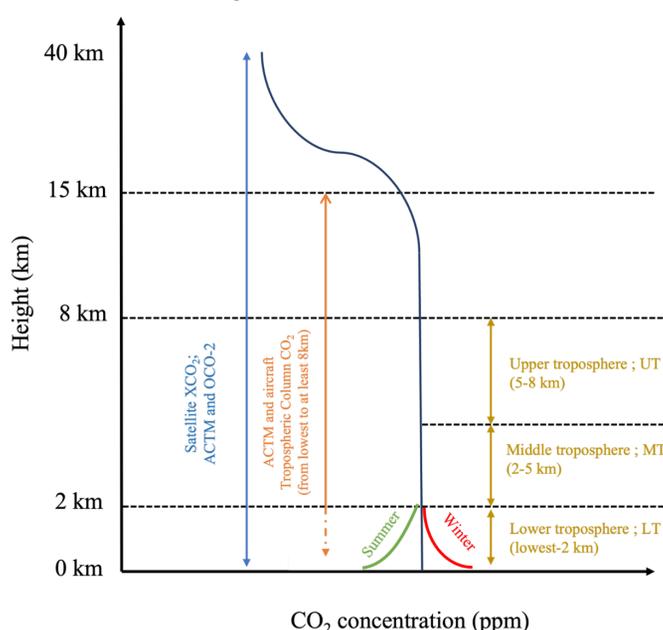
CO ₂ dataset	Data source
1. Aircraft vertical profile (VP) of CO ₂	1. Atmospheric Tomography Mission (ATom) and JAL/NIES CONTRAIL from obspack_co2_1_GLOBALVIEWplus_v6.1 (Schuldt et al., 2021) 2. Brazilian Amazon VP of CO ₂ from four aircraft sites (SAN, ALF, RBA, TEF) (Gatti et al., 2021)
2. Surface CO ₂	1. Global in-situ sites from obspack_co2_1_GLOBALVIEWplus_v6.1 2. WMO WDCGG (World Data Centre for Greenhouse Gases)
3. Column CO ₂ (XCO ₂)	1. NASA/JPL OCO-2 Level-2, version 10r (O'Dell et al.)

3. Transport model

Transport Model	MIROC4-ACTM
Prior	$FG = CO_{2,ff}(\text{GridFED}) + CO_{2,ld}(\text{CASA-3hr}) + CO_{2,ocn}(\text{Taka-Ocn})$
ACTM XCO ₂	$XCO_2^{ACTM} = \sum_i CO_2^{priori} \cdot dp_i + \sum_i A_i \cdot dp_i \left(\sum_i CO_2^{ACTM} - \sum_i CO_2^{Priori} \right)$

Details on forward and inversion can be found in Chandra et al. (2022); Das et al. (2022), whereas XCO_2^{ACTM} calculation is in Patra et al. (2017)

4. Data analysis



Aircraft vertical profiles of CO₂ is subdivided into three tropospheric layers LT, MT, UT, and tropospheric Column or Aircraft XCO₂

Mismatch of ACTM against OCO-2 XCO₂ and aircraft CO₂ at vertical layers are evaluated considering airmass/pressure layers of partial columns (Fig. 1)

Fig. 1: Vertical profile of CO₂ from aircraft and ACTM for total column calculation, and the definitions of tropospheric layers LT (lowest-2 km), MT (2-5 km) and UT (5-8 km) are denoted with different colors.

5. Results

5.1 MIROC4-ACTM comparison with OCO-2 and surface CO₂

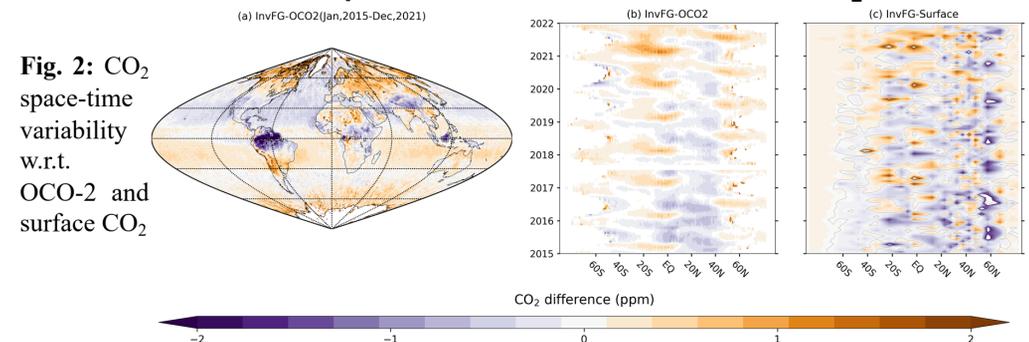


Fig. 2: CO₂ space-time variability w.r.t. OCO-2 and surface CO₂

Mismatch between XCO_2^{ACTM} and OCO-2 are observed over land than ocean likely due to uncertainty in prior flux in land biosphere than the oceanic CO₂ exchange [Fig. 2a]

XCO_2^{ACTM} shows seasonally repeating difference with respect to OCO-2 observed XCO₂ [Fig. 2b], as opposed to surface CO₂ [Fig. 2c]

5.2 CO₂^{ACTM} difference against XCO₂ from Aircraft and OCO-2

5.2.1 North America, Southern Ocean, Pacific, and Atlantic

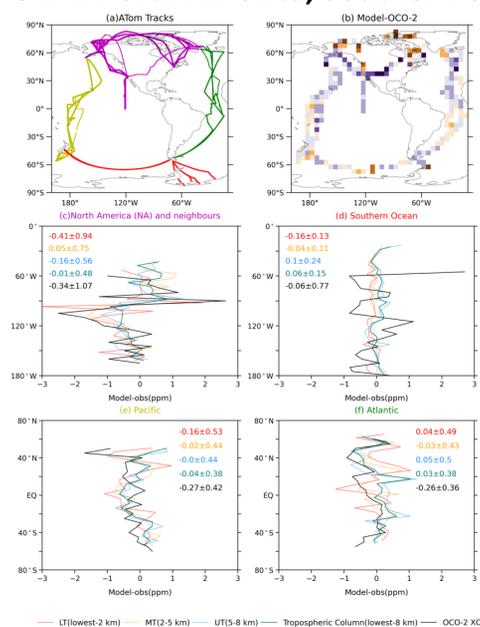


Fig. 3: XCO_2^{ACTM} difference with OCO-2 and aircraft CO₂ along several ATom aircraft campaign tracks.

Largest mismatch (mean and variability) of XCO_2^{ACTM} against OCO-2/Aircraft XCO₂ observed over North America and neighbours in LT are due to uncertainty in prior land flux.

Best match is observed over remote tropospheric air in Southern Ocean, also in Pacific and Atlantic Ocean sectors due to less impact of land airmass.

5.2.2 Amazon

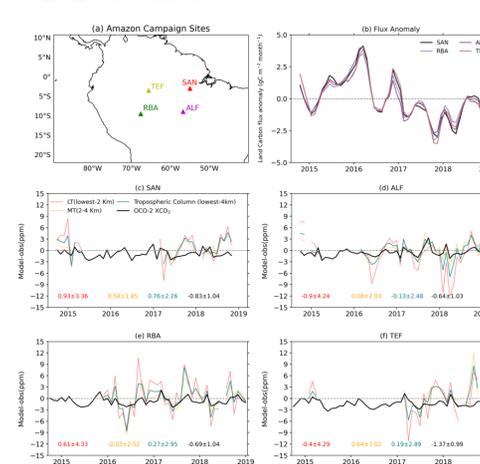


Fig. 4: CO₂ differences at the Amazon aircraft sites

Largest mismatch of ±12 ppm is found in LT at ALF, RBA and TEF specific aircraft measurement site. Also, a systematic difference against OCO-2 XCO₂ is evident, but mostly less than 1 ppm.

Differences also arise due to illrepresentation of sub-grid scale process in MIROC4-ACTM due to coarse resolution 2.8°×2.8°, which are further evident at the CONTRAIL airport sites

5.3 Conclusion

Systematic retrieval bias is evident in OCO-2 across time period with pronounced effect in southern hemisphere ocean region.

Most model-CO₂ differences exist in LT because of uncertainty in prior flux, and coarse model resolution. MIROC4-ACTM simulated CO₂ at MT, UT well by capturing the large scale dynamical transport.